TECHNICAL MEMORANDUM

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On the effect of model resolution on numerical simulation of blocking

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ABSTRACT

A numerical case study of blocking is discussed briefly for which model resolution seems to play a major role in the success of the simulation. It is argued that the effect of the initial conditions is somewhat less important than normally considered. The model used for all the experiments is the ECMWF spectral model (Baede, Jarraud and Cubasch, 1979).

1. INTRODUCTION

The horizontal scale of atmospheric blocking patterns is large and their dynamics has, with few exceptions (e.g. Green, 1977 and Bengtsson, 1981), always been investigated in terms of planetary ultra-long waves. In trying to model numerically this phenomenon, therefore, model resolution would perhaps be considered to be a non-critical modelling characteristic. If, however, as pointed out in different contexts by these authors, the interactions between smaller scale transient eddies and the quasi-stationary system are essential for the maintenance of the blocking pattern, the partial absence of such smaller eddies would strongly decrease the ability of a model to represent this phenomenon. A satisfactory representation of these eddies in terms of resolution would then become of fundamental importance. The purpose of this work is to report a numerical case study of blocking that showed a high degree of sensitivity to model resolution, so as to suggest that the scales of motion unresolved by the low resolution run but resolved by the high one played a physically important role in the onset and maintenance of the blocking circulation.

The model used was the ECMWF spectral model (Baede, Jarraud and Cubasch, 1979) at the two triangular truncations T40 (low resolution) and T63 (high resolution). The scales of motion resolved explicitly by the T63 model but not by the T40 could be thought as spanning the range between 300 and 500 km, expressing the concept of "scale" in terms of half wavelength. The parameterization of physical effects (radiation, condensation and effects of scales of motion beyond the model truncation, (see Tiedtke, Geleyn, Hollingsworth and Louis, 1979) was exactly the same in the two runs and reflects more or less the status of the "physics" of the ECMWF operational grid-point model as at August 1981.

2. THE CASE STUDY

The case selected took place during December 1978 and partially overlaps with the period covered by the WGNE Case Study No.1 (see JSC, 1981), since we will discuss ten-day integrations starting from December 17, 1978 at 12GMT. This development initiated the blocked-type circulation which characterized most of the winter period in 1978/79. The initial data used are the ECMWF FGGE global initialized (non-linear normal mode initialization, see Williamson and Temperton, 1981) analysis. Fig. 1 shows the 500 mb and 1000 mb height maps for the Northern Hemisphere for the initial data (day 0).

Figs. 2a to 5a show that, by day 5 of the integration period, a very large anticyclone had established itself over the North Atlantic, diverting completely the flow of the westerlies. This large scale situation intensified at first and then was maintained, with little or no substantial change, for another week. By day 10 of the forecast period, the blocking pattern had invaded the whole of the Atlantic region, from 30 degrees north to the North Pole (Fig. 5a).

Figs. 2 to 5 show the different synoptic behaviour that the two different resolution runs produced compared with the verifying FGGE analyses. Figures on the left, (a), show the verifying analysis; centre figures, (b), show the high resolution T63 run; figures on the right, (c), describe the low resolution T40 run. 500 mb height maps are above 1000 mb height maps. It is clear from the figures that the T63 forecast succeeds in capturing accurately most of the main features of the blocking event. The day 5 forecasts predicts correctly the build-up of the blocking, down to several synoptic details of the flow; the amplitude is, however, underestimated. The day 8 forecast is characterized by a marked erroneous eastward displacement of the blocking feature. It, nevertheless, still shows a blocking pattern of realistic intensity with a very good representation of the split of the westerly jet at 500 mb. The day 10 forecast, although showing a much earlier weakening of the block, still contains some synoptic value; note, for example, the correct split of the block into a double feature, with two block-like patterns side by side, both at 500 and 1000 mb. On the other hand, the lower resolution run fails completely to accomplish the onset of the block, despite a very reasonable short range (e.g. day 3) forecast. By day 8 the two forecasts have completely diverged synoptically from each other, with the T40 run having erroneously drifted towards a much more zonal state.

It is interesting, at this point, to note that another low resolution run initiated from initial conditions of 12 hours later (1978-12-17-12GMT, experiment T40B) shows little improvement (see Fig. 6) in describing the development in the Atlantic, ending up by day 8 (actually day 7.5 for this experiment) with a block-like pattern in the wrong hemispheric quadrant. This would suggest that, in this particular case study, the model is showing to be as sensitive to the resolution as it is to the specification of the initial conditions, if not more. It also shows, however, that the low resolution model is capable of producing blocks, since the one developed at day 7.5 of the T40B forecast (Fig. 6c) will last up to the end of the 10 day integration (not shown).

Comparing the T63 and T40 objective scores (see Fig. 7), it can be noticed that the two experiments start diverging from each other only after day 4.5, except, of course, for the shorter waves upon which the impact of resolution is more immediate (day 1.5). After day 4, however, the main difference (to the advantage of the high resolution run) is to be found in the ultra-long wave components. For a more complete spectral resolution comparison study of this type, the reader is referred to Jarraud, Girard and Cubasch (1981).

The time evolution of the kinetic energy (Fig. 8) also shows that there is a sharp increase of KE in the wavenumber band 4 to 9 during the first two days of the integration, followed by a decrease taking place in parallel with a rapid growth of the KE in the wavenumber band 1 to 3. This might be interpreted in the following way: during a first stage (first two days) an intense development of a synoptic scale wave takes place which favours the onset of the block; subsequently, an exchange of kinetic energy from medium (4 to 9) to long and ultra-long (1 to 3) waves takes place in order to maintain the blocking situation. An adequate description of synoptic and subsynoptic scale waves (hence the sensitivity to resolution) would therefore be essential during both stages; see also Källén (1982).

Figure 9 shows the difference fields of 500 mb geopotential height between the T63 experiment and the T40 experiment days 2, 4, 6 and 8. It is here possible to see how the errors in the low resolution run originate from an underestimation of the development off the east coast of North America (Fig. 9a) and then grow to larger amplitudes. At day 4, a pronounced centre in the difference fields located on the North American continent, in the lee of the Rocky Mountains (Fig. 6b), suggests that the T40 run might not give an adequate description of the mountain effects which cause successive downstream developments of troughs (Simmons and Hoskins, 1979); this might, in turn, contribute to the onset and maintenance of the blocking (Kalnay-Rivas and Merkine, 1981).

3. CONCLUSIONS

The sensitivity to resolution that the ECMWF spectral model has shown in this Atlantic blocking case study indicates that the cumulative effects of scales of motion unresolved by a spectral triangular 40 truncation (but resolved by a T63 one) can have a crucial influence on the onset of such an important dynamical structure as blocking. These effects are most likely due to non-linear interactions between ultra-long waves and shorter waves and seem to affect the ability of the model to enter into a locally blocked state at the right time and in the right area. This confirms, on a particular blocking case study, general results previously obtained in a larger sample comparison, see Jarraud, Girard and Cubasch (1981). This, of course, has an immediate bearing on the importance of model resolution in numerical weather prediction, mainly on time scales greater than 2 to 3 days.

This case study would also seem to lend more support to those regarding blocking as a large scale circulation system supported vitally by comparatively smaller scale eddy disturbances than it would to those regarding it purely as an "almost inescapable" state for both the real atmosphere and hightly truncated models.

These two points of view need not, however, be mutually incompatible (cf. Källén, 1982). When one moves from a highly truncated model to a realistic GCM (and, possibly, to the real atmosphere), the presence of smaller scale eddies (and of smaller scale forcings) could, on one hand, change substantially the "attraction" properties and the number and position of the possible quasi-steady states in the phase space of the system. Smaller eddies could also, on the other hand, contribute locally to the creation of conditions for the transition from one local quasi-steady state into another (Malguzzi and Speranza, 1981) and, moreover, contribute to the stability (and "steadiness") of a blocked-type state.

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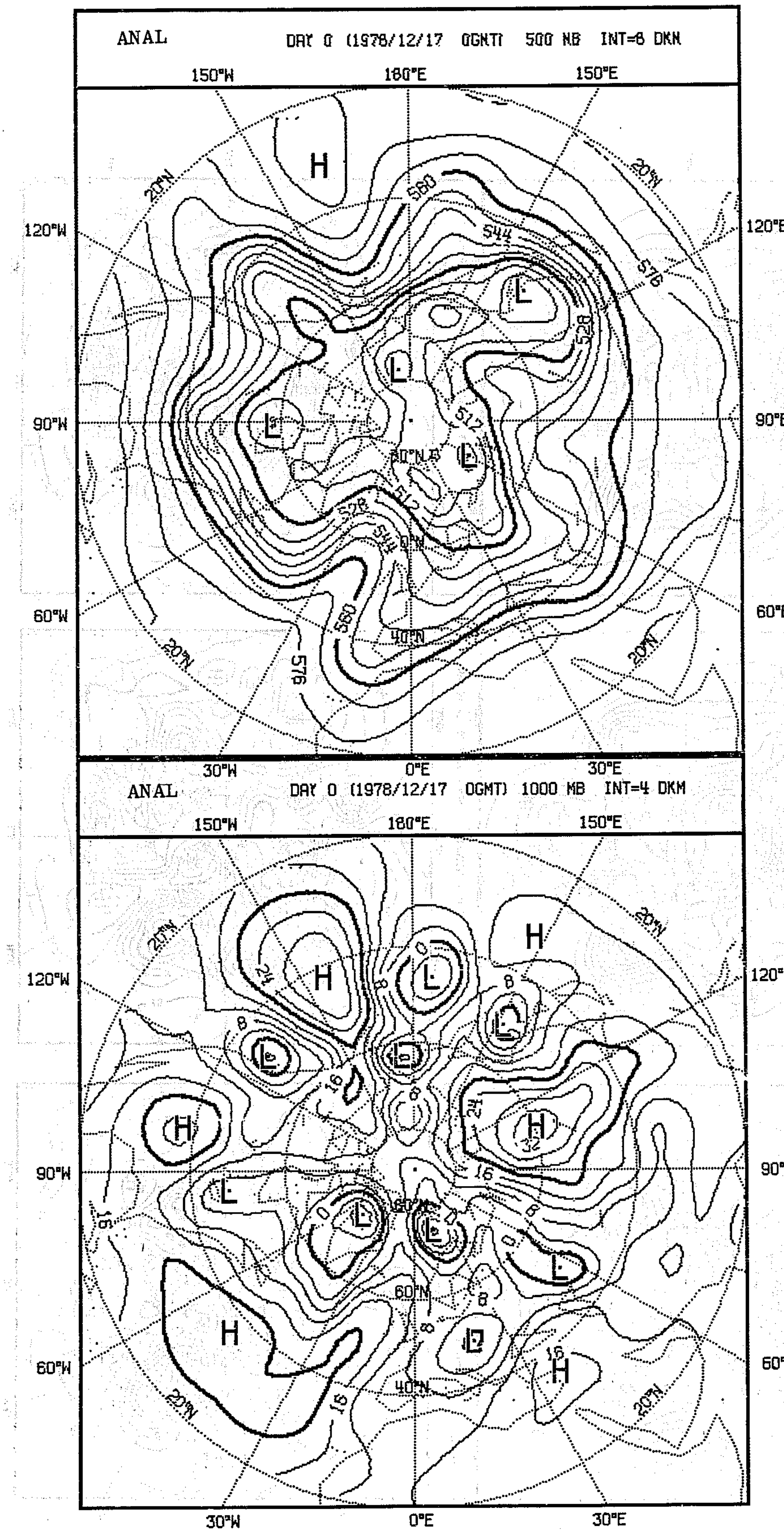


Fig. 1 FGGE analysis for 00 GMT 17.12.1978. Upper, 500 mb geopotential height; lower, 1000 mb geopotential height. Initial conditions used for the two numerical experiments of different resolution.

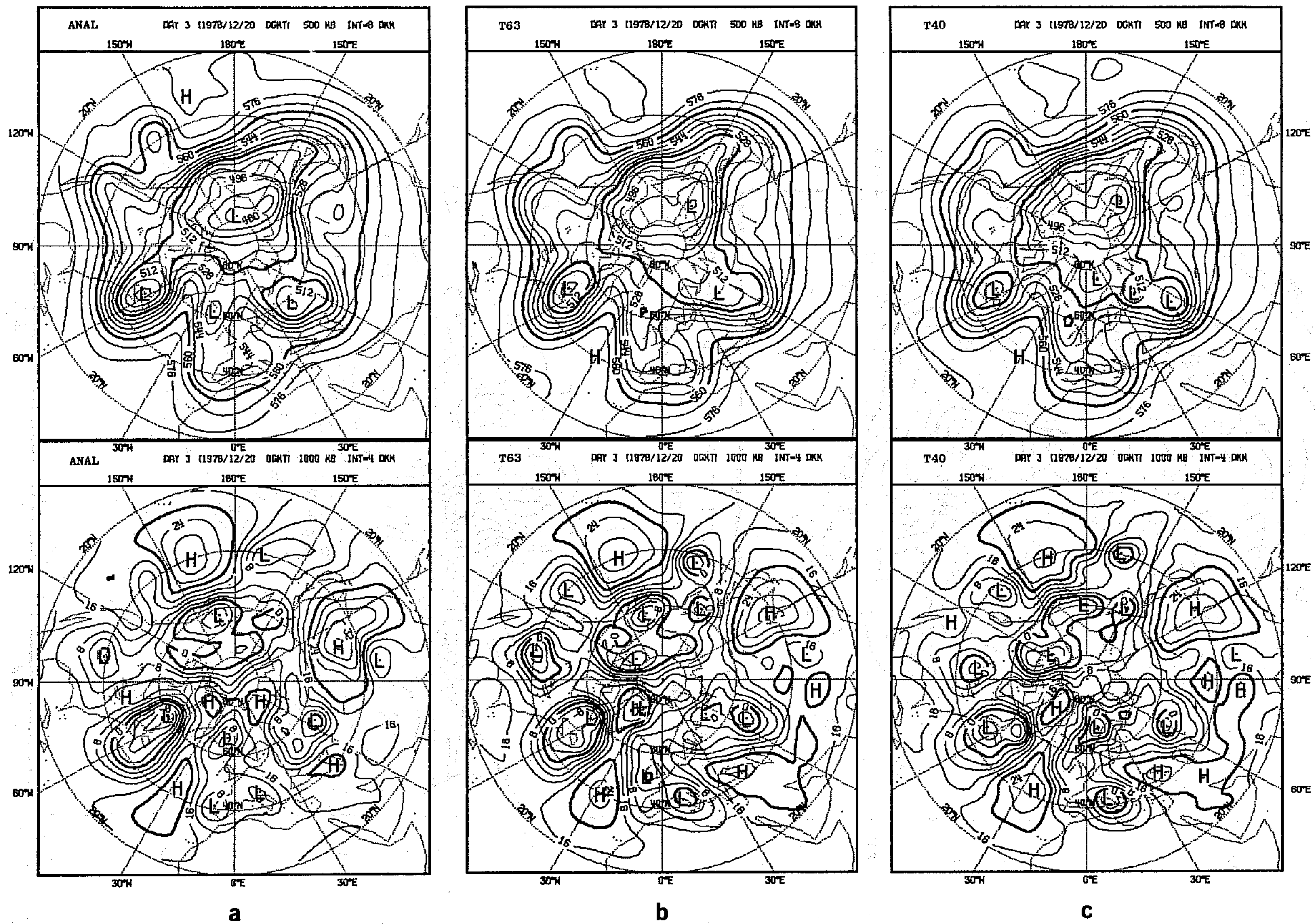
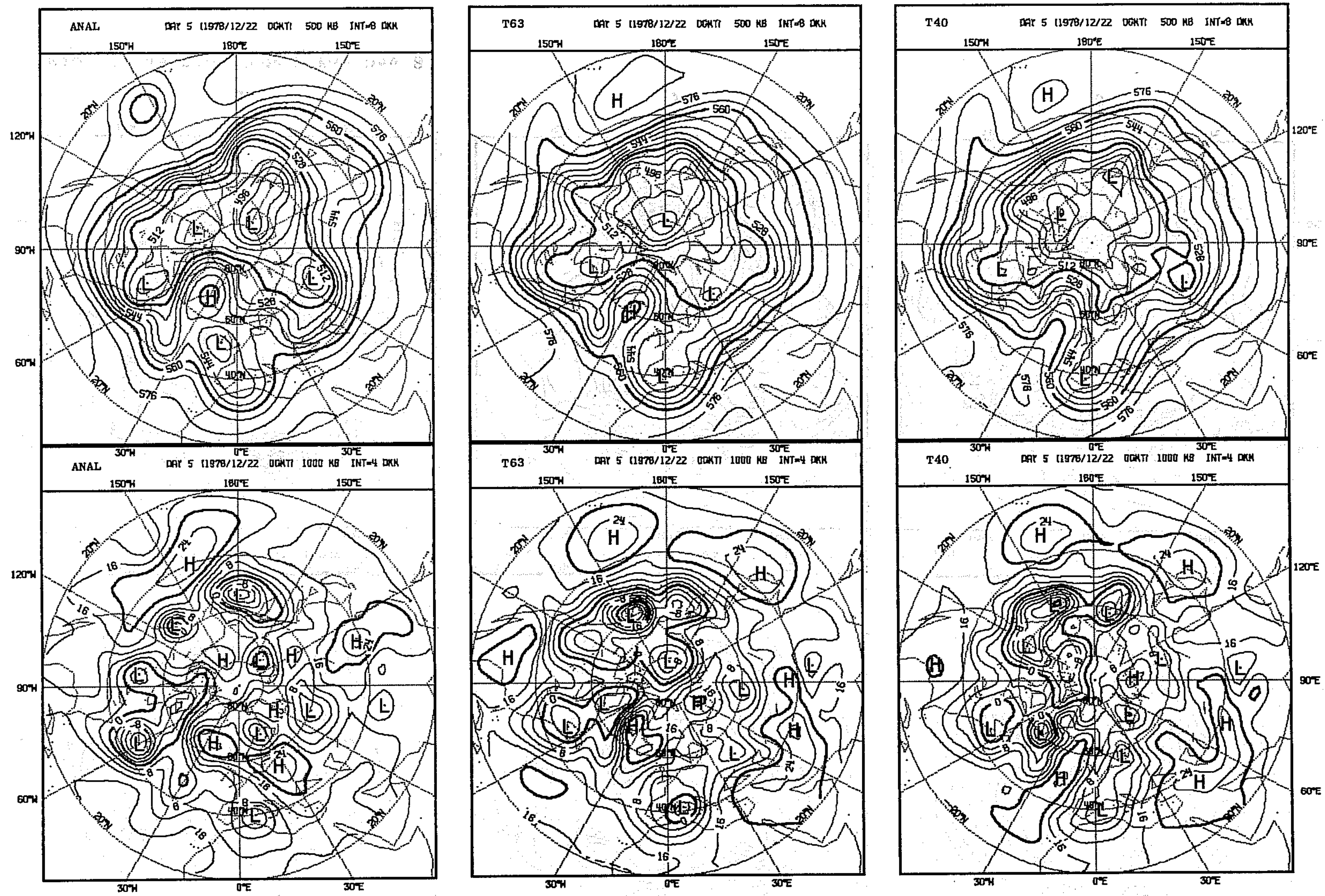
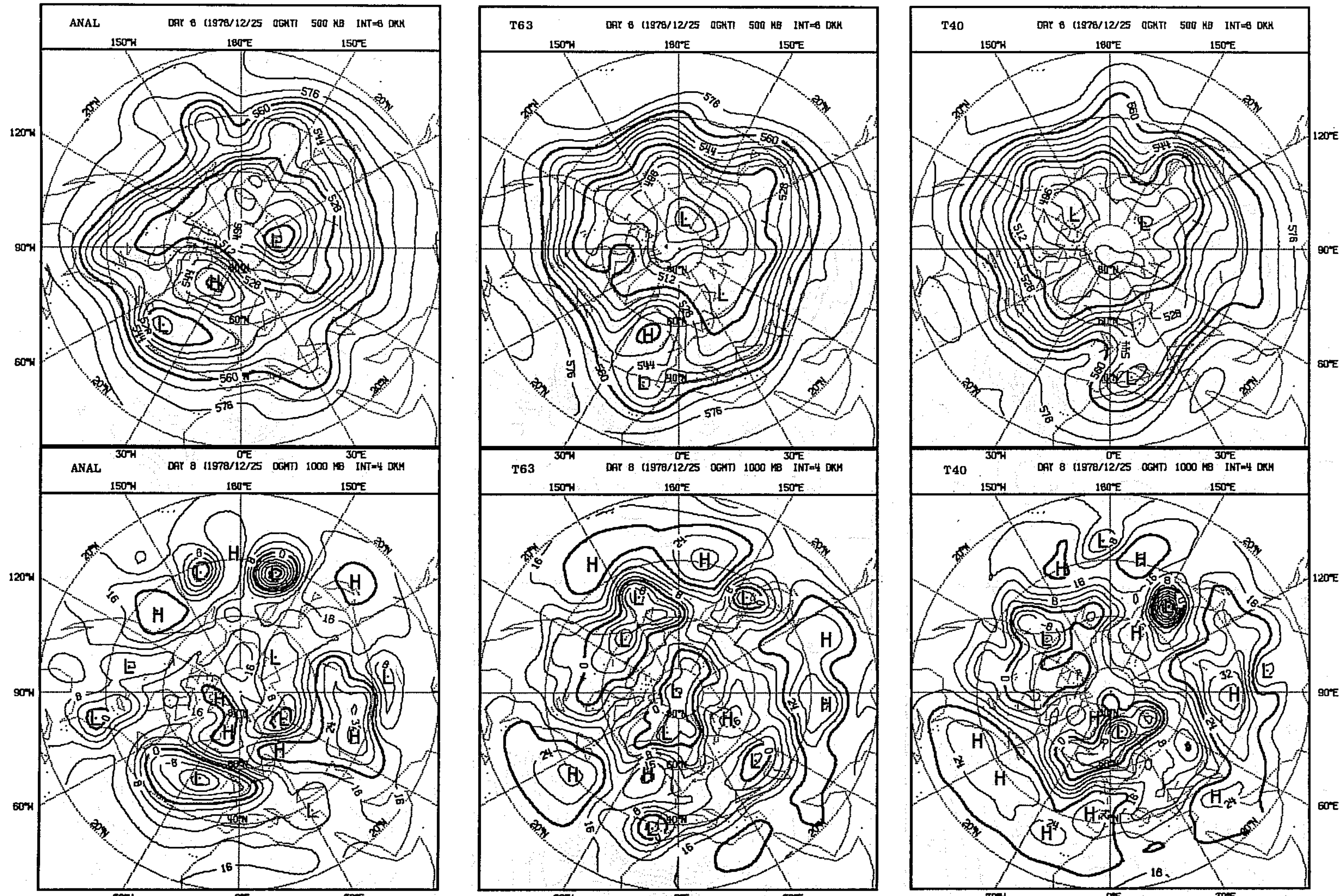


Fig. 2 Day 3 of the experiments. Left a) is verifying analysis; centre b) is the high resolution (T63) experiment; right c) is the low resolution (T40) experiment. Upper and lower as Fig. 1.



a b c

Fig. 3 As Fig. 2 but for Day 5.



a b c

Fig. 4 As Fig. 2 but for Day 8.

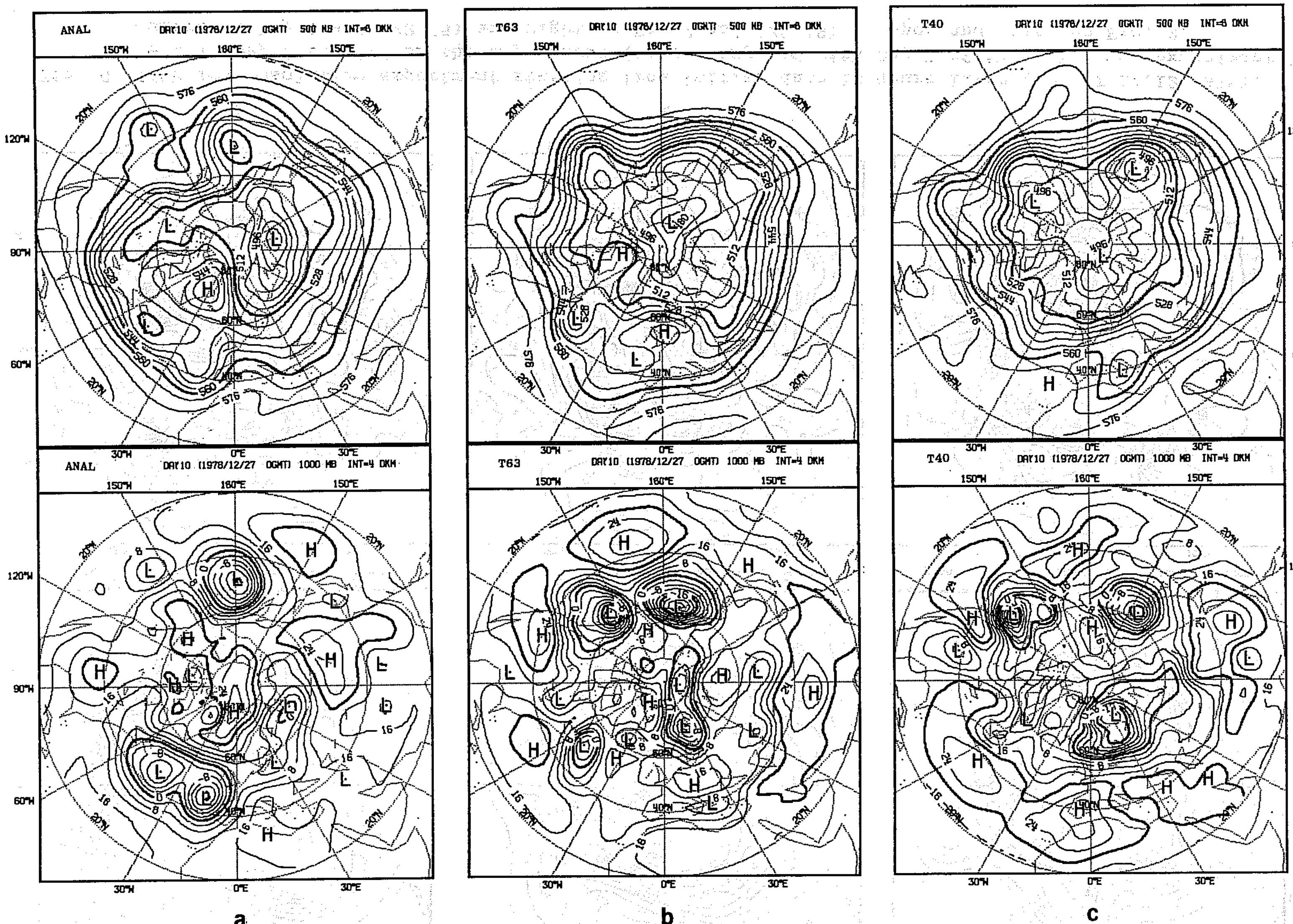


Fig. 5 As Fig. 2 but for Day 10.

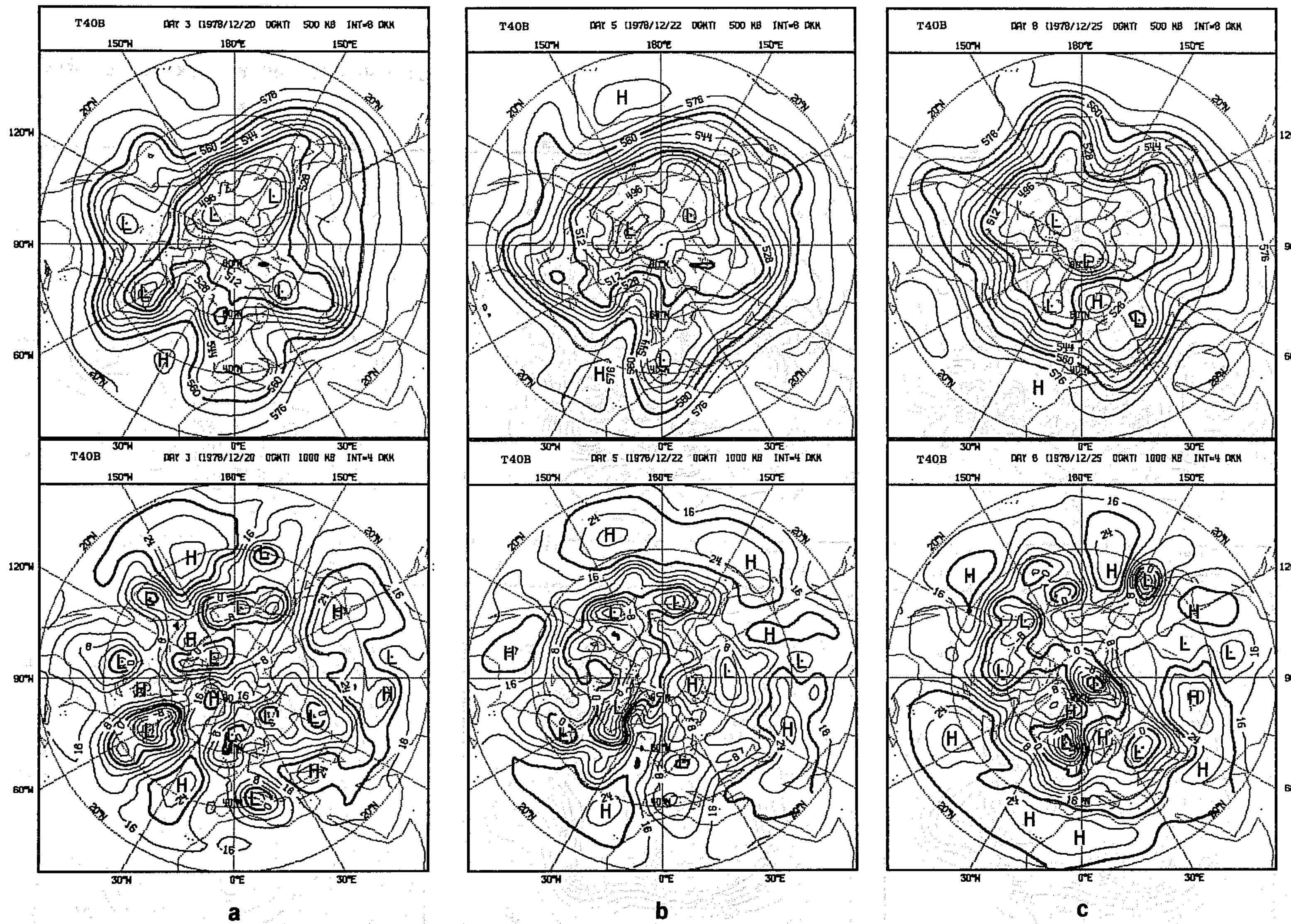


Fig. 6 T40B low resolution experiment starting from initial data 12 hours later (00 GMT 17.12.1978). Left is Day 2 and 1/2 of the experiment (to be compared with Day 3 of the other two experiments); centre is Day 4 and 1/2 (5) and right is Day 7 and 1/2 (8). Upper and lower as Fig. 2.

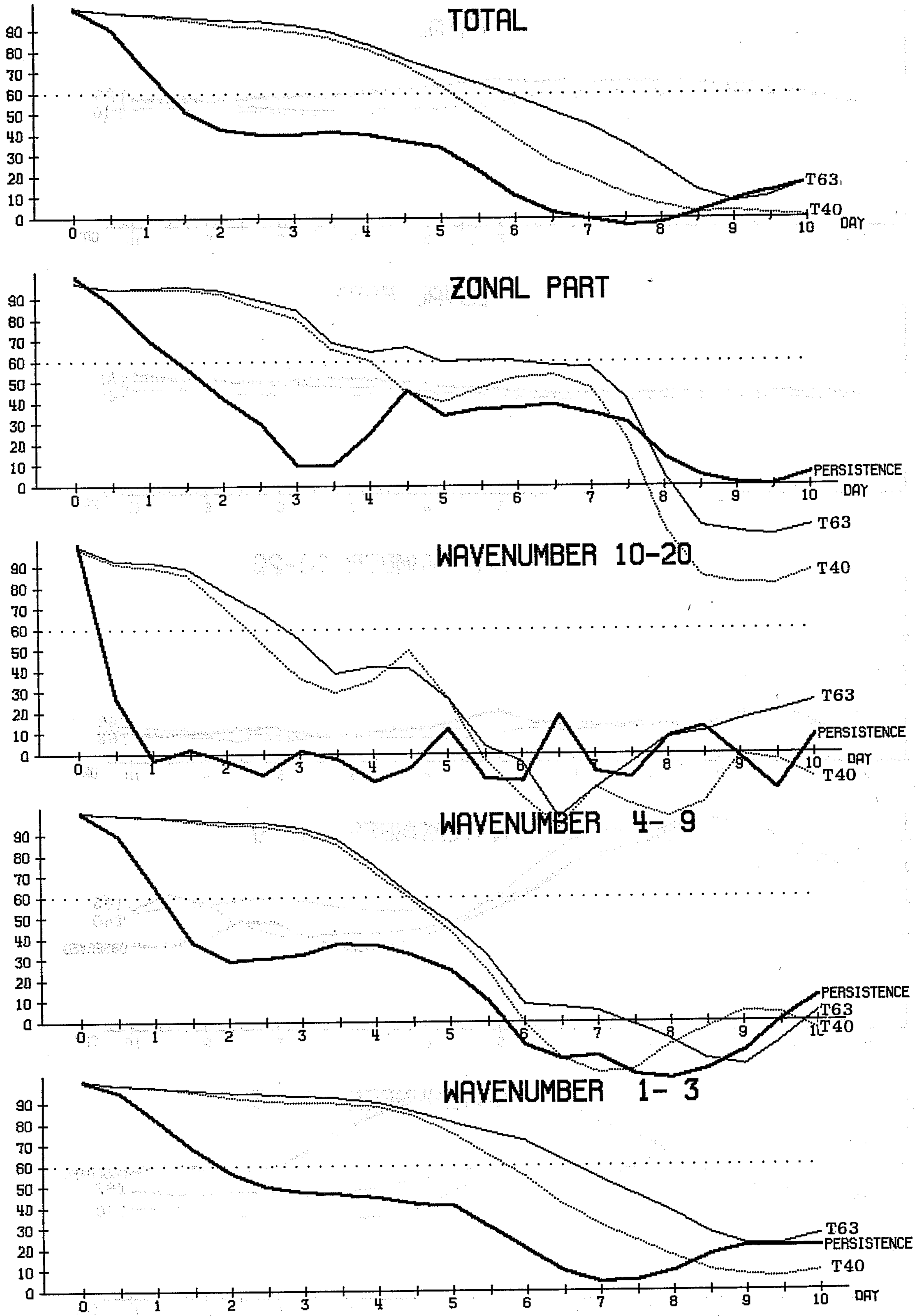


Fig. 7 Anomaly correlation of heights for the T63 and the T40 experiments as a function of time into the forecast. A spectral breakdown is shown, together with the total score. The correlation coefficient is computed between observed anomaly fields (climatology having been subtracted), from 20° to 82.5°N and from 1000 to 200 mb.

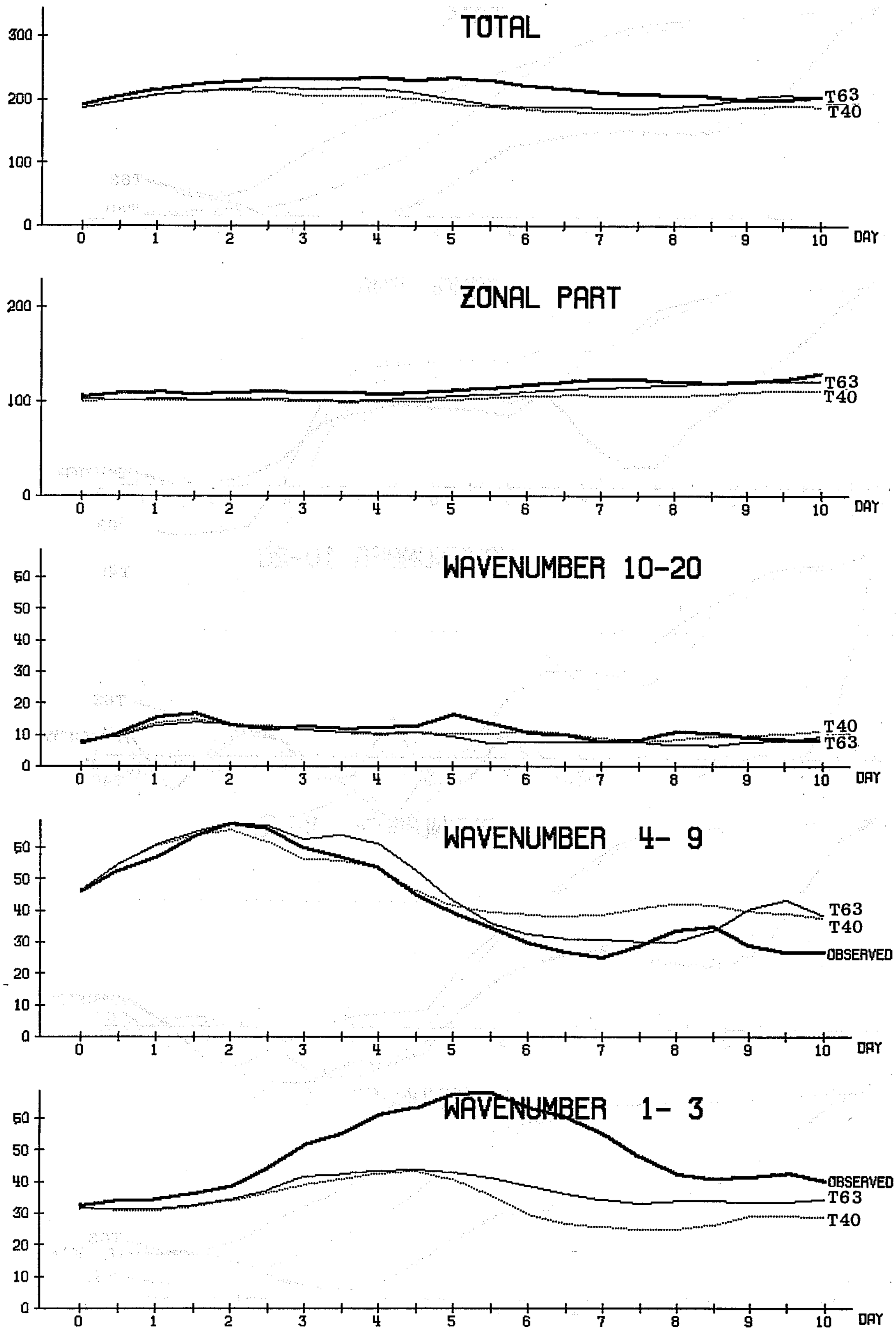


Fig. 8 Total kinetic energy as a function of time. Spectral breakdown and volume mean as for Fig 7.

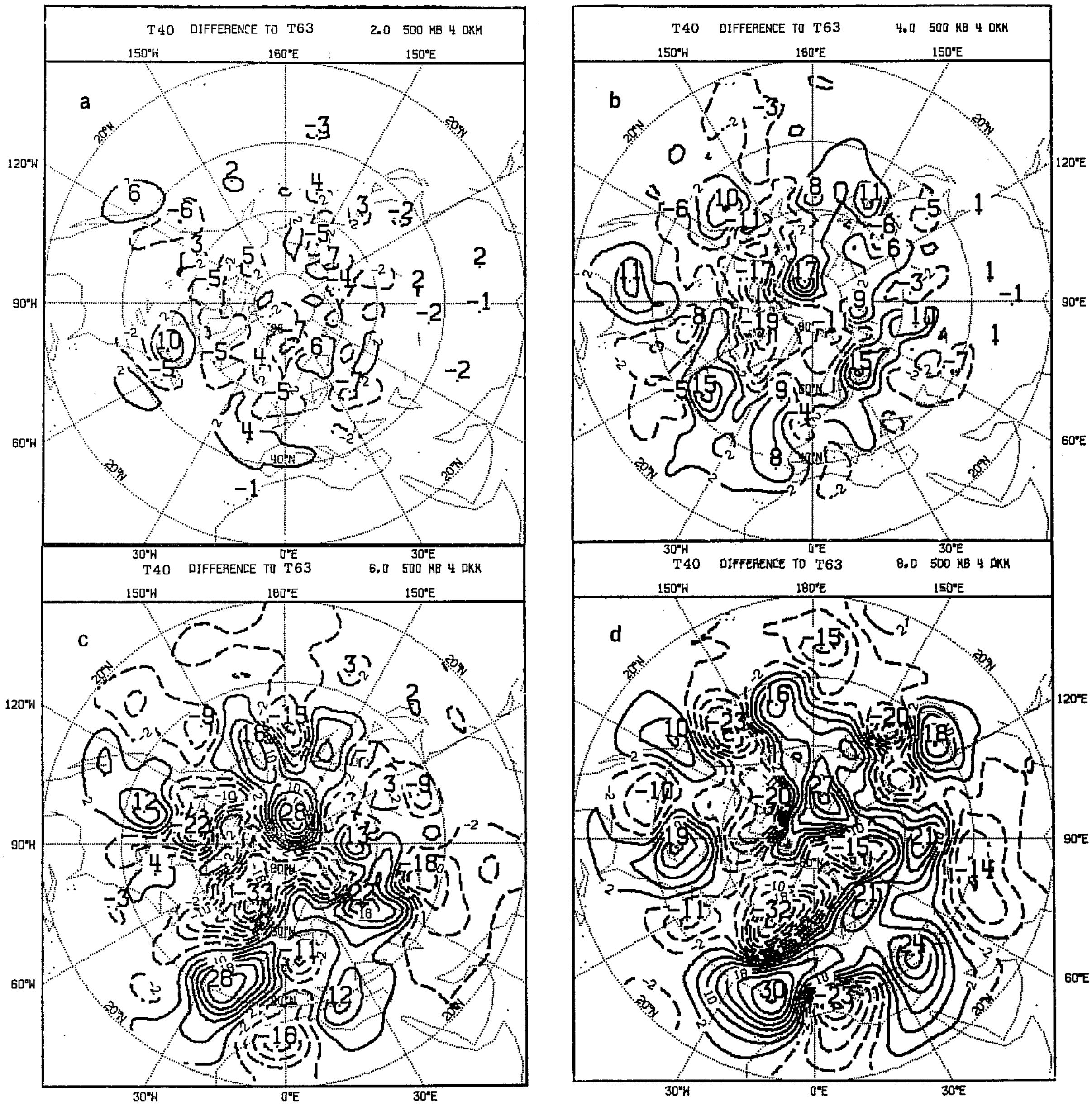


Fig. 9 Difference maps T63 run minus T40 run. a) Day 2, b) Day 4, c) Day 6 and d) Day 8. only 500 mb geopotential height is shown.