

Using Meteorological Analyses for Off-Line Chemical Transport Modelling

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- Description of Off-Line Models
- Example CTM results
- Formulation of a TOMCAT/SLIMCAT models
 - Mass conservation
 - Vertical motion
 - Chemical data assimilation
- Summary

Off-Line Chemical Transport Modelling

Off-line CTM:

- Meteorology (winds, temperature, humidity?) specified from analyses.
- Model integrates chemical species.
- Model constrained to 'real' meteorology - good for comparison with observations.
- Computationally simpler (no internal variability) and cheaper than a coupled chemistry-GCM.

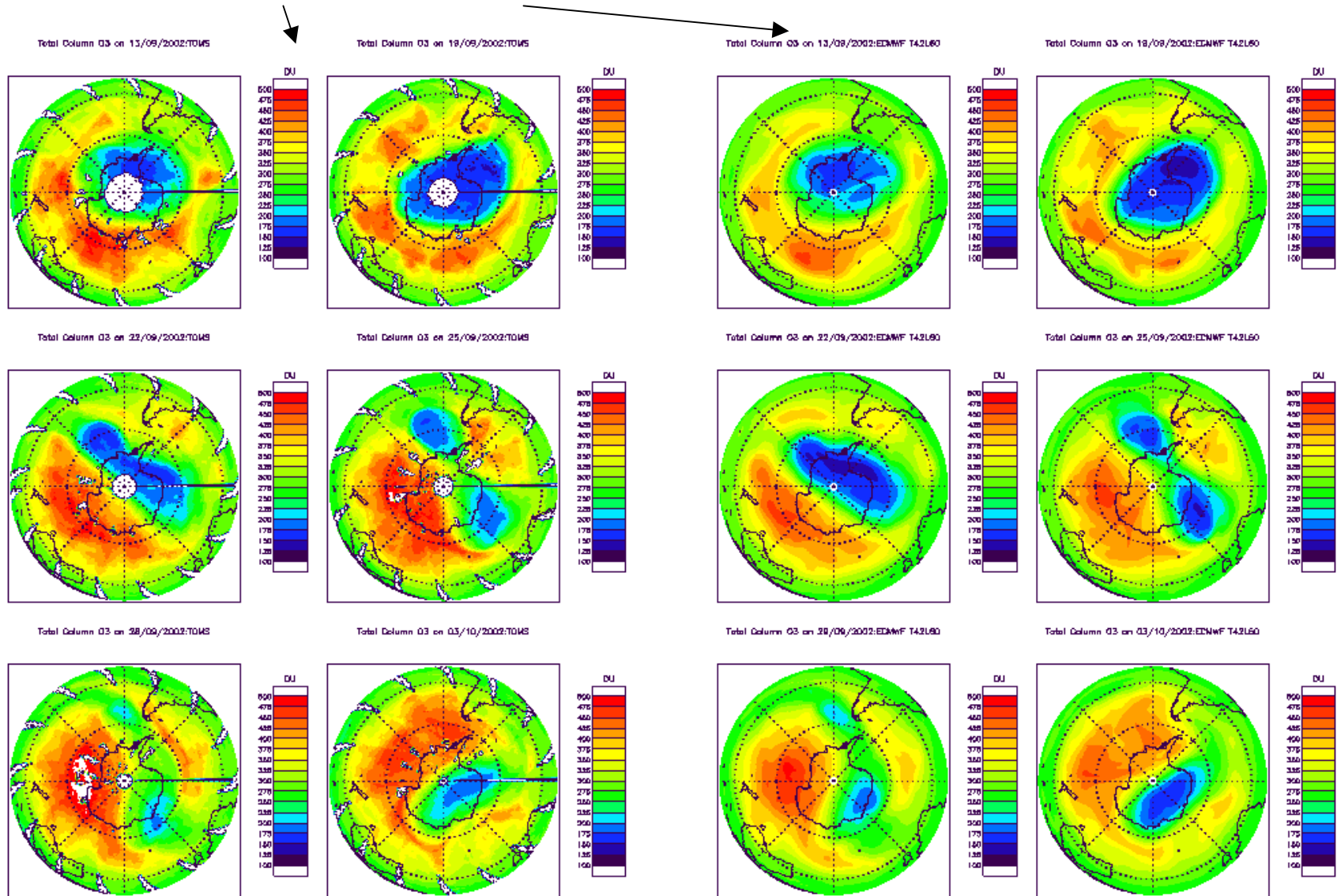
Principle of how a CTM works is simple - but in practice exactly how analyses are used is critical.

Off-line Chemical Modellers (+ trajectory modellers) are a large user group for meteorological analyses.

Experiences of CTM modellers provide useful information on quality of analyses.

Example CTM results 1: 2002 Antarctic O₃ Hole

TOMS v 3D CTM (SLIMCAT, ECMWF analyses)



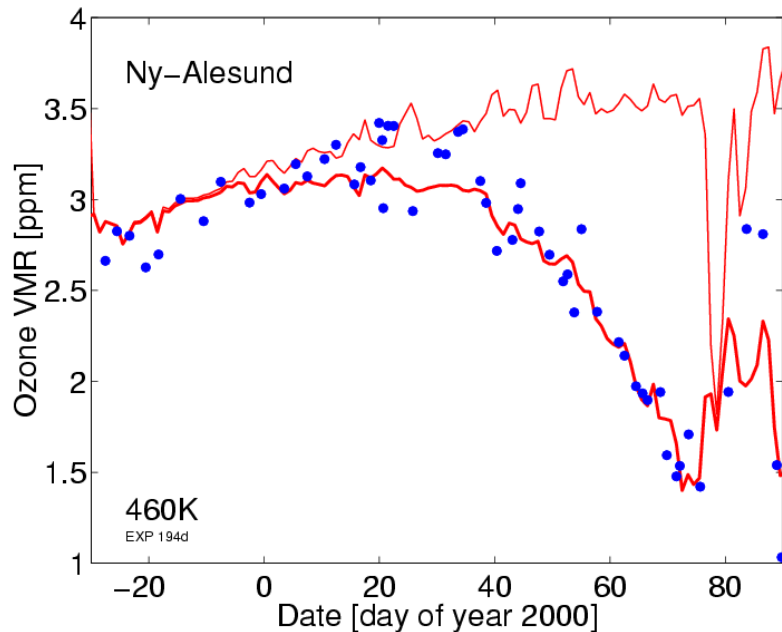
Analyses appear to capture unusual 2002 Antarctic dynamics well

Example CTM Results 2: Arctic O₃ Loss

*SLIMCAT CTM. 5° x 5°.
UKMO winds.*

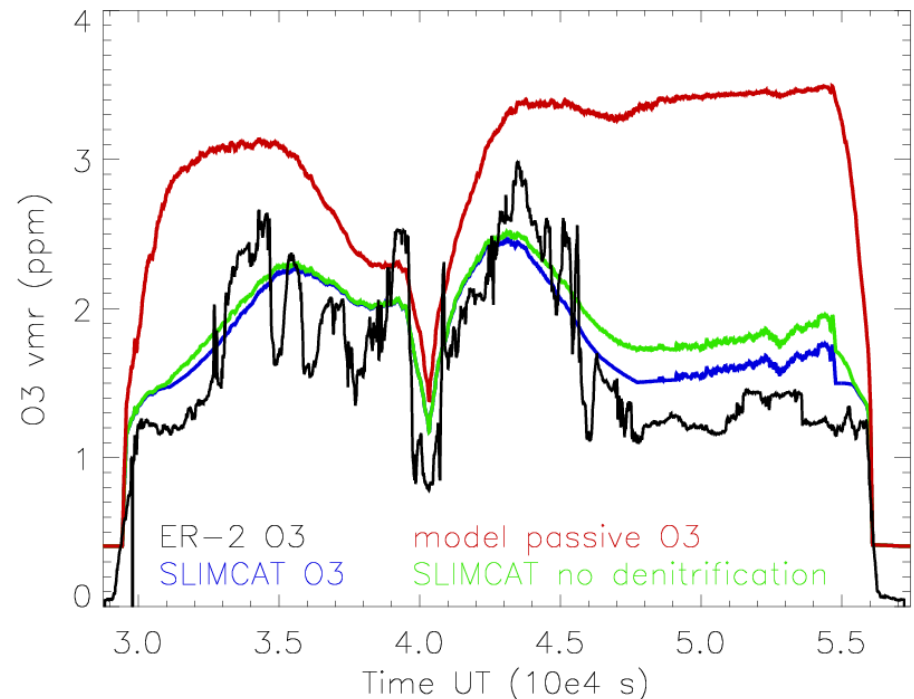
Arctic O₃ loss 1999/2000

3D CTM v O₃ sonde at Ny Alesund



3D CTM v ER-2 in-situ observations

(Arctic vortex flight March 2000 ~18 km)



CTM 3: Long-Term NO₂ variations

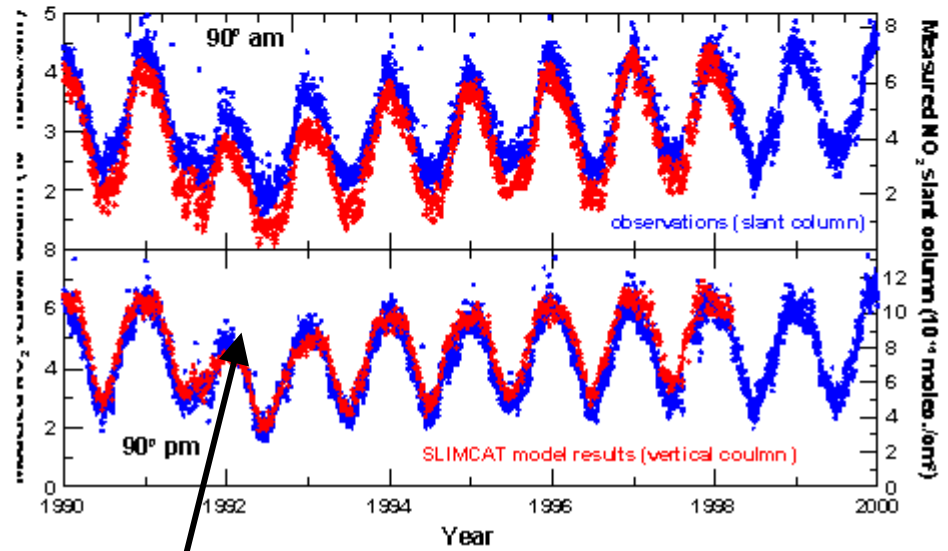
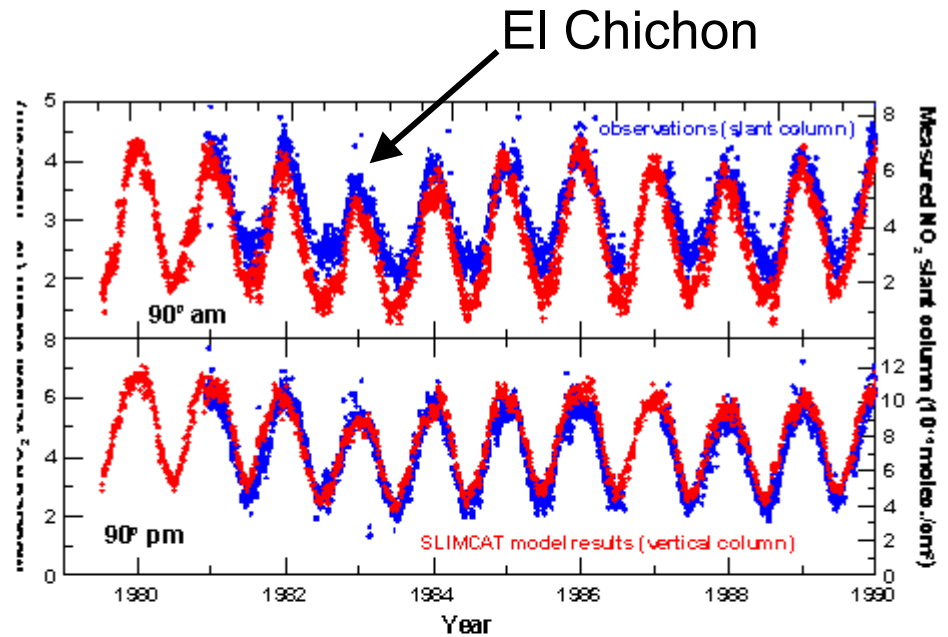
Groundbased column NO₂
at Lauder (45°S)

(P. Johnston + K. Kreher)

v SLIMCAT 3D CTM
(ECMWF ERA15 run to
exploit long Lauder time
series) (assuming AMF=17)

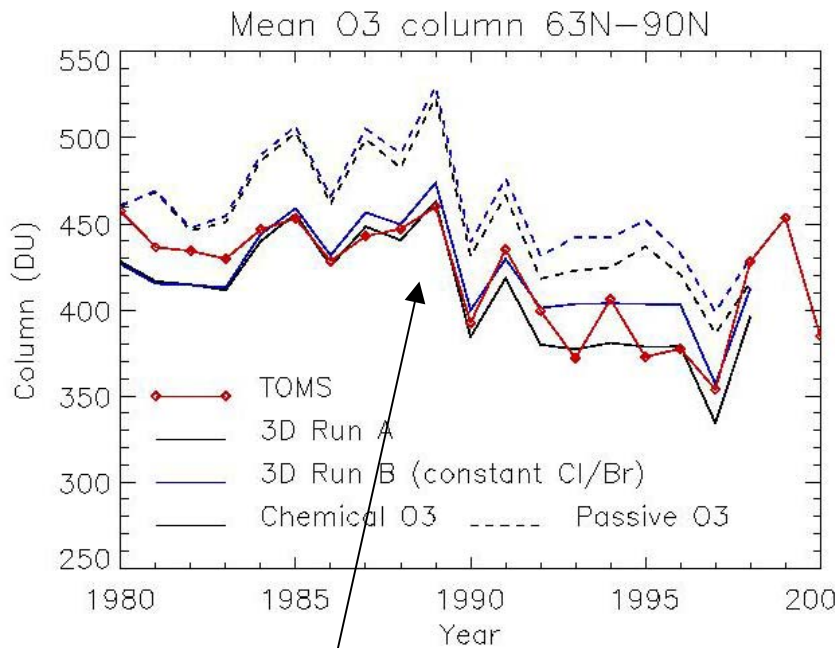
3D model captures: (i)
seasonal cycle and (ii)
aerosol-induced variability.

Also seems to capture a lot
+6%/year trend (study on-
going - see Kreher et al.
AGU 2002).

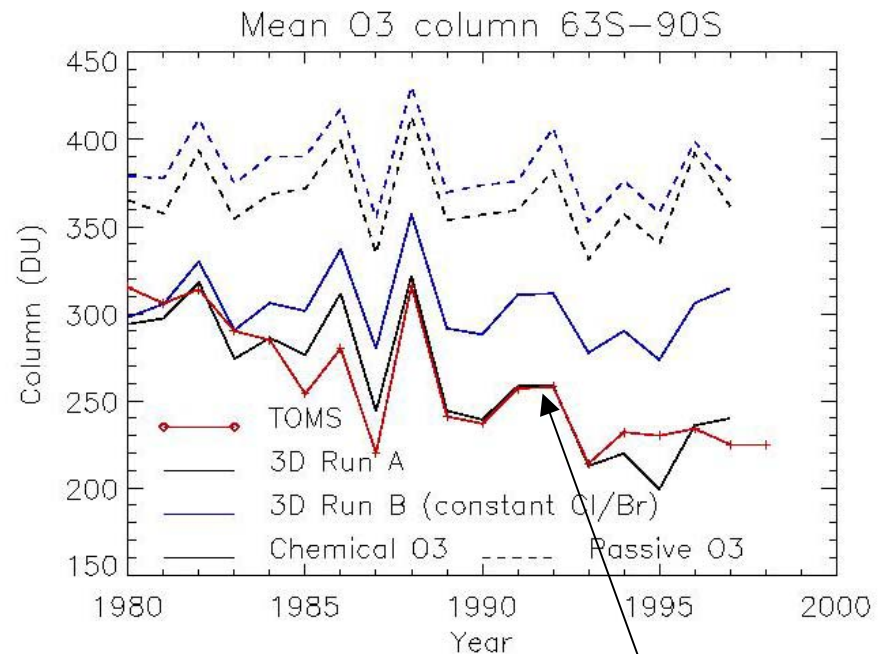


CTM Example 4: Polar O₃ Variability and Trend

Mean NH March/ SH October Column O₃ 63°-90° v TOMS data (P. Newman)
SLIMCAT CTM forced by ERA15 (+ operational) winds.



Model captures interannual variability (esp. pre-1994 ERA period).

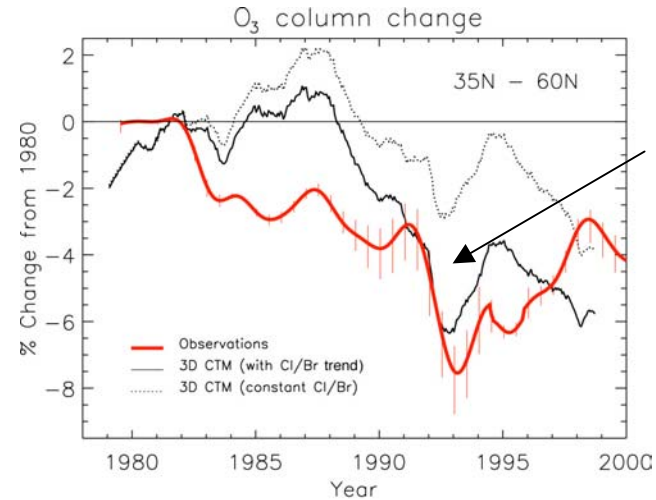
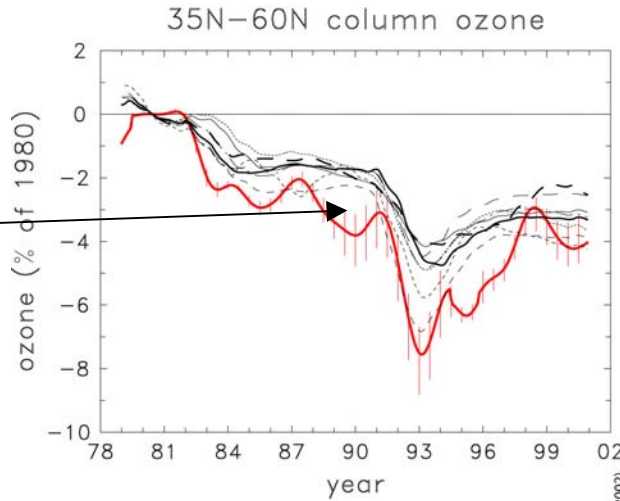


Chemical loss (trend) in SH ok. Models known to underestimate Arctic (polar) loss (see posters/Wednesday)

CTM Examples 5: Mid-Latitude O₃ Trends

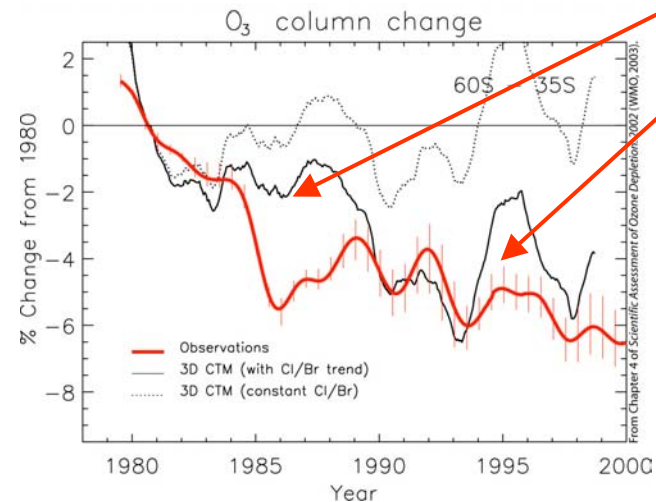
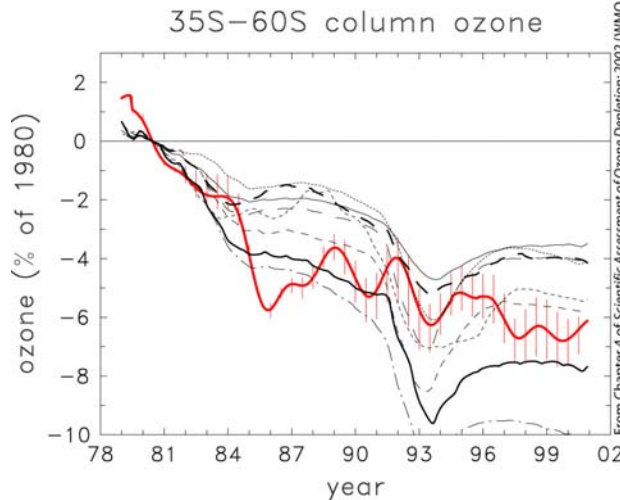
2D Models - Cly/Br trends;
climatological winds; poor
polar treatment.

SLIMCAT 3D CTM - analysed
winds, better polar treatment.
(ERA15 + op winds)



Chemistry/
dynamics

Aerosol
chemistry



Solar
cycle???

SH - 2D
poor

From Chapter 4 of Scientific Assessment of Ozone Depletion: 2002 (WMO, 2002).

From Chapter 4 of Scientific Assessment of Ozone Depletion: 2002 (WMO, 2002).

TOMCAT/SLIMCAT 3D CTMs

*Chipperfield and Simon,
CNRM Toulouse, 1991/92.*

Two related off-line CTMs (in fact same library):

- Horizontal/vertical resolution variable.
- Horizontal winds/temperatures from analyses (ECMWF, UKMO).
- Vertical coordinate: σ -p (TOMCAT), σ - θ (SLIMCAT).
- Vertical motion from either:
 - Divergence (of mass flux).
 - Diagnosed heating rates (SLIMCAT in θ coordinates).
- Default **advection** scheme: Prather [1986] - finite volume.
- **Convection**: Tiedtke [1989] - (only requires T, u, v, q).
- 'Full' **chemistry** schemes stratosphere/troposphere (TOMCAT Cambridge).
- Chemical **data assimilation** scheme: sub-optimal Kalman filter [Khattatov]

Why SLIMCAT?

- θ coordinates better in stratosphere (correct separation of horizontal/vertical motion; reduces numerical diffusion).
- Analysed vertical winds can be 'noisy'.

SLIMCAT/TOMCAT Stratospheric Chemical Scheme

Integrated shorter lived species O_x ($O_3 + O(^3P) + O(^1D)$)
 NO_x ($N + NO + NO_2$), NO_3 , N_2O_5 , HNO_3 , HO_2NO_2
 ClO_x ($Cl + ClO + Cl_2O_2$), $OCIO$, HCl , $ClONO_2$, $HOCl$
 BrO_x ($Br + BrO$), HBr , $BrONO_2$, $HOBr$, $BrCl$
 CH_3OOH , CH_2O

Steady-state H , OH , HO_2
 CH_3O , CHO , CH_3O_2

Long-lived species CH_4 , N_2O , H_2O , CO
 $CFCl_3$, CF_2Cl_2 , CH_3Br

100 gas-phase reactions

20 photolysis reactions

~9 heterogeneous reactions on liquid sulphate aerosols and solid polar stratospheric clouds (PSCs).

CTM Vertical Coordinate

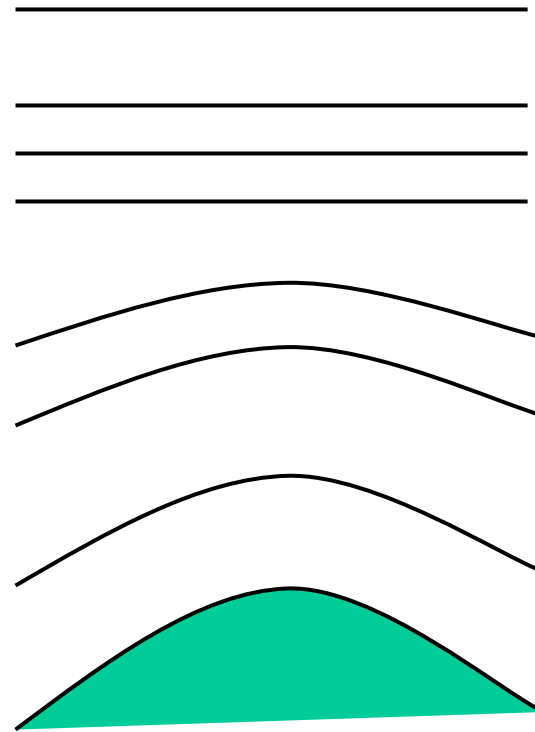
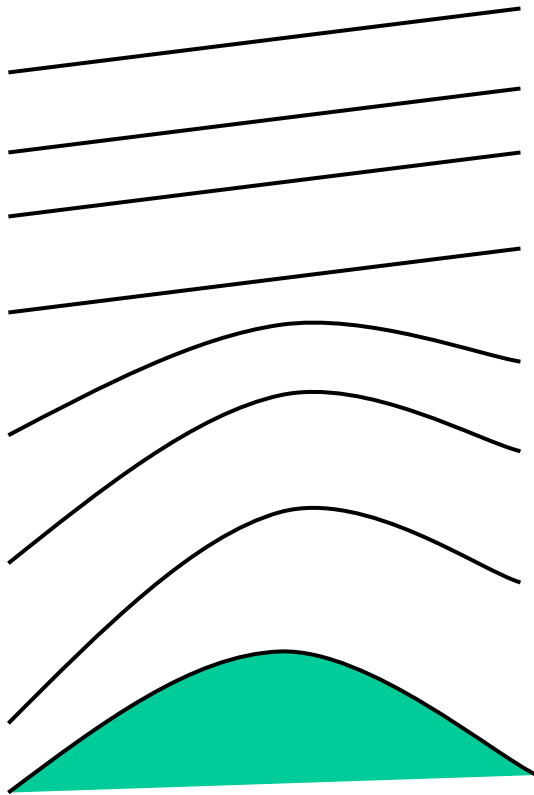
'SLIMCAT' mode

hybrid σ - θ

'TOMCAT' mode

hybrid σ -p

Lowest
pure θ
level = 350
K typically



Surface

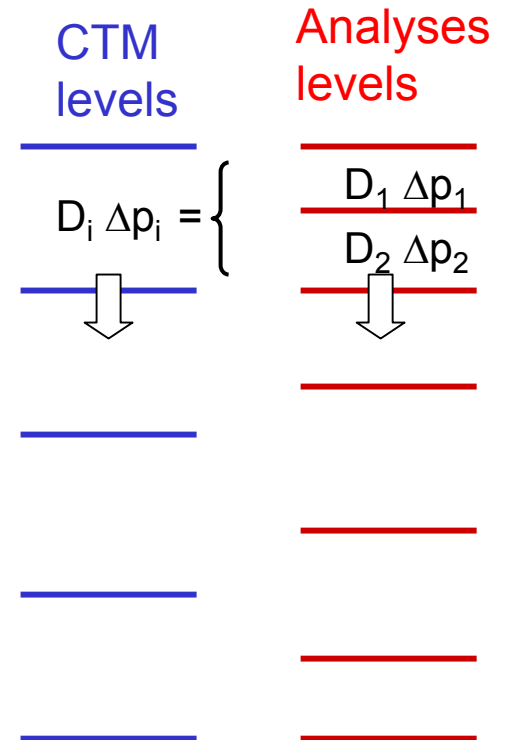
'Philosophy' of Off-Line Approach

Off-line model should replicate meteorological model (advection, convection)

Minimise manipulations (interpolation etc) of analyses, but CTM will likely want different grid (θ levels, irregular latitudes etc).

In SLIMCAT/TOMCAT when forced by ECMWF analyses:

- Model reads spectral coefficients of V , D , T , q , p_s (at any resolution - ideally on original analysis levels).
- V and D of winds converted to V and D of mass fluxes.
- Horizontally: Spectral transform of mass fluxes averaged over model grid cell edges.
- Vertically: Horizontal mass fluxes distributed over CTM model levels. ($\int D dp$ preserved).



Mass Conservation

A problem for off-line models. A few different aspects to this problem:

Necessary pre-requisites:

- Mass-conserving advection scheme.
- Conserving chemistry scheme.

Then, in addition, need good formulation of model:

- Transport defined in terms of mass fluxes.
- Vertical mass fluxes from divergence of horizontal mass fluxes.
- Conversion of analyses from analysis grid to model grid done in conserving way.

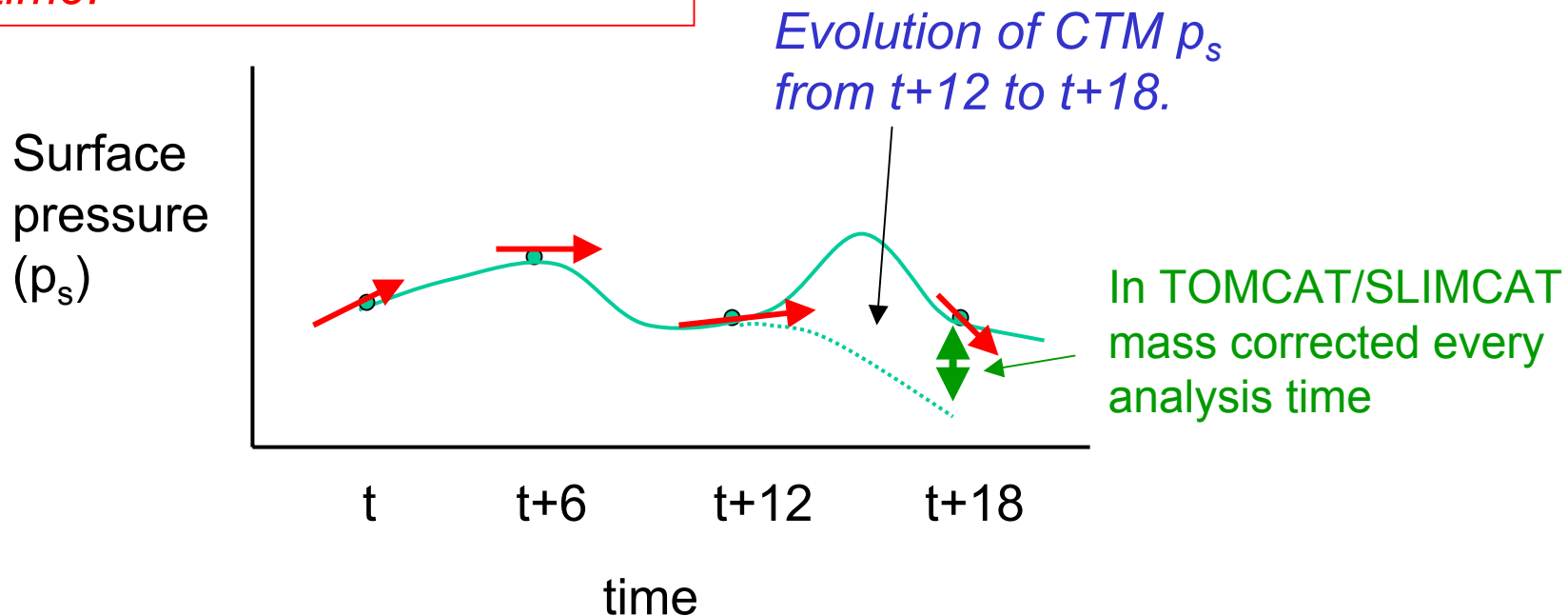
Mass Conservation (cont)

Even though mass fluxes may balance instantaneously, time discretisation will prevent complete mass conservation:

CTM only has 6-hourly fields (p_s , dp_s/dt from divergence).

CTM interpolates fields linearly in time.

This local mass correction possible so long as analyses conserve mass (which they don't !)



Three Laws of Off-Line Modelling

1. All good CTMs should conserve mass exactly.
2. A CTM can only conserve mass properly if the analysis (meteorological) model conserves mass.
3. Analysis (meteorological) models don't conserve mass.

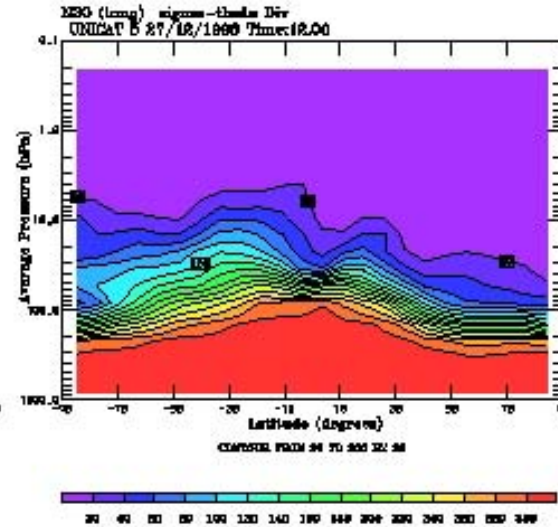
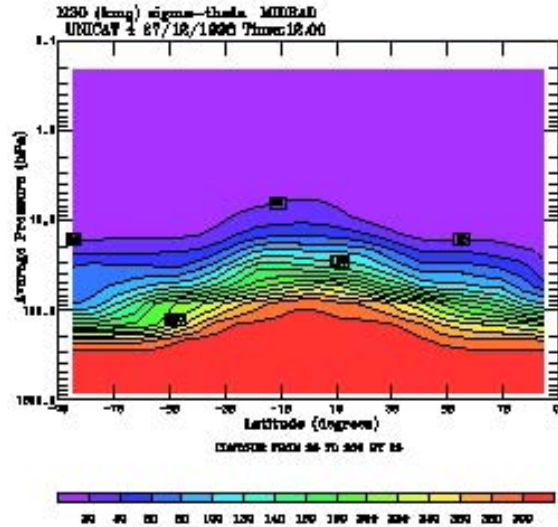
SLIMCAT/TOMCAT 3D CTM forced by ERA40 Analyses

Different vertical coordinates and vertical motion

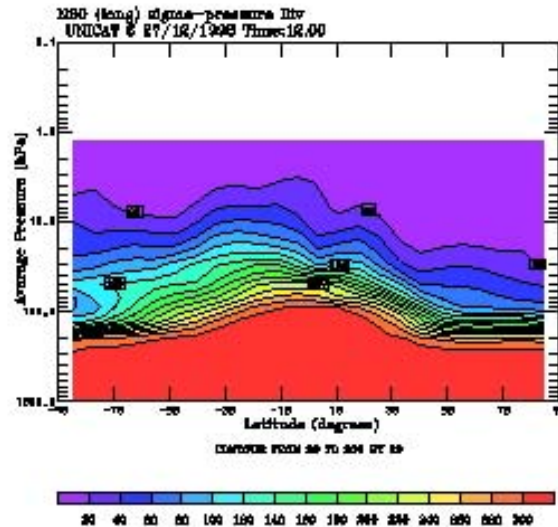
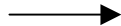
N_2O -like tracer.

27/12/1998.

Sigma-theta ($\sigma-\theta$)
Heating rates (MIDRAD)



Sigma-pressure ($\sigma-p$)
Divergence



Sigma-theta ($\sigma-\theta$)
Divergence



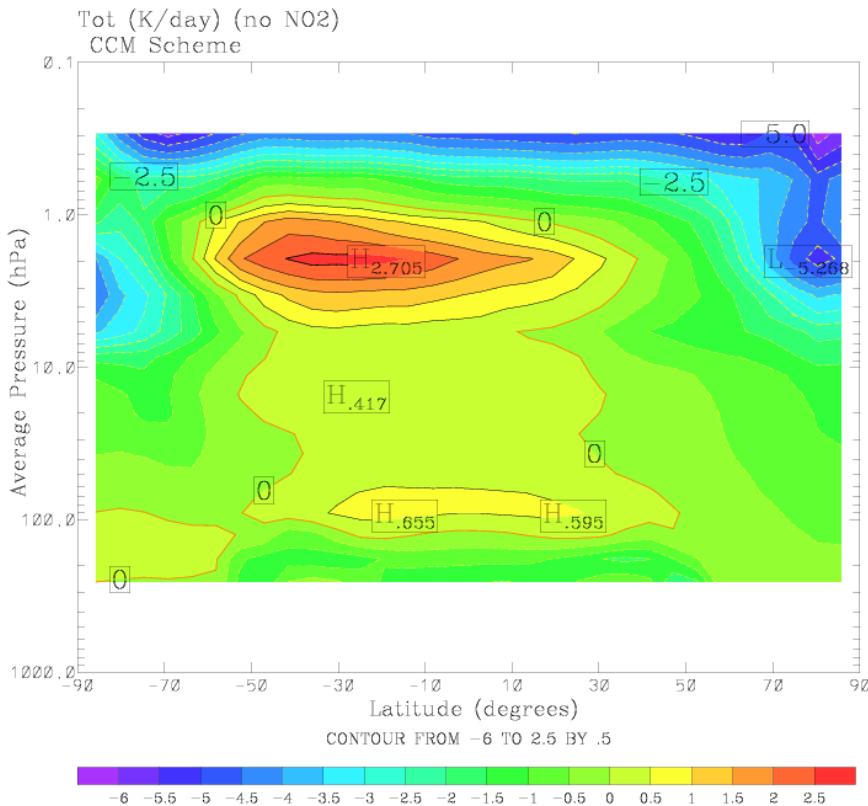
Diagnosing Vertical (Diabatic) Transport (Heating Rates)

SLIMCAT CTM.

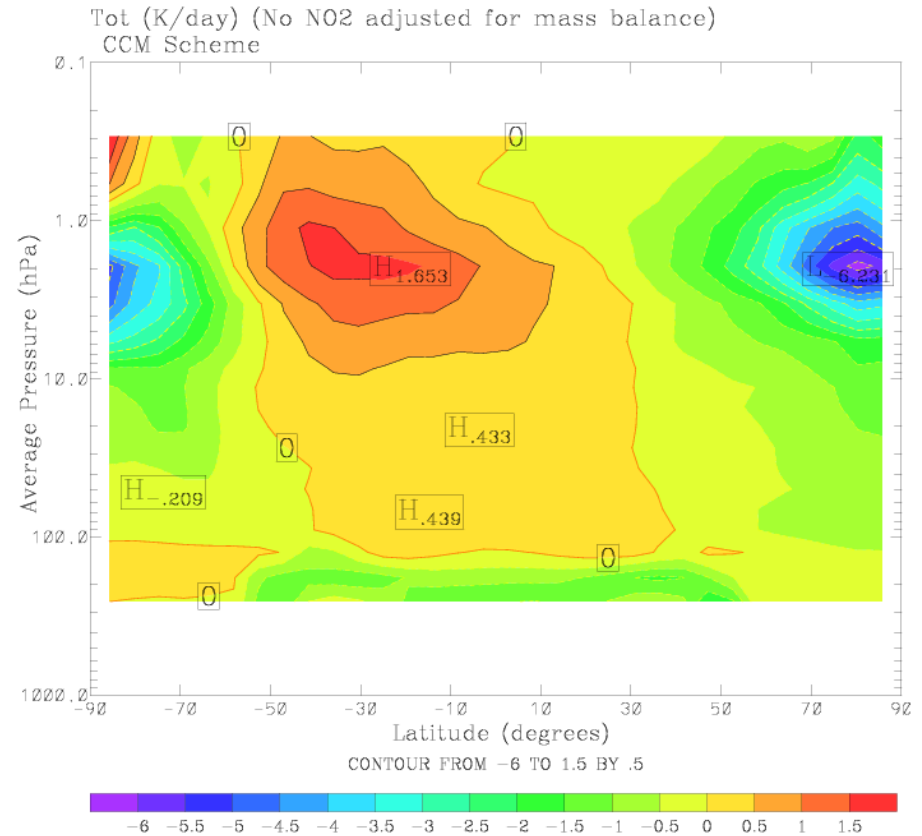
CCM radiation scheme.

ERA-40 temperatures.

Net heating rate



Net heating rate - with imposed mass balance



Sequential Chemical Data Assimilation in a 3D CTM

Included sub-optimal Kalman filter chemical assimilation scheme (Khattatov, NCAR) into a full chemistry 3D atmospheric model (SLIMCAT).

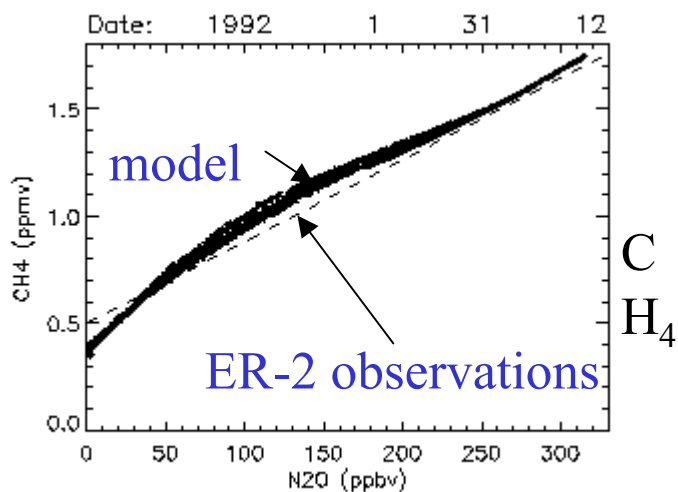
A unique improvement of our scheme over previous uses of this scheme is the preservation of compact correlations during assimilation process.

Scheme is useful for assimilation long-lived tracers (not short-lived species)

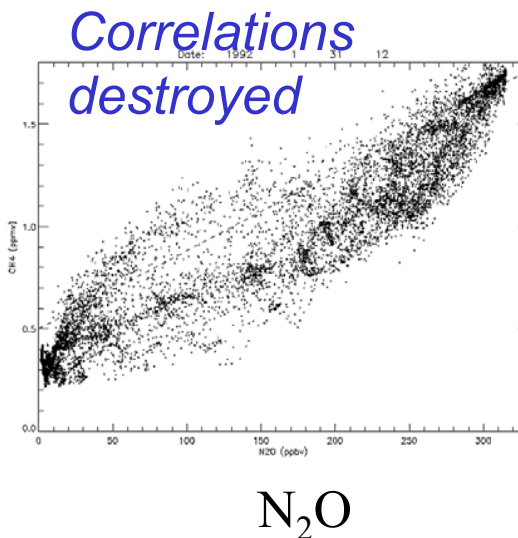
See Chipperfield et al., JGR, [2003].

Tracer-tracer correlations: CH₄ v N₂O from SLIMCAT 3D CTM

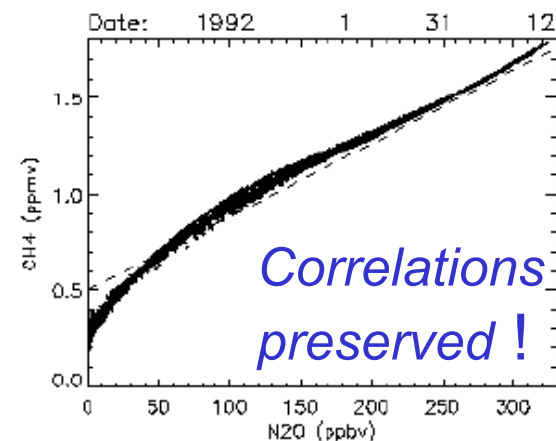
- ⇒ Long-lived tracers display compact correlations (e.g. Plumb + Ko, 1992)
- ⇒ Total abundance of chemical species limited (e.g. ΣCl)
- ⇒ **Chemical assimilation should make use of these constraints!**



Basic (non -
assimilation model)



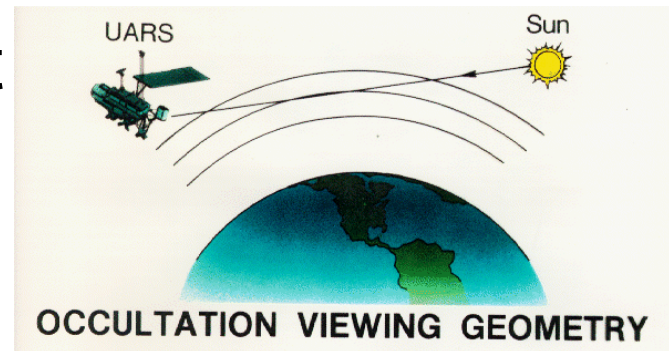
Naïve assimilation
of HALOE CH₄ only.



Assimilation of
HALOE CH₄, and
adjustment of N₂O to
preserve correlation.

HALogen Occultation Experiment

(Russell et al.)



- Launched on Upper Atmosphere Research Satellite (UARS) - Sept 1991

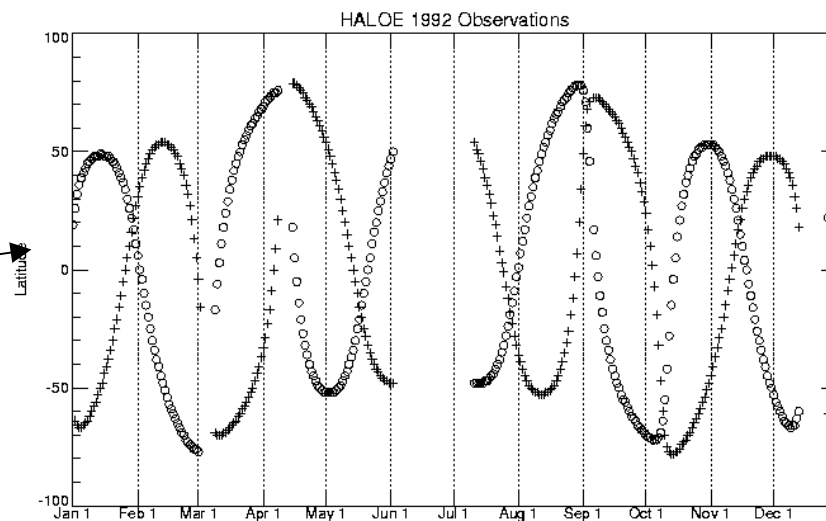
- Solar occultation observations

- Latitude coverage over ~1 month

- Here we use profiles of O_3 , CH_4 , H_2O , and HCl

Start Date: Observed
TPA: 30. km

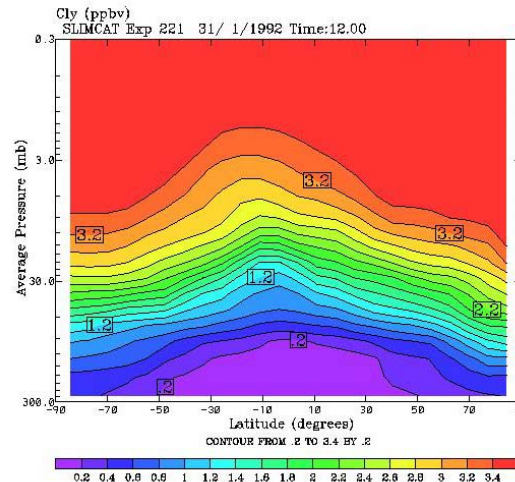
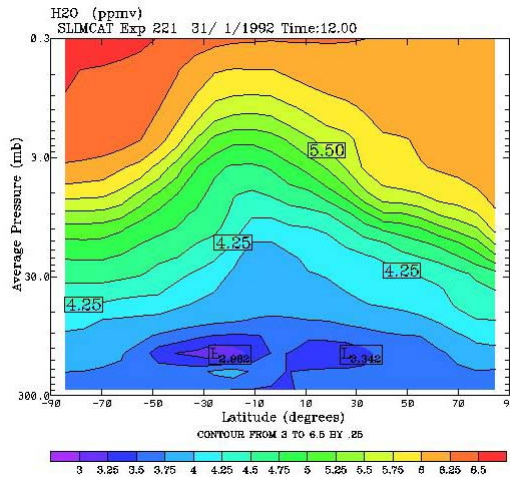
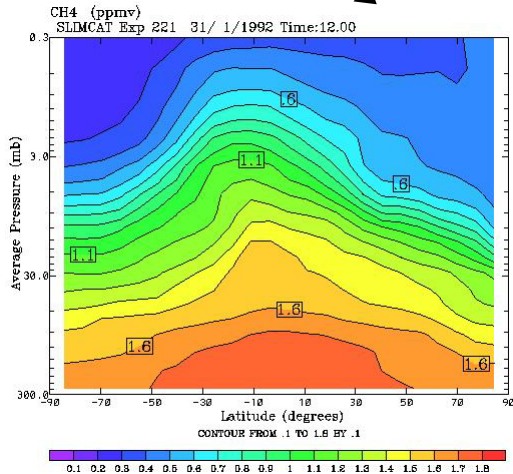
+ SUNSET
o SUNRISE



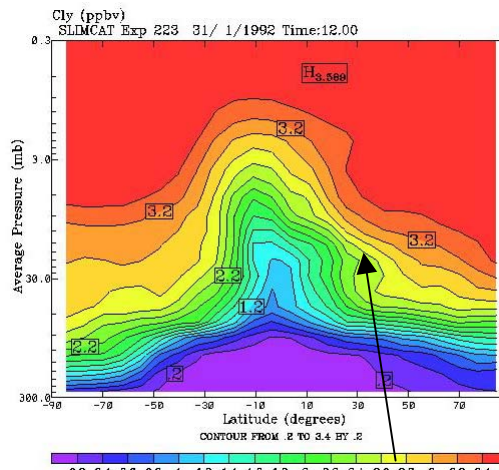
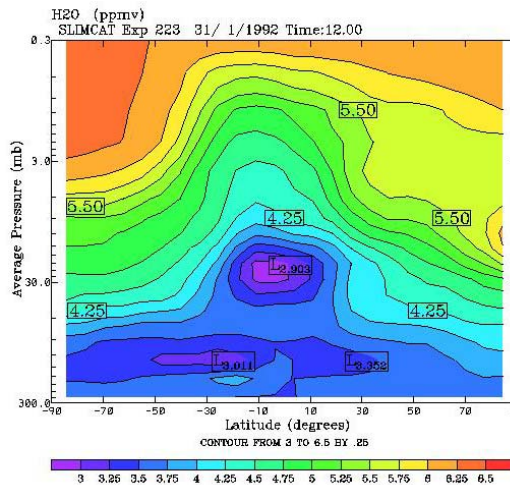
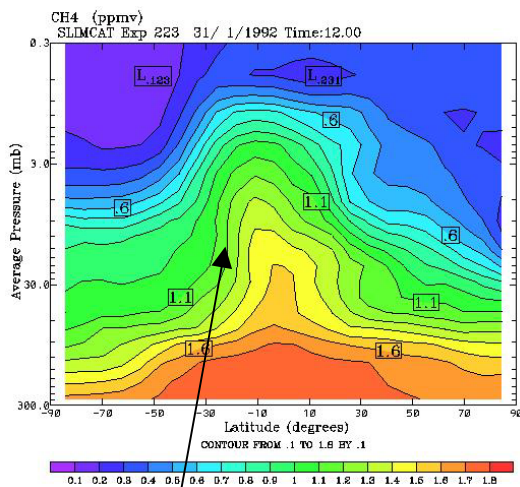
Latitude Progression

Effect of Assimilation: CH₄ , H₂O, Cly (not observed)

UKMO
winds



No
Assim.



Assim.

Large change in
subtropics

Zonal mean fields 31/1/1992

Change to Cly occurs
through tracer-tracer
correlations with CH₄

Suggestions for Analysis Products

Quality

Temperatures in LS: < 1K error?

Horizontal winds in tropics

Vertical winds (BD circulation)?

Products

Diabatic heating rates (LW + SW components, contributions from latent heat release in troposphere)

Convective mass fluxes.

Higher top boundary (transport from mesosphere?)