Parameterization of Boundary Layer Clouds:

A GCSS perspective

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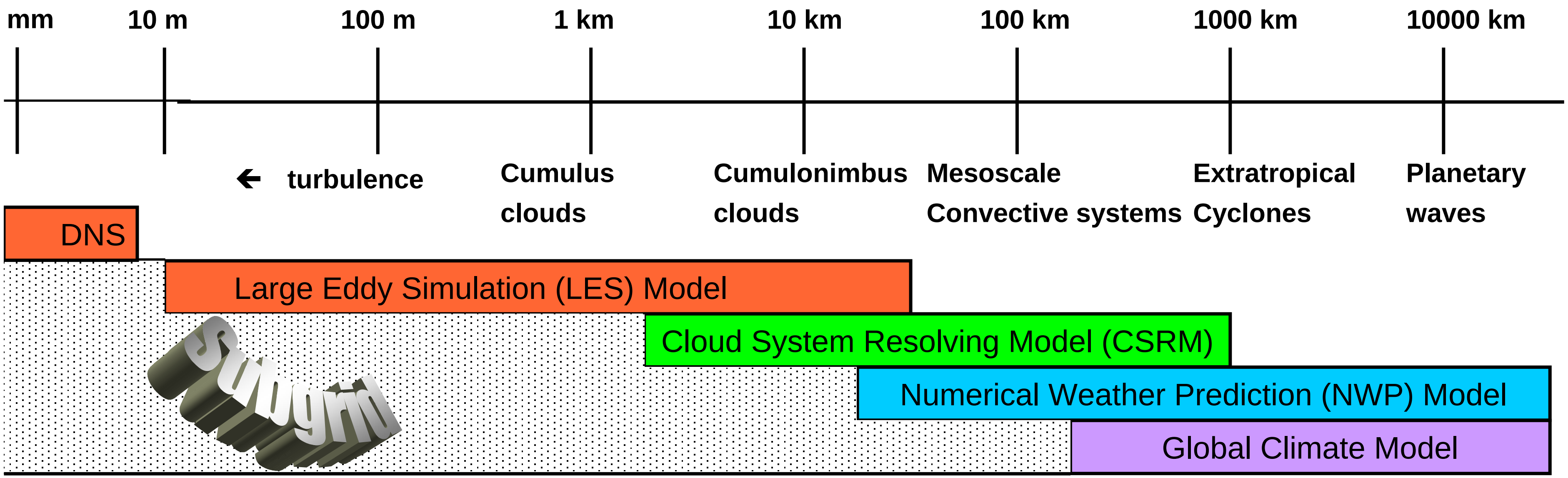
The Zoo of Atmospheric Models



mm | 10 m | 100 m | 1 km | 10 km | 100 km | 1000 km | 10000 km

← turbulence

Cumulus clouds

Cumulonimbus clouds

Mesoscale Convective systems

Extratropical Cyclones

Planetary waves

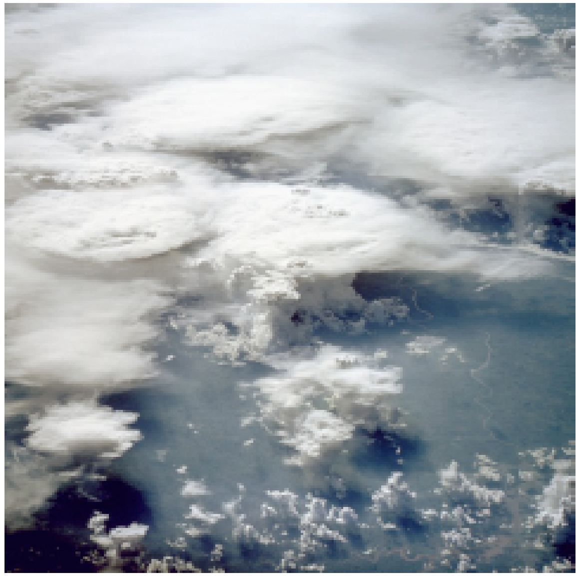
DNS

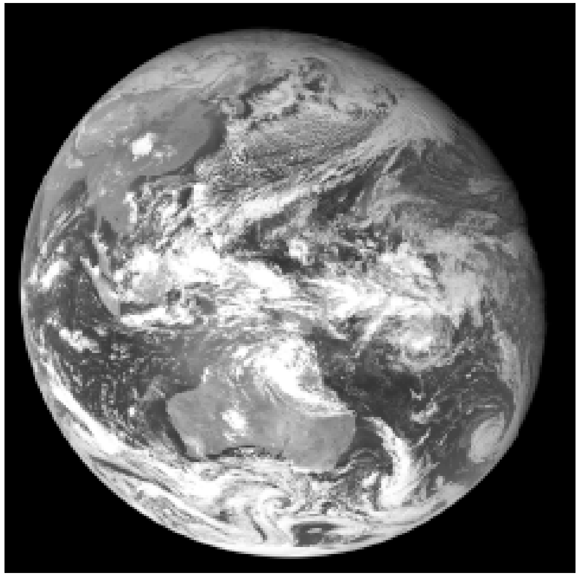
Large Eddy Simulation (LES) Model

Cloud System Resolving Model (CSRM)

Numerical Weather Prediction (NWP) Model

Global Climate Model

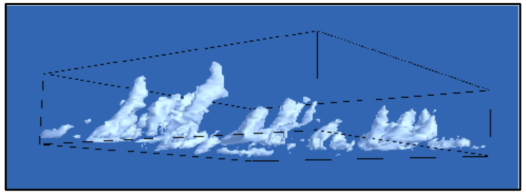


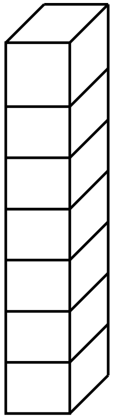


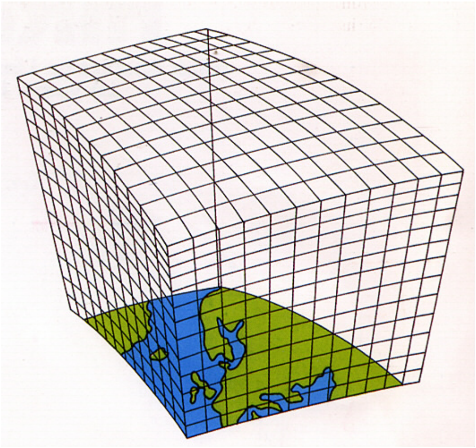
GEWEX Cloud Systems Studies (GCSS)

(Simplified) Working Strategy

See http://www.gewex.org/gcss.html







Large Eddy Simulation (LES) Models

Cloud Resolving Models (CRM)

Single Column Model

Versions of Climate Models

3d-Climate Models

NWP's

Observations from
Field Campaigns

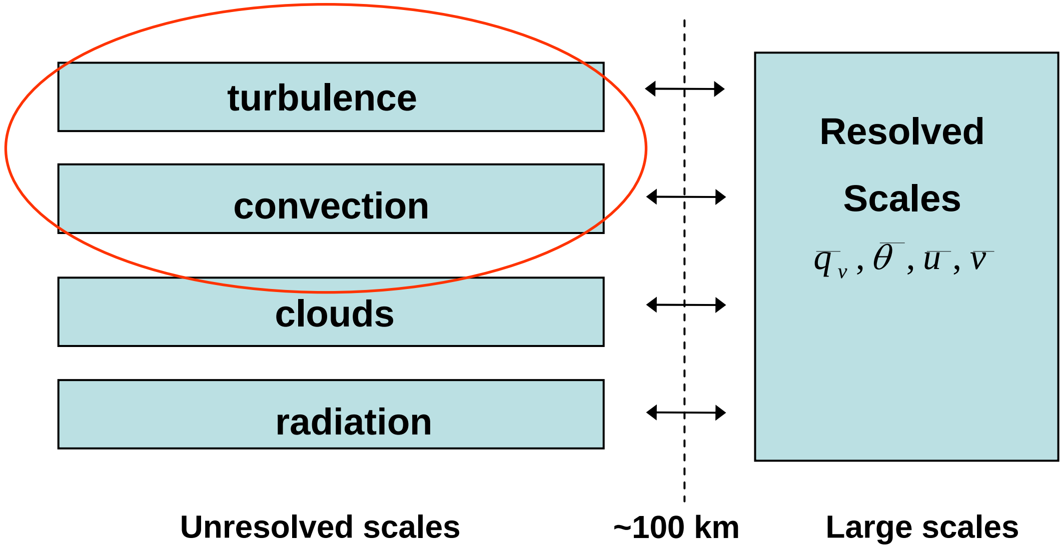
Global observational
Data sets

Development

Testing

Evaluation

History and Progress in Convential Parameterizations for the Cloudy PBL



turbulence

convection

clouds

radiation

Resolved

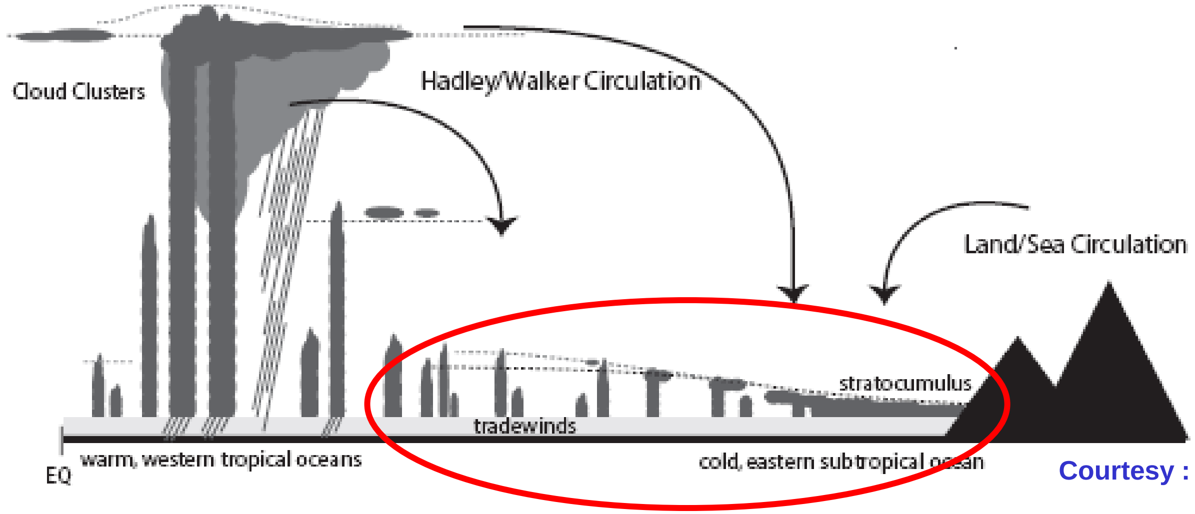
Scales

$\overline{q_v}, \overline{\theta}, \overline{u}, \overline{v}$

Unresolved scales

~100 km

Large scales



Courtesy : Bjorn Stevens

Grid Averaged Equations of thermodynamic variables

$$\frac{\partial\overline{\theta}}{\partial t} = -\mathbf{v}\cdot\nabla\overline{\theta} - w\frac{\partial\overline{\theta}}{\partial z} - \frac{\partial}{\partial x_i}\overline{u_i'\theta'} + \frac{L}{\pi c_p}(c-e) + Q_{rad}$$

$$\frac{\partial\overline{q}_v}{\partial t} = -\mathbf{v}\cdot\nabla\overline{q}_v - w\frac{\partial\overline{q}_v}{\partial z} - \frac{\partial}{\partial x_i}\overline{u_i'q_v'} - (c-e)$$

$$\frac{\partial\overline{q}_l}{\partial t} = -\mathbf{v}\cdot\nabla\overline{q}_l - w\frac{\partial\overline{q}_l}{\partial z} - \frac{\partial}{\partial x_i}\overline{u_i'q_l'} + (c-e) - P_r$$

Large scale advection

Large scale subsidence

turbulent transport

Net Condensation Rate

Introduce moist conserved variables!

$$\theta_l \approx \theta - \frac{L}{c_p \pi} q_l$$

- Liquid water potential Temperature

$$q_t \equiv q_v + q_l$$

- Total water specific humidity

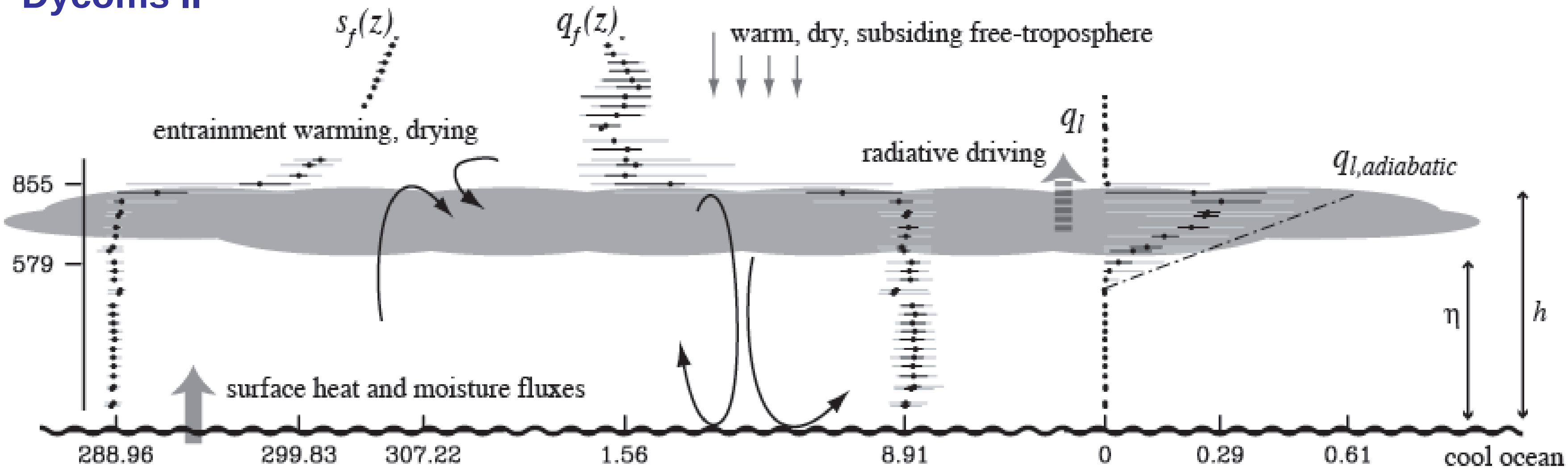
$$\boxed{\begin{aligned}
\frac{\partial \overline{\theta}_l}{\partial t} &= -\overline{\mathbf{v}} \cdot \nabla \overline{\theta}_l - \overline{w} \frac{\partial \overline{\theta}_l}{\partial z} - \frac{\partial}{\partial z} \overline{w' \theta_l'} + Q_{rad} \\
\frac{\partial \overline{q}_t}{\partial t} &= -\overline{\mathbf{v}} \cdot \nabla \overline{q}_t - \overline{w} \frac{\partial \overline{q}_t}{\partial z} - \frac{\partial}{\partial z} \overline{w' q_t'} - P_r
\end{aligned}}$$

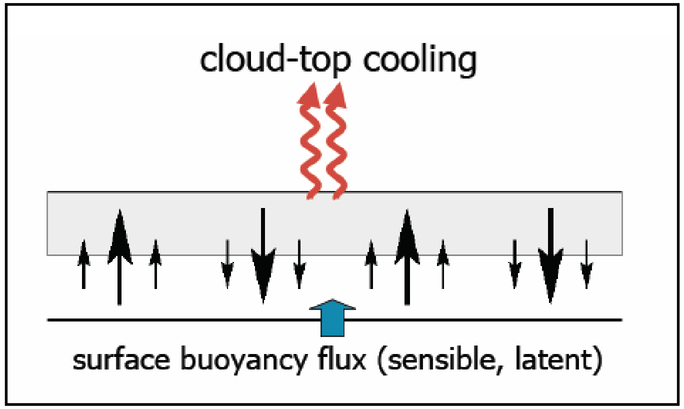
Parameterization issue reduced to finding the subgrid fluxes

Stratocumulus : characteristics and used variables

Courtesy : Bjorn Stevens

Dycoms II





Convection is driven from the top and the bottom

Stratocumulus (2)

A long history in GCCS.

Experiment	Case	year
FIRE	Nocturnal Scu	1994
Idealized Smoke case		1995
ASTEX	Langrangian case	1995
ASTEX	Nocturnal	1996
FIRE	Diurnal cycle	2002
DYCOMSII	Nocturnal Scu	2003
DYCOMSII	Nocturnal Scu Precipitating	2005

Why?

Stratocumulus (3)

Experiment	Case	year
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ASTEX	Langrangian case	1995
ASTEX	Nocturnal	1996
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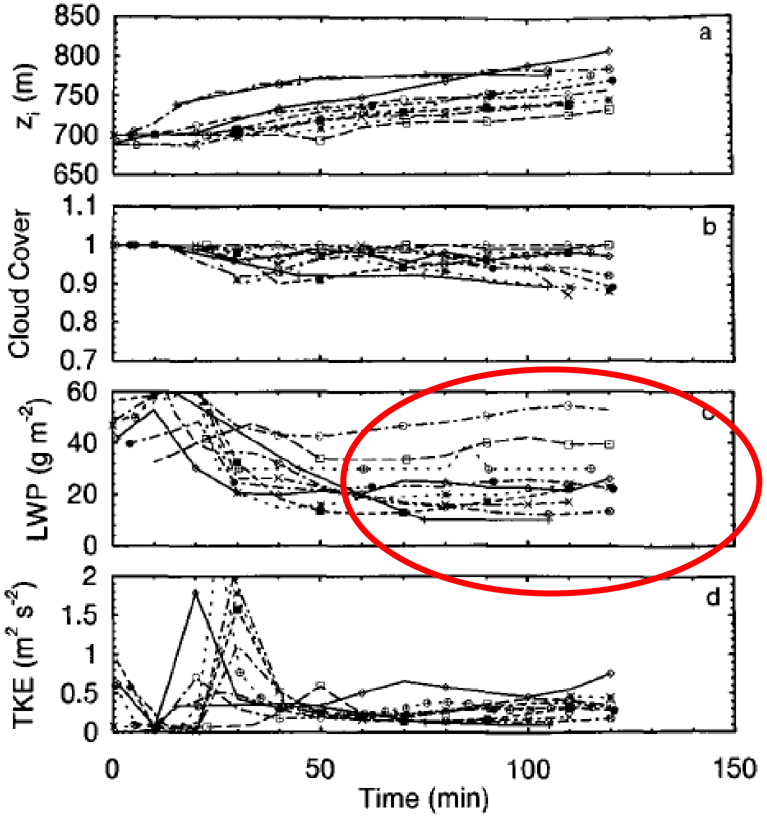
DYCOMSII	Nocturnal Scu	2003
DYCOMSII	Nocturnal Scu Precipitating	2005

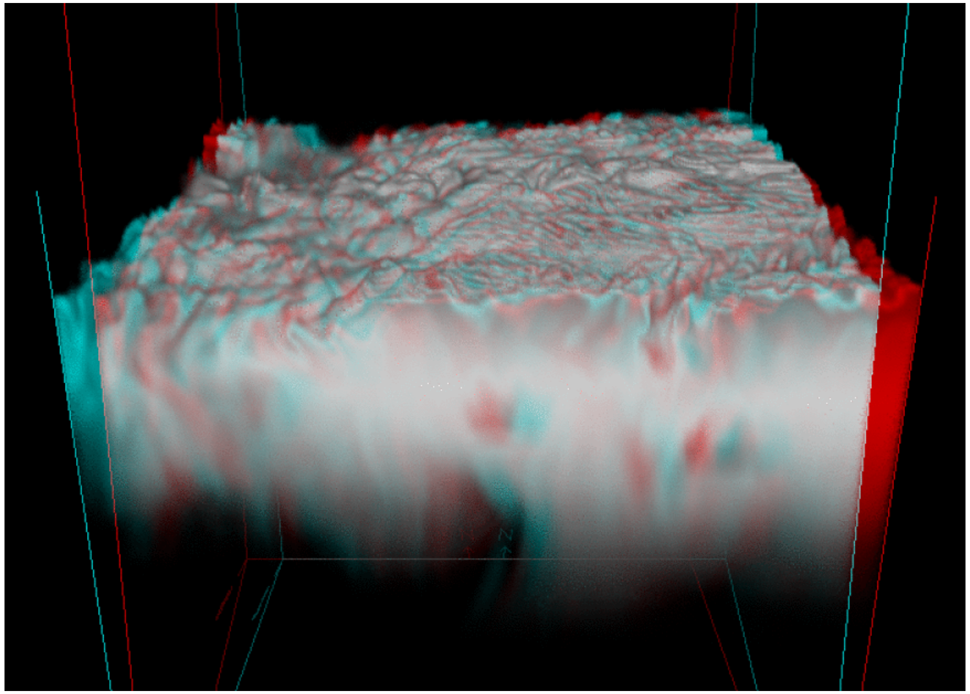
Spread of LWP in LES too large to constrain SCM's and parameterizations due to :

- case not well constrained.
- Numerics and resolution of the LES models not good enough to deal with strong inversion.

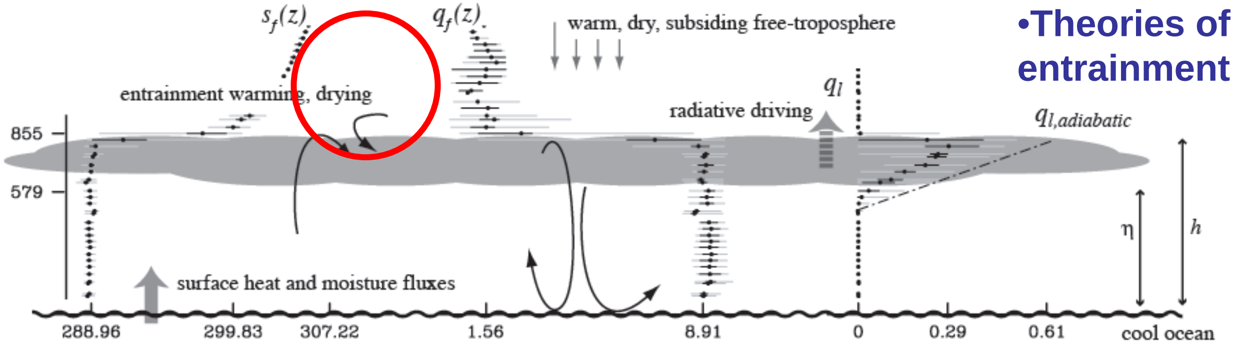
LES Results (first case 1994)



Stratocumulus (4)



Courtesy: Steve Krueger



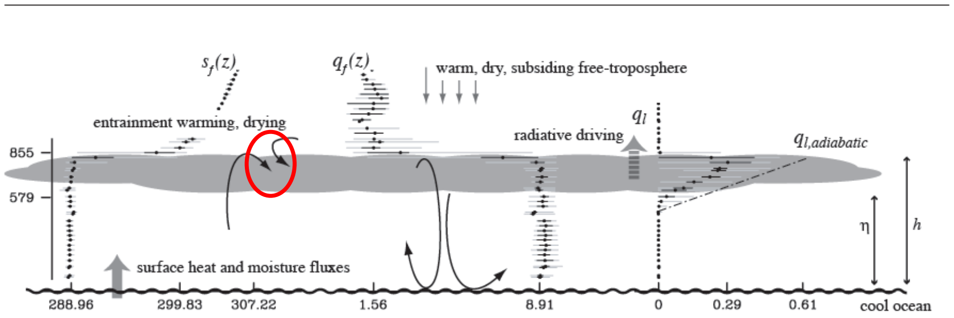
Era of maturing (1995-2002):

- Better constraint cases
- Improved advection schemes for LES
- Higher Resolution.

Making of the theory and Parameterizations:

- Identification of top-entrainment as a key process
- Theories and parameterizations of entrainment.
- Theories of decoupling of Scu./ cloud-top entrainment instability (Randall 1980)

Stratocumulus : Top-entrainment (1)



Computation of the flux $\overline{w'\psi'} = -K_\psi \frac{\partial \overline{\psi}}{\partial z}$

$$\psi \in \{q_t, \theta_l\}$$

Representation of entrainment rate w:

$$\overline{w'\psi'}_e = w_e \Delta\psi$$

w_e from parametrization

Analogous to the dry PBL: $w_e = A\frac{w_* \theta_{v,*}}{\Delta\theta} \approx \frac{\text{"energetics"}}{\text{"jump"}}$

In Scu many more parameters enter into the energetics:.

Surface moisture flux.

Surface sensible heat flux.

Condensation/evaporation processes.

Long-wave radiative cooling.

Temperature and humidity jumps at inversion

No lack of rules/parameterizations of the entrainment velocity

- Nicholls and Turton (1986)

$$w_e = \frac{2.5AW_{NE}}{\Delta\theta_{v,NT} + 2.5A\left(T_2\Delta\theta_{v,dry} + T_4\Delta\theta_{v,sat}\right)}$$

- Lilly (2002)

$$w_e = \frac{A_{DL}W_{NE,DL}}{\Delta\theta_{v,DL} + A_{DL}\left(L_2\Delta\theta_{v,dry} + L_4\Delta\theta_{v,sat}\right)}$$

- Stage and Businger (1981)
  Lewellen and Lewellen (1998)
  VanZanten et al. (1999)

$$w_e = \frac{AW_{NE}}{T_2\Delta\theta_{v,dry} + T_4\Delta\theta_{v,sat}}$$

- Lock (1998)

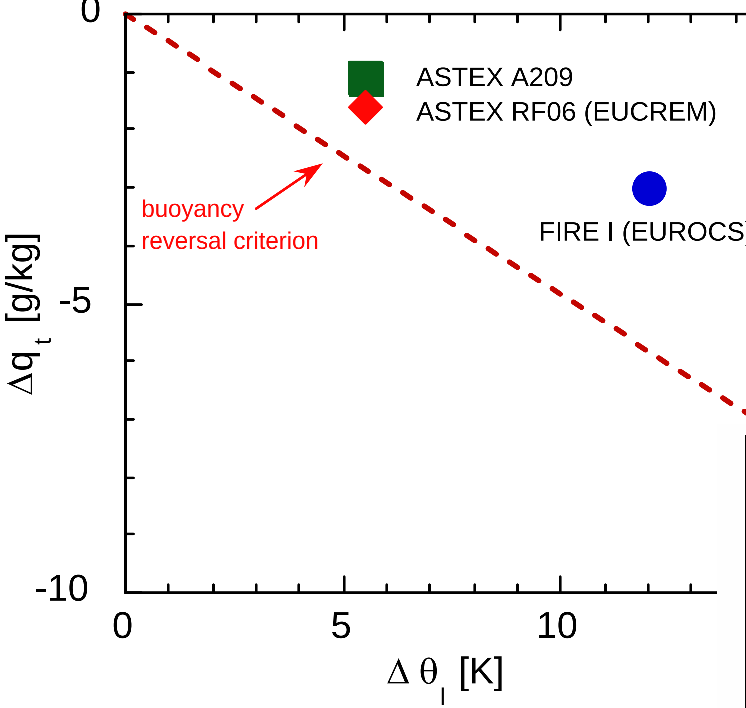
$$w_e = \frac{2A_{AL}W_{NE} + \alpha_t A_W \Delta F_L/(\rho c_p)}{\Delta\theta_v}$$

- Moeng (2000)

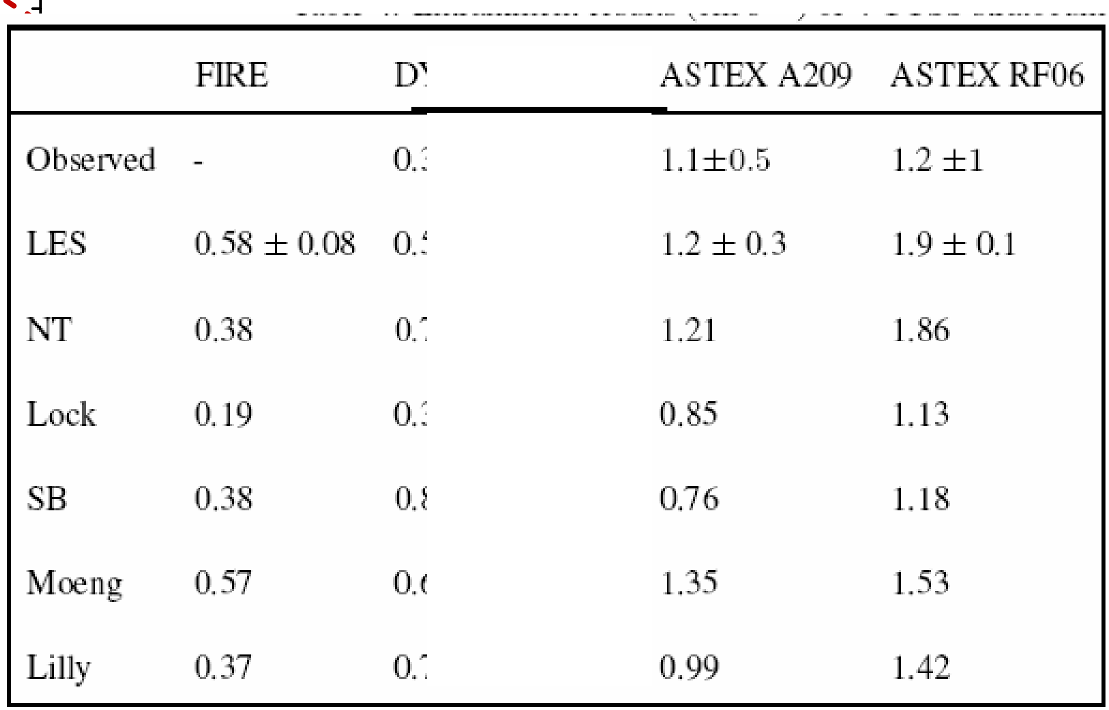
$$w_e = \frac{A_M\overline{w'\theta_l'} + \Delta F_L\left(3 - e^{-\sqrt{b_m L}}\right)/(\rho c_p)}{\Delta\theta_l}$$

Stratocumulus : Entrainment velocities: Observations vs Parameterizations

initial jumps for different GCSS stratocumulus cases

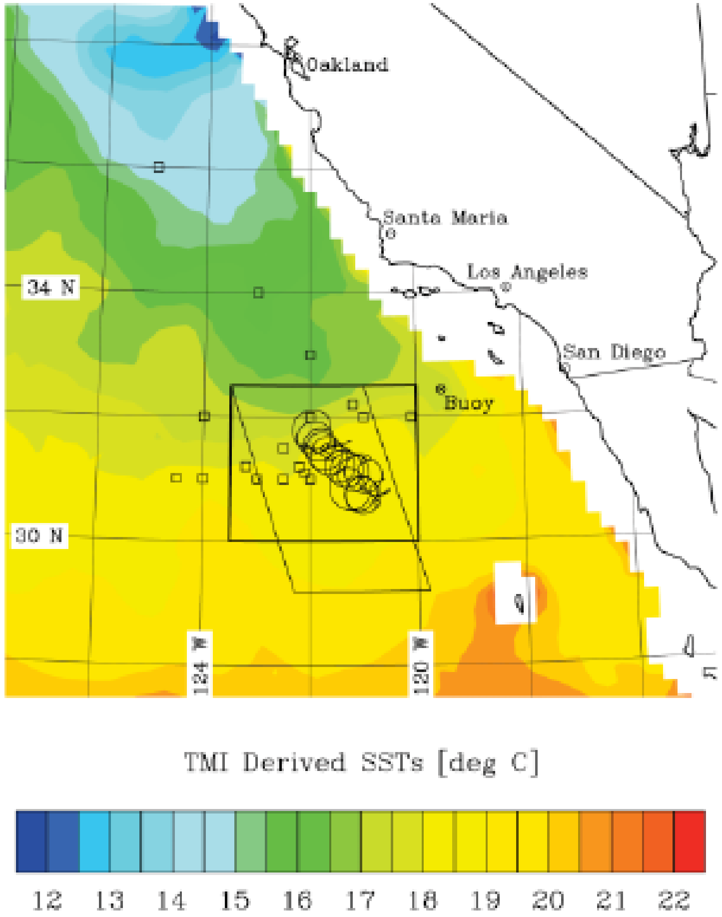


Entrainment velocities (cm/s) of 3 GCSS Cases



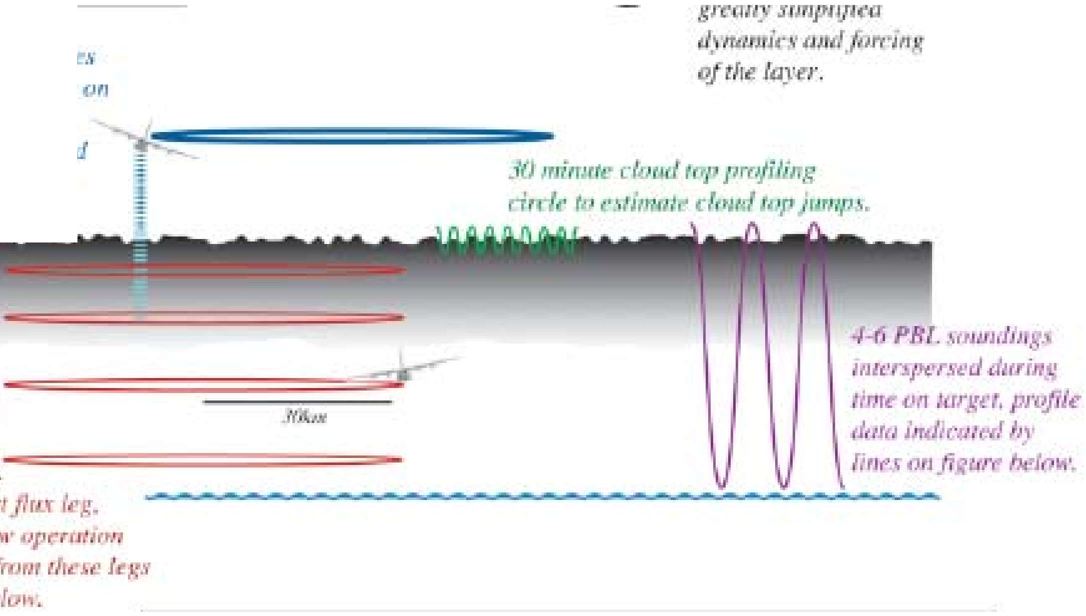
	FIRE	DY	ASTEX A209	ASTEX RF06
Observed	-	0.	1.1±0.5	1.2 ±1
LES	0.58 ± 0.08	0.	1.2 ± 0.3	1.9 ± 0.1
NT	0.38	0.	1.21	1.86
Lock	0.19	0.	0.85	1.13
SB	0.38	0.	0.76	1.18
Moeng	0.57	0.	1.35	1.53
Lilly	0.37	0.	0.99	1.42

Uncertainty in entrainment rate has inspired the GCSS-community to design a special dedicated field experiment to narrow down the uncertainty of this key process



DYCOMS II

B. Stevens et al. BAMS 84 (2003)



greatly simplified dynamics and forcing of the layer.

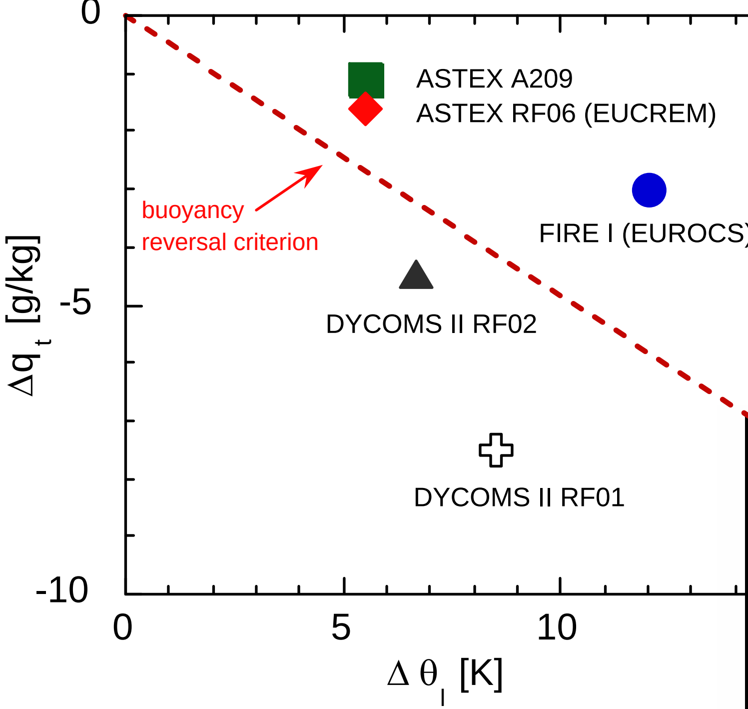
30 minute cloud top profiling circle to estimate cloud top jumps.

4-6 PBL soundings interspersed during time on target, profile data indicated by lines on figure below.

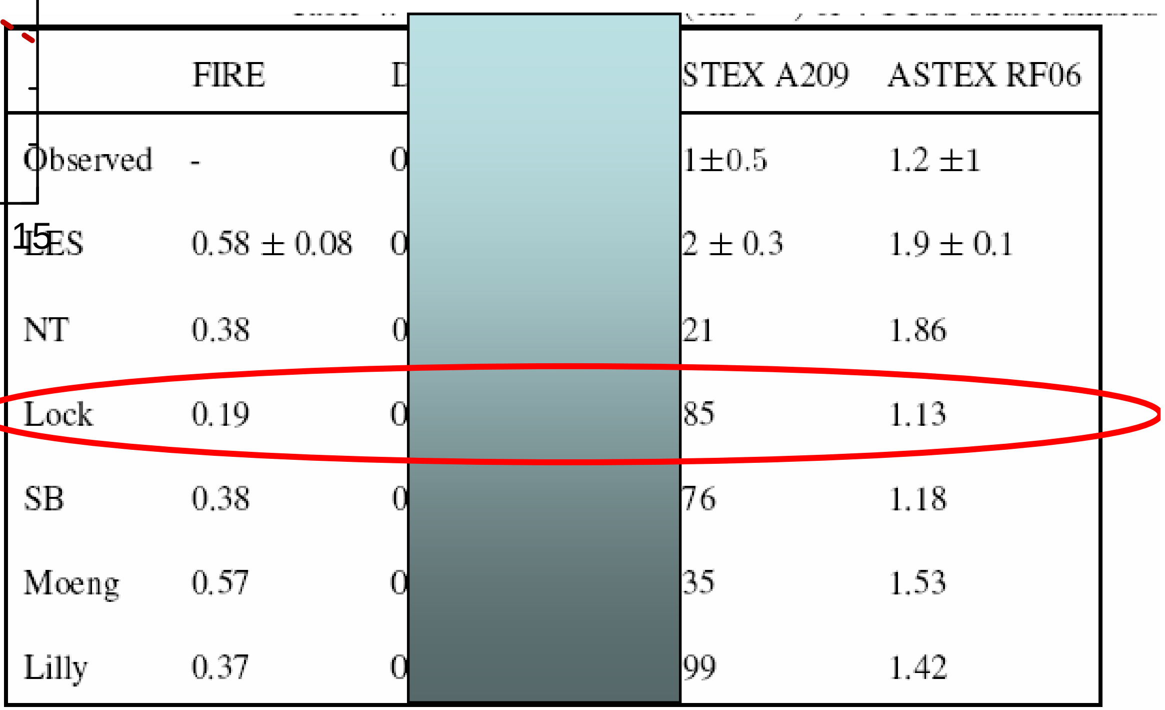
30 minute circles flown in alternating directions are used to estimate divergence and vorticity. Flights timed so surface leg is last flux leg, just around sunrise, so as to allow operation of aircraft nearer surface. Data from these legs indicated by red dots on figure below.

Incorporating DYCOMS results: narrowing down parametrizations!

initial jumps for different GCSS stratocumulus cases



Entrainment results (cm/s) of 4 GCSS Cases



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Did it made a difference?

Yes, especially for those operational centres that actively participated in this process: i.e. ECMWF, Met. Office, Meteo France.

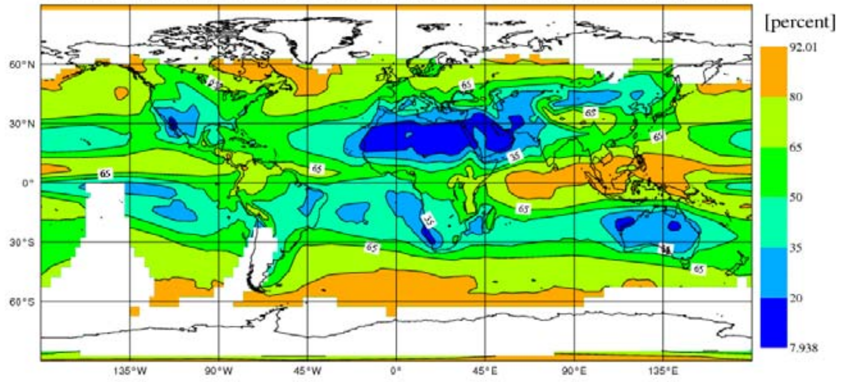
Example:

ECMWF: cloud fraction climatology

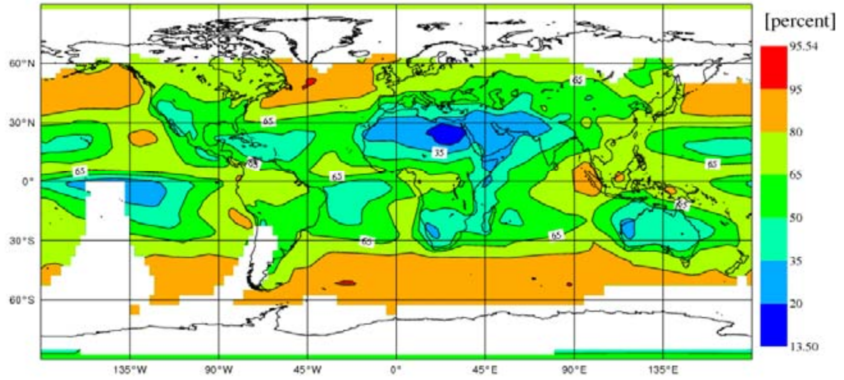
2002: underestimation of Scu

(general GCM-problem)

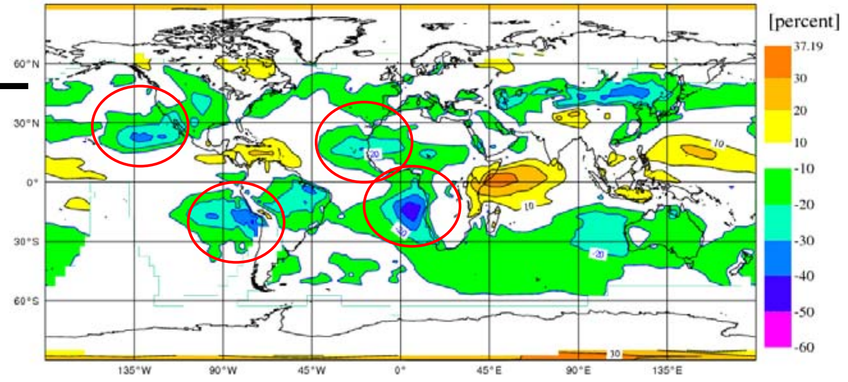
Total Cloud Cover ej3z September 2000 nmonth=12 nens=3 Global Mean: 57.9 50N-S Mean: 55.6



Total Cloud Cover ISCCP D2 September 2000 nmonth=12 50N-S Mean: 62.2



Difference ej3z - ISCCP 50N-S Mean err -6.6 50N-S rms 13.1



model - obs

9/2/2008

Courtesy: Martin Kohler

Did it made a difference?

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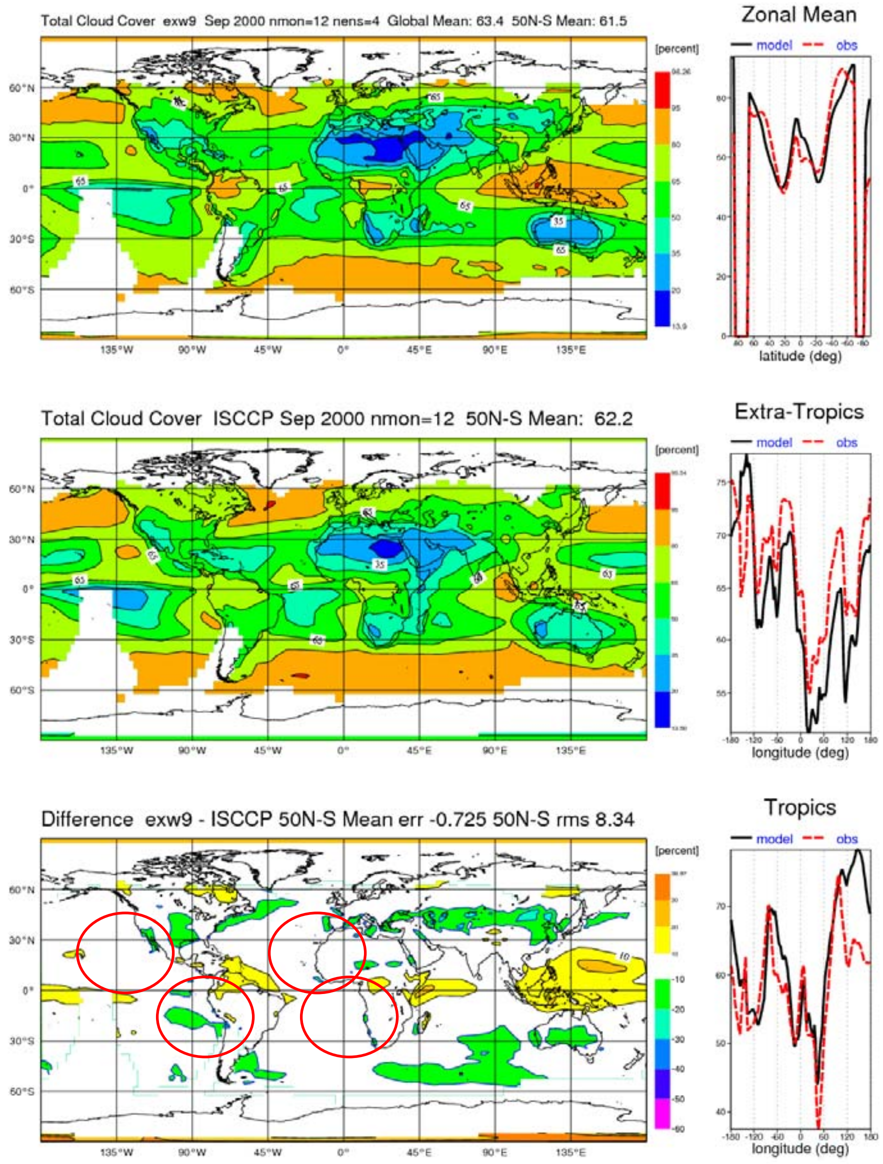
Example:

ECMWF: cloud fraction climatology

2007: Scu underestimation problem resolved.

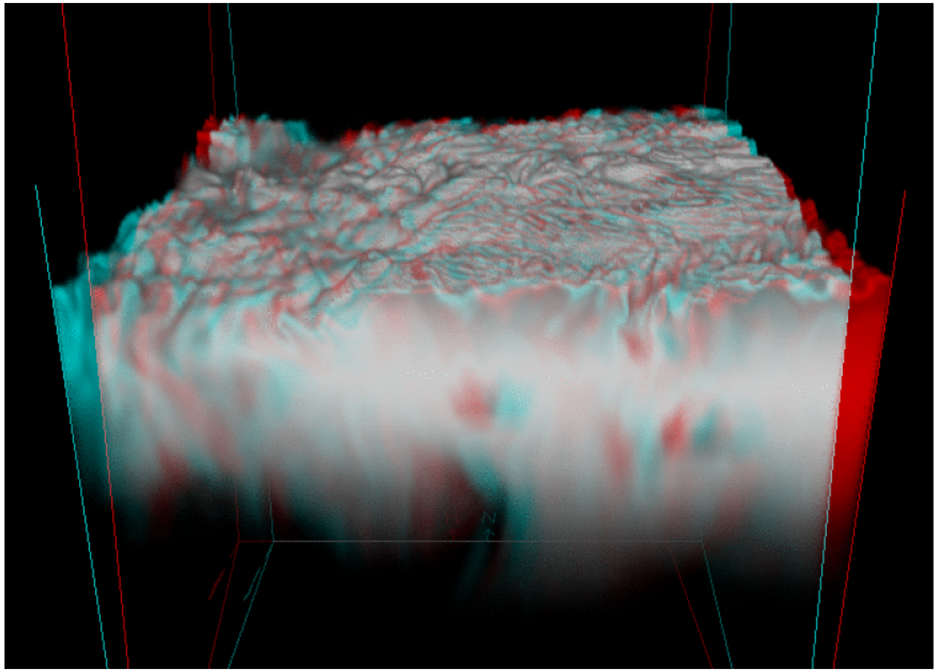


model - obs ⟵

Courtesy: Martin Kohler

Lessons to be learned!!

- **use observations and models to identify the weak spots** **(top-entrainment)**
- **advance theories to improve representation** **(entrainment closures)**
- **design critical field experiments** **(DYCOMS)**
- **Implement the findings in Large-scale models** **(ECMWF)**
- **Critically evaluate the result on a global scale** **(ISSCP,CERES,SSMI)**



LES results

S. Krueger, Univ of Utah

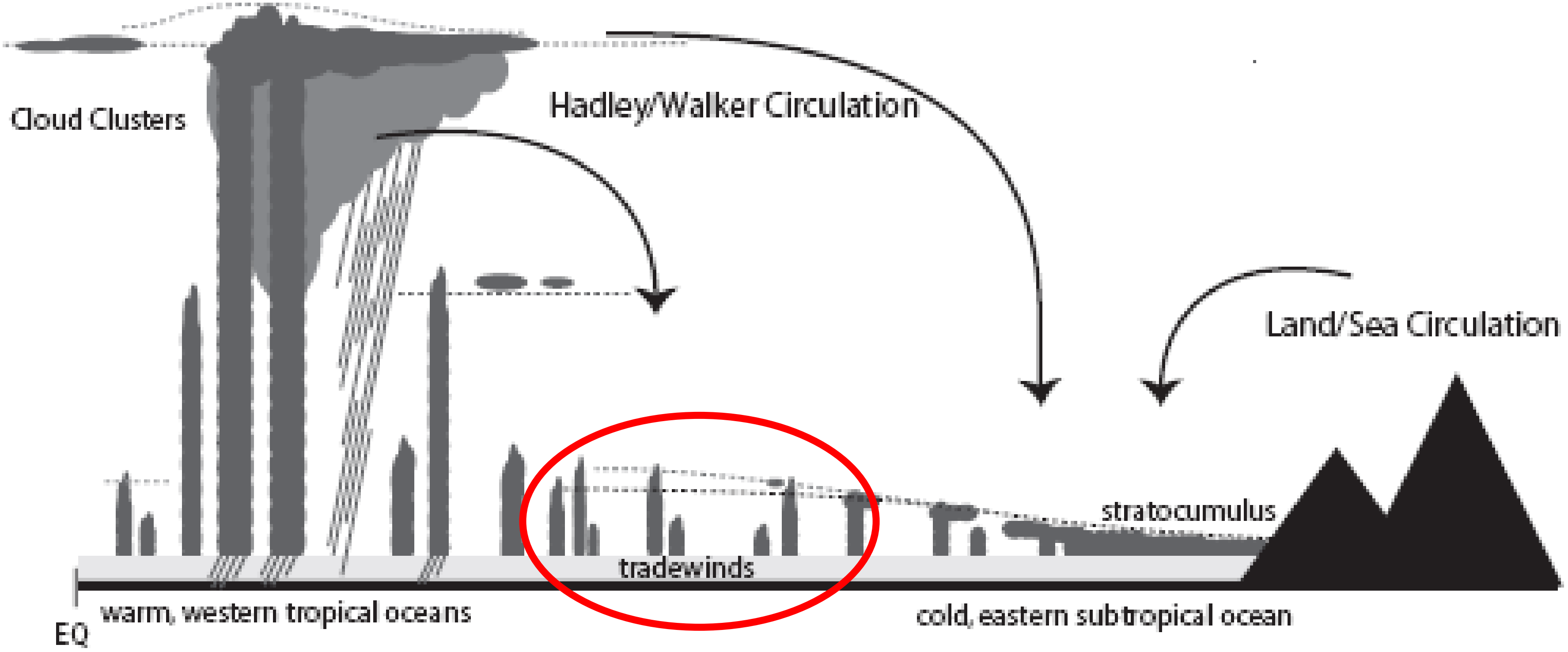
Conclusions (stratocumulus)

- Mixing in Scu should be done in moist conserved variables

- Key problems : Regime changes : Break up of Scu / decoupling

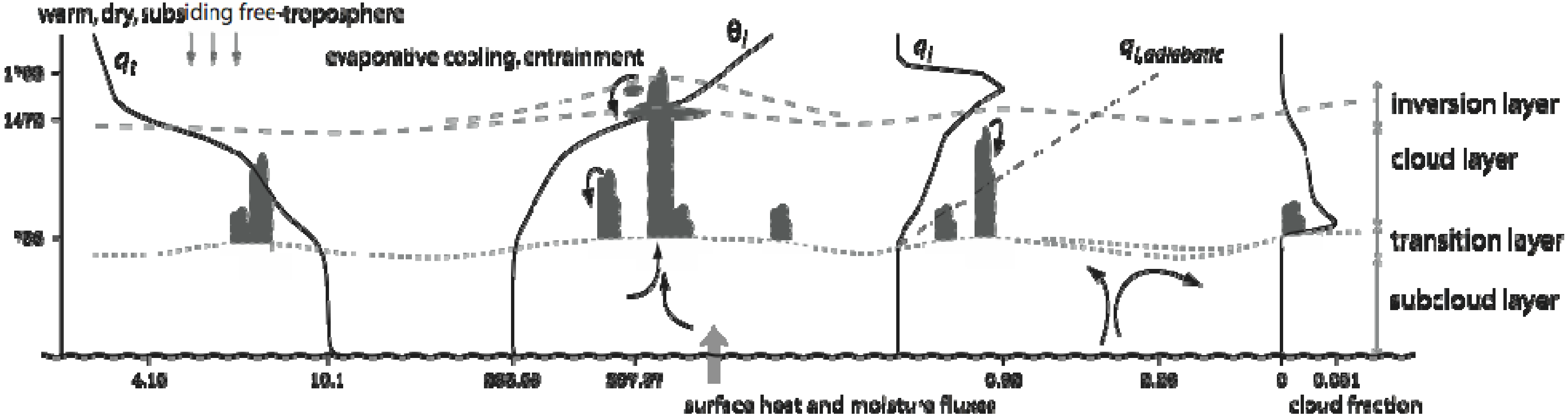
- For higher(vertical) resolution (dz~100m), TKE-schemes without explicit top-entrainment seem to be an acceptable alternative for parameterizations with explicit top-entrainment parameterizations.

Key Cloud-types that have been studied in GCCS



Shallow Cumulus: Characteristics

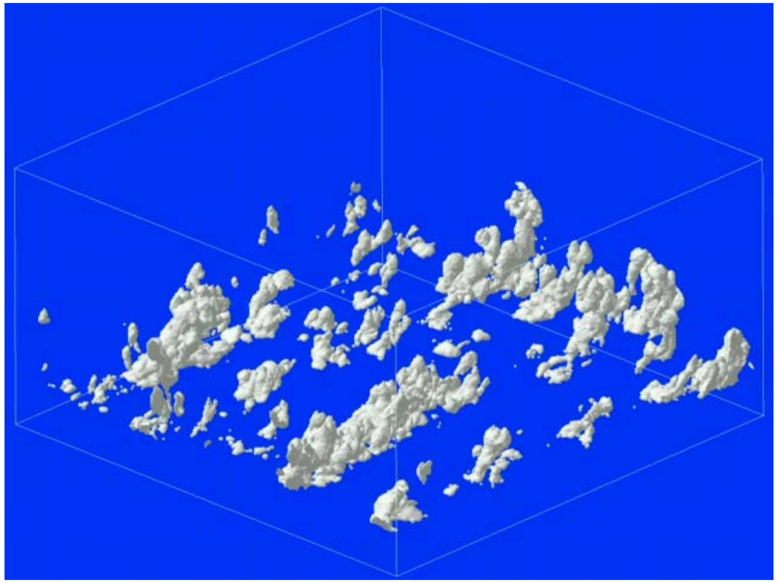
Courtesy Bjorn Stevens



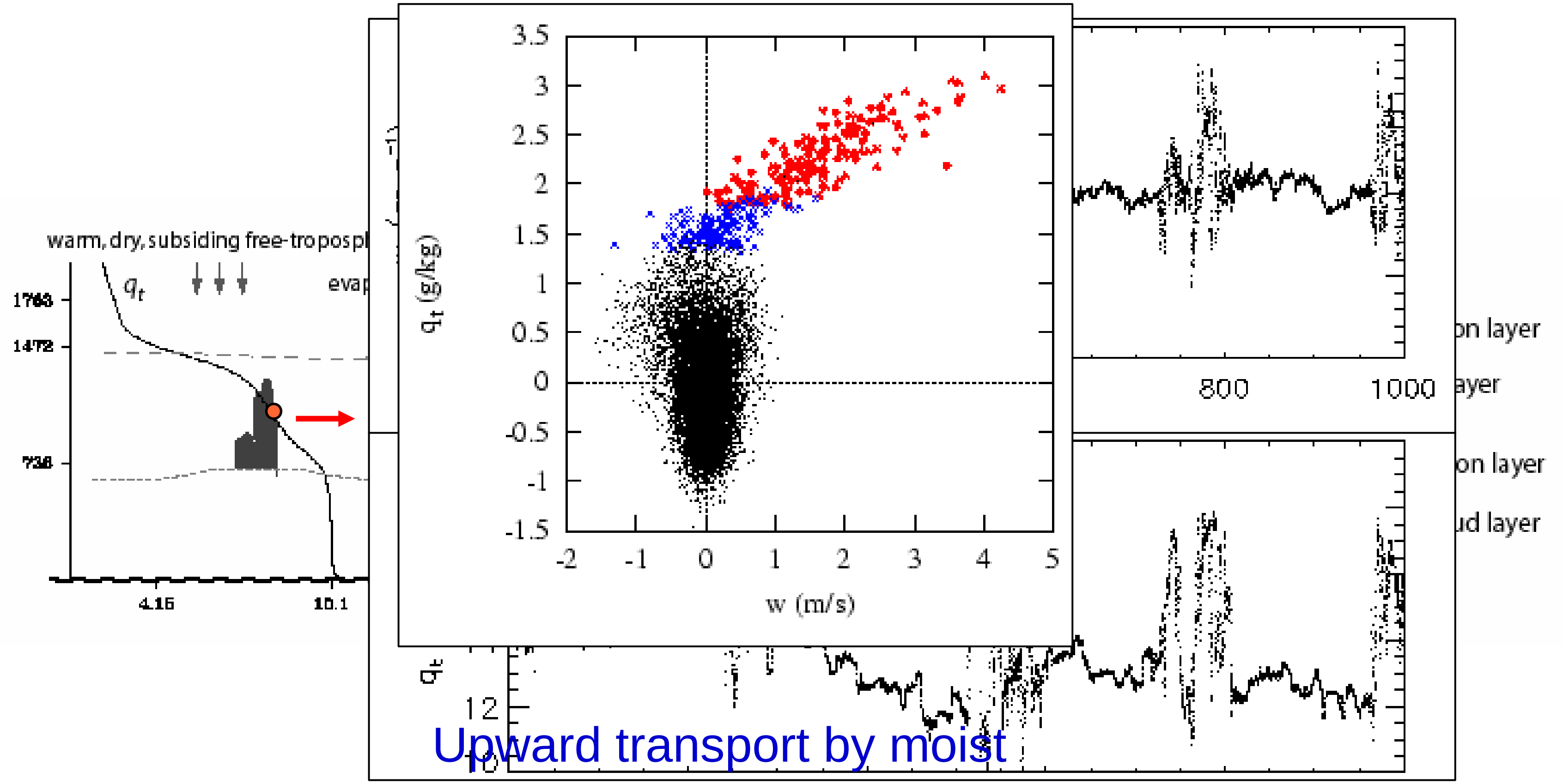
LES

Heus

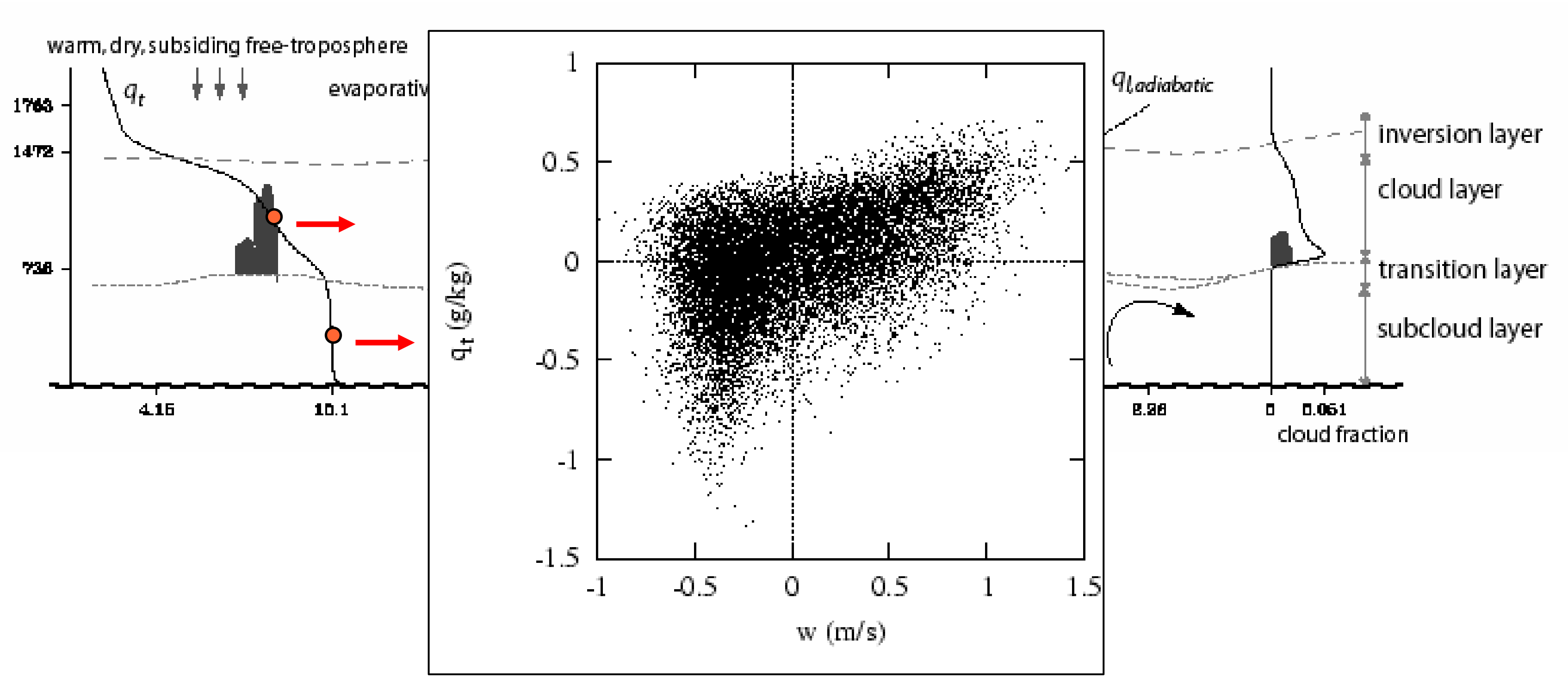
TU Delft



Horizontal Variability



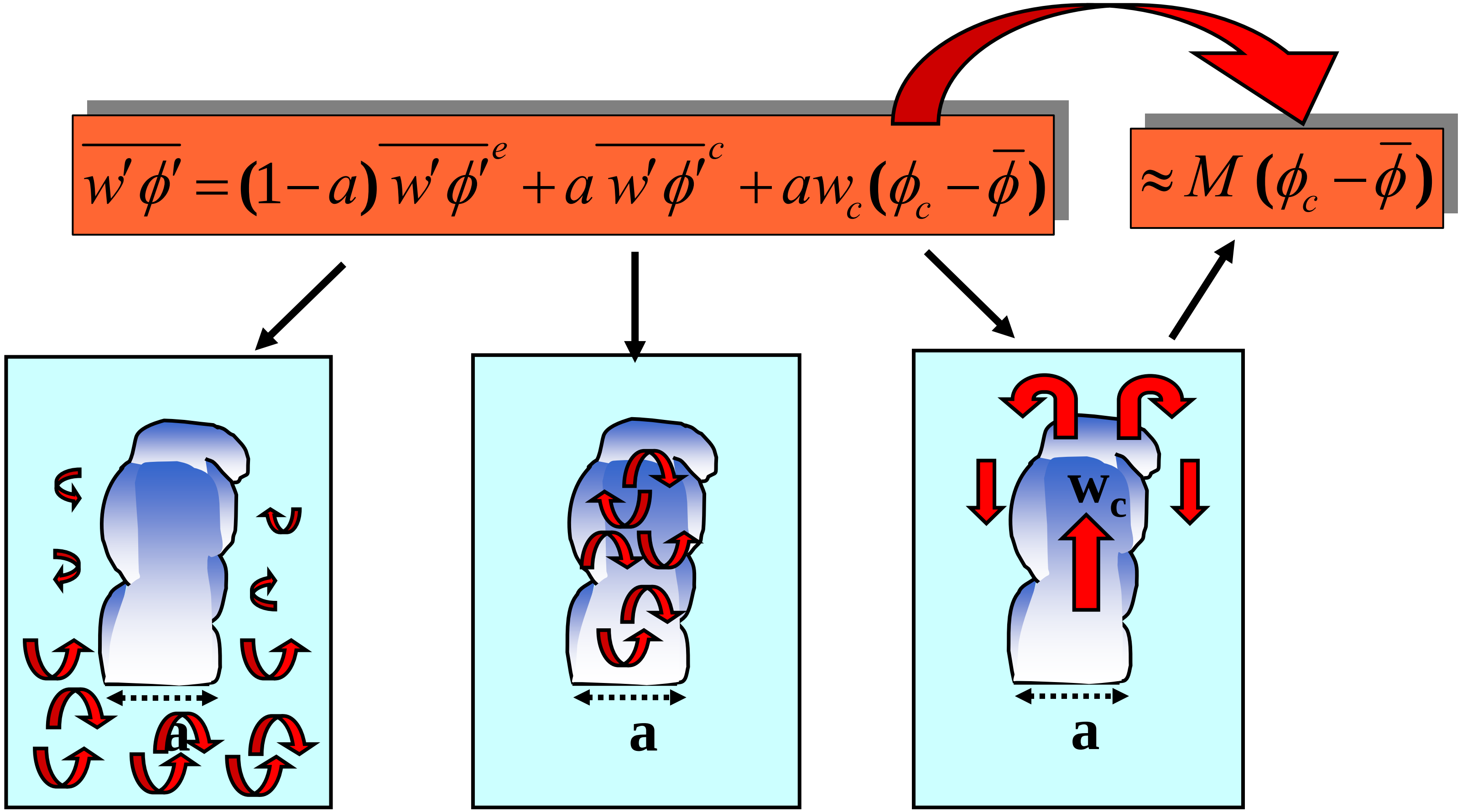
Upward transport by moist buoyant cumulus cores



less distinct updrafts in subcloud layer

Strong bimodal character of joint pdf has inspired the design of mass flux parameterizations of turbulent flux in Large scale models

(Betts 1973, Arakawa& Schubert 1974, Tiedtke 1988)



$$\overline{w'\phi'} = (1-a)\,\overline{w'\phi'}^{e} + a\,\overline{w'\phi'}^{c} + a w_c(\phi_c - \bar{\phi})$$

$$\approx M\,(\phi_c - \bar{\phi})$$

How to estimate updraft fields and mass flux?

Betts	1974 JAS
Arakawa&Schubert	1974 JAS
Tiedtke	1988 MWR
Gregory & Rowntree	1990 MWR
Kain & Fritsch	1990 JAS
And many more……..	

The old working horse:

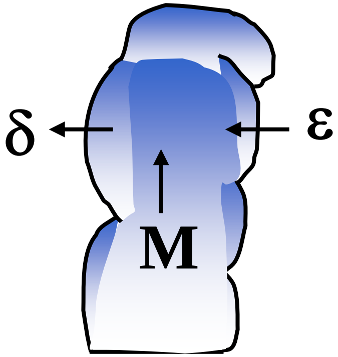
Entraining plume model:

$$\frac{\partial \phi_c}{\partial z} = -\varepsilon(\phi_c - \overline{\phi}) \text{ for } \phi \in \{\theta_l, q_t\}$$

$$\frac{1}{M}\frac{\partial M}{\partial z} = \varepsilon - \delta$$

$$\frac{1}{2}\frac{\partial w_c^2}{\partial z} = -b\varepsilon\, w_c^2 + aB, \quad B = \frac{g}{\theta_0}\left(\theta_v - \overline{\theta_v}\right)$$

••••

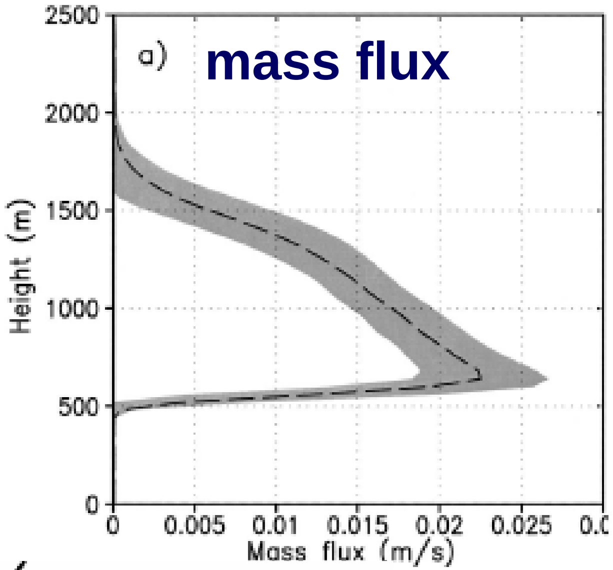


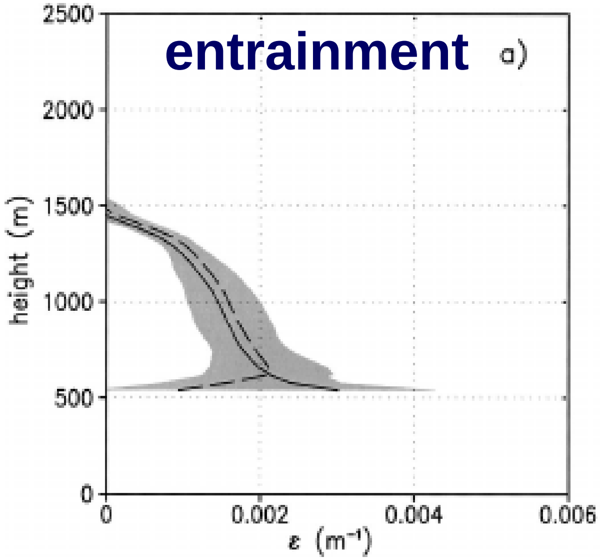
Plus boundary conditions at cloud base.

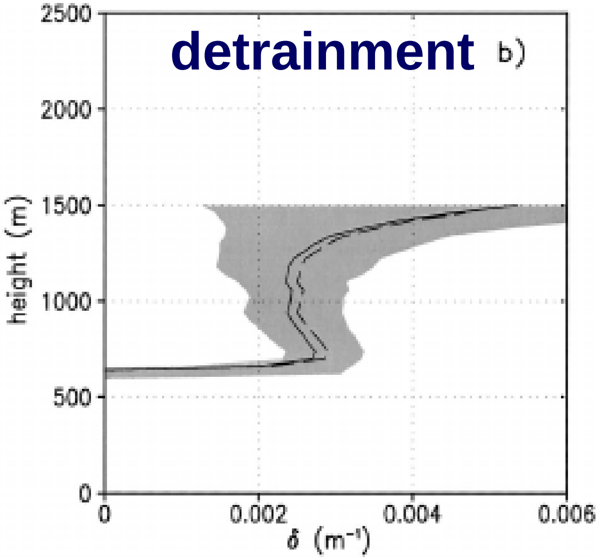
GCSS cases

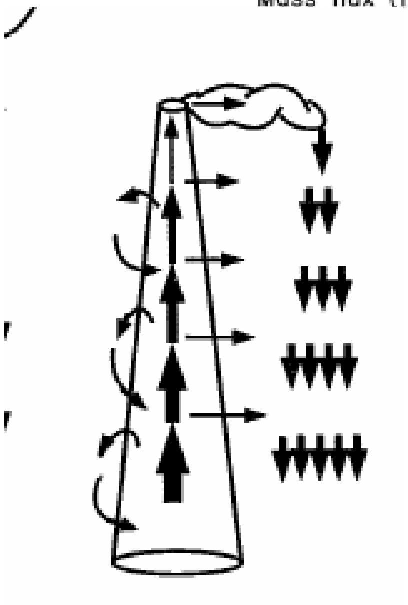
Experiment	Case	year
BOMEX	Steady state Trade wind cu	1997
ATEX	Trade wind cu topped with Scu	1998
ARM **(June 1997)**	Diurnal Cycle Cumulus	2000
RICO	Precipitating trade wind cu	2006

Typical LES results from GCSS intercomparison studies









Main Results:

1. **Lateral entrainment and detrainment rates typically of the order of 10^{-3} m^{-1}**
2. **Detrainment rates typically larger than entrainment rates or**
3. **Mass flux decreases with height**

Siebesma and Cuypers JAS 95

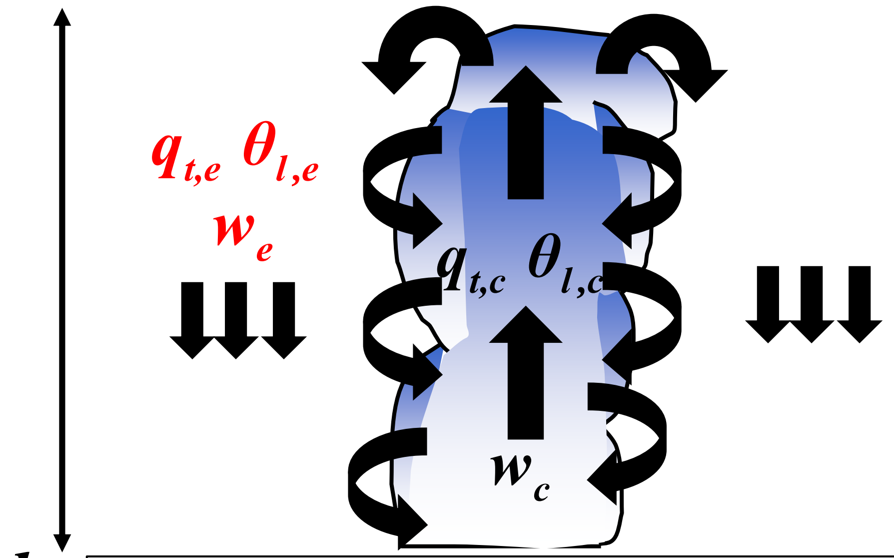
Siebesma 1998

Grant and Brown QJRMS 1999

Gregory QJRMS 2000

Neggers et al JAS 2002

Led to simple conceptual models for entrainment rates



$$for\ \phi \in \{\theta_l, q_t\}:$$

$$\frac{d\phi_c}{dt} = F_{mixing}$$

$$w_c \frac{\partial \phi_c}{\partial z} = -\frac{(\phi_c - \phi_e)}{\tau}$$

$$\boxed{\frac{\partial \phi_c}{\partial z} = -\varepsilon(\phi_c - \phi_e) \quad \text{where} \quad \varepsilon = \frac{1}{w_c \tau} \approx \frac{1}{h_c}}$$

Siebesma 1997

Bretherton and Grenier JAS 2003

Shallow convection: $h_c \sim 1000\text{m}$

$\varepsilon \sim 10^{-3}\ \text{m}^{-1}$!!

Alternative: $\boxed{\varepsilon \approx \frac{1}{w_c \tau_0} \propto \frac{1}{\text{w}_\text{c}}}$

Neggers et al 2001 JAS

Cheinet 2003 JAS

Shallow Cumulus: Lateral Detrainment Rates

- **Detrainment has received less attention than entrainment.**
- **Varies much more from case to case so is probably more important to parameterize mass flux correctly**

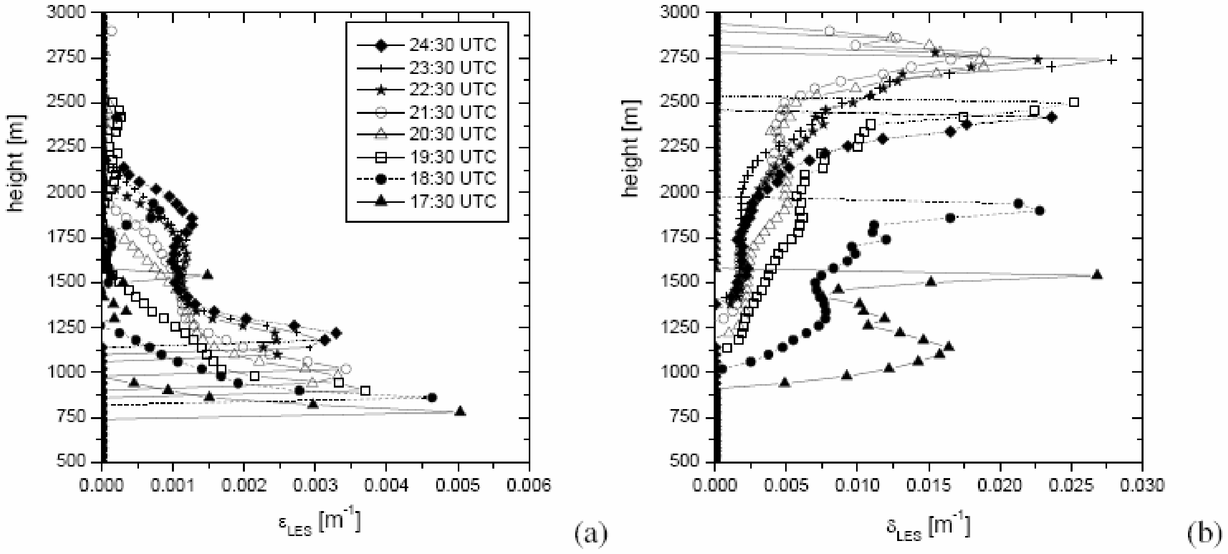
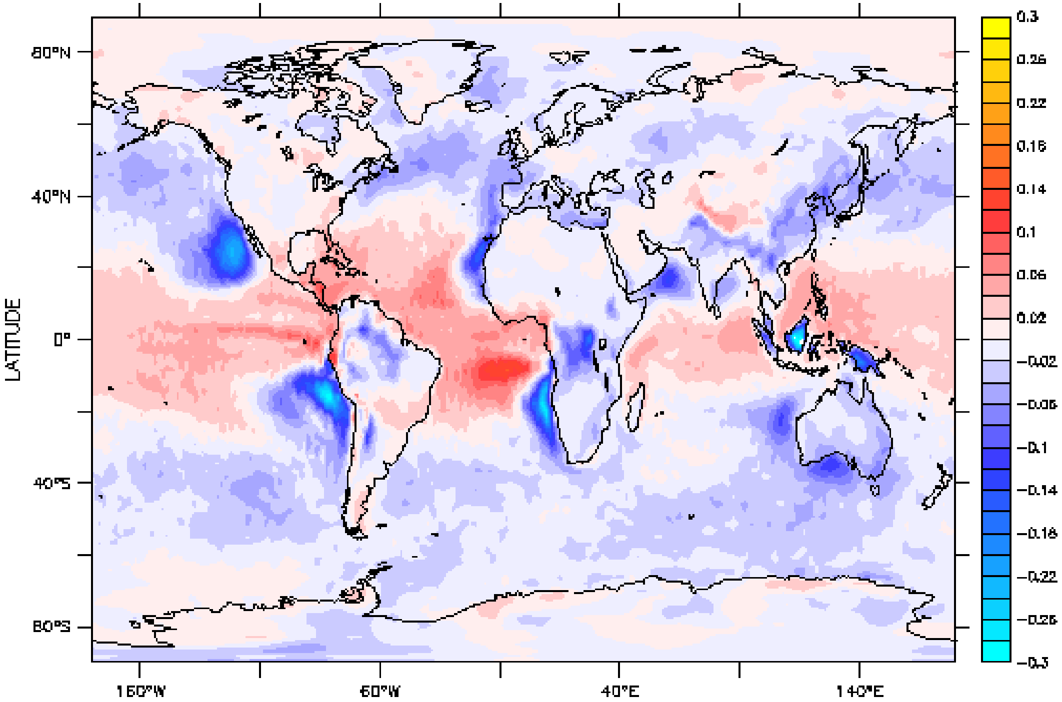


FIGURE 3: Hourly averaged fractional entrainment (a) and detrainment (b) rates diagnosed from LES results for the ARM case. Note the different x-axis scale for (a) and (b).

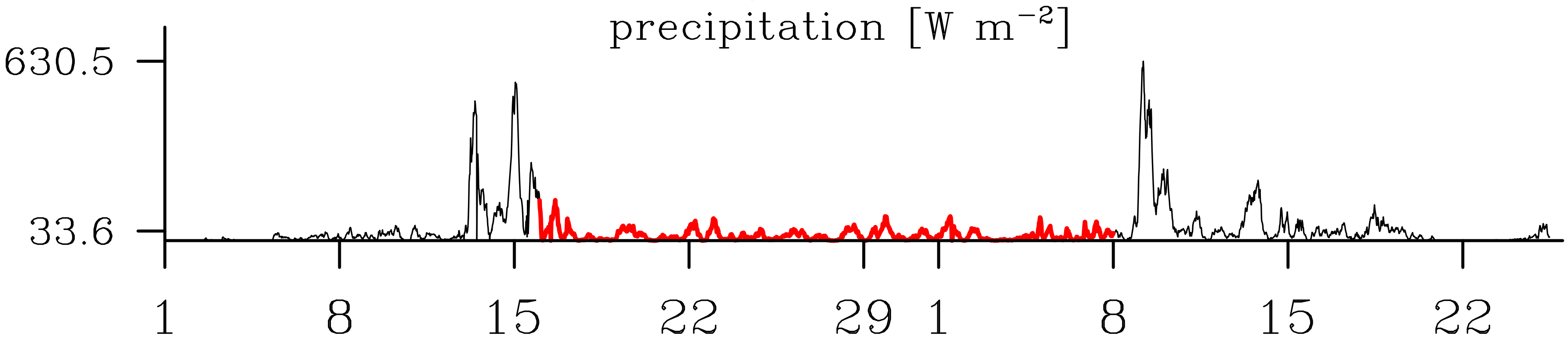
Sensitivities in ECMWF

Change in cloudcover when setting the entrainment rate of the updraft in the subcloud layer to zero:

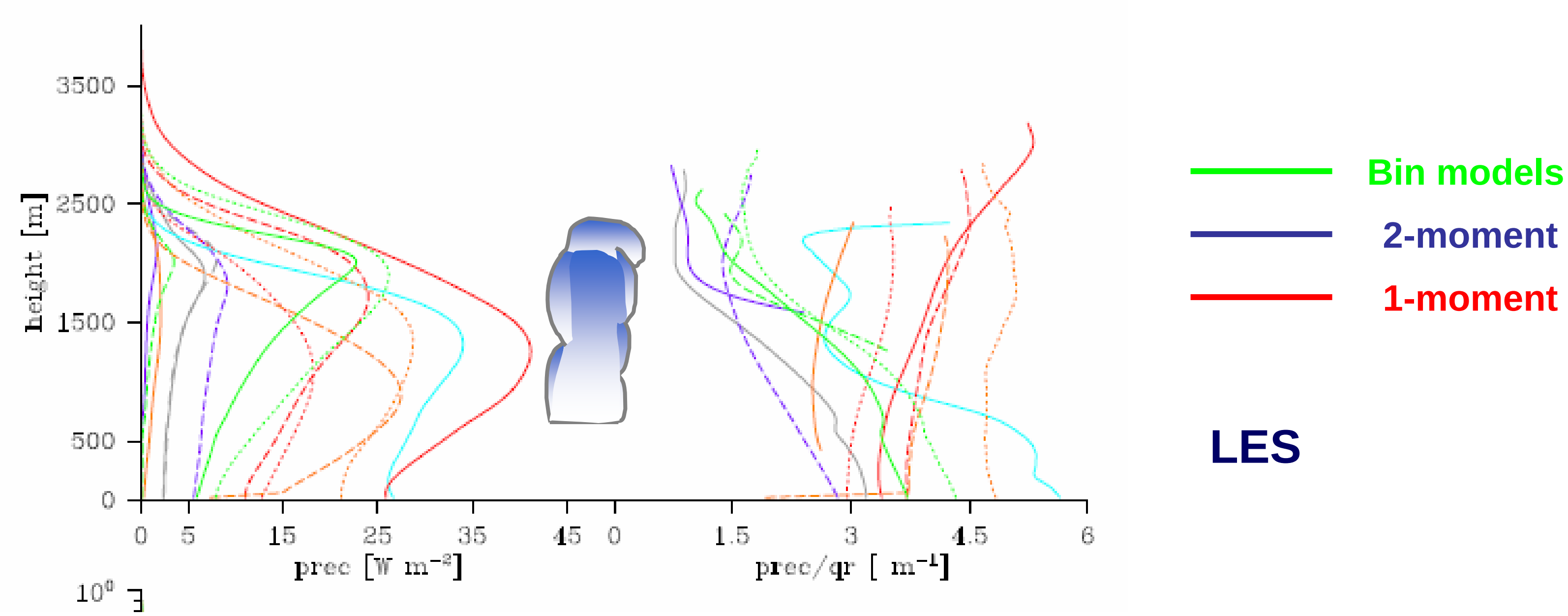


Intercomparison case based on precipitating cumulus observed during field campaign RICO:

- **Period: 2004 December 16 – 2005 January**

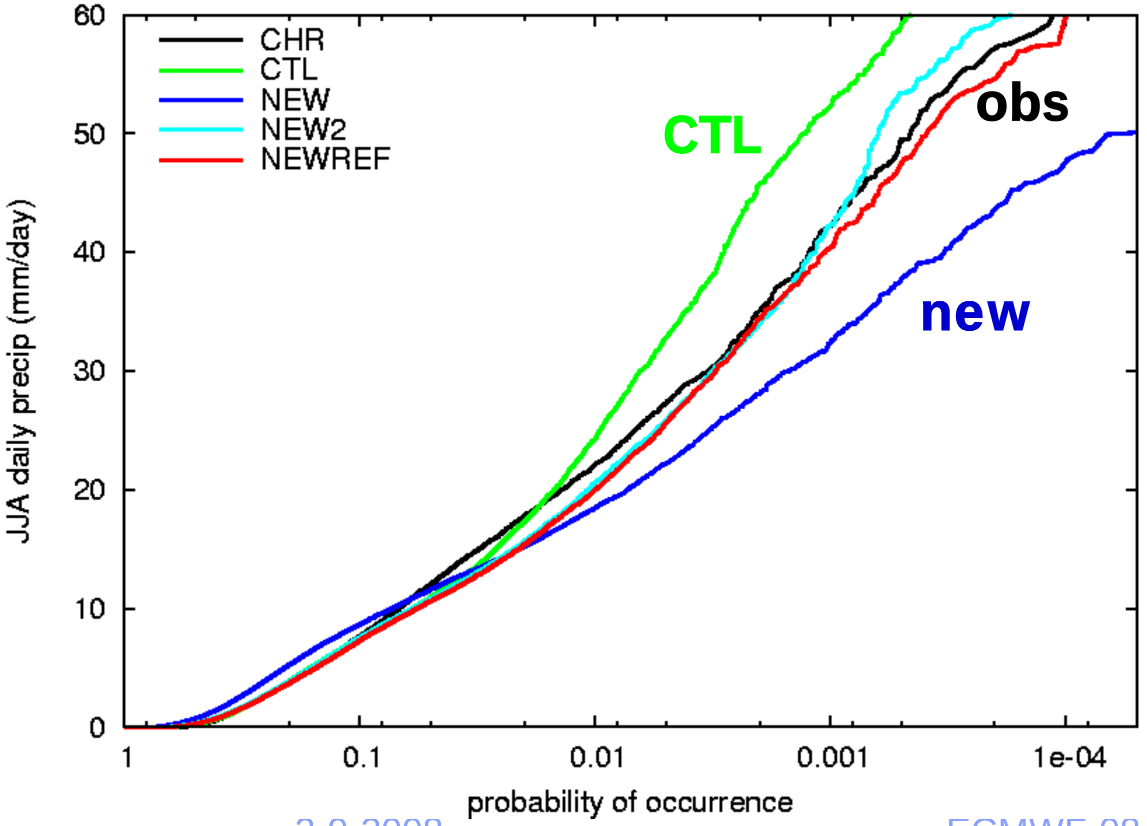


Observed precipitation rate during suppressed period : ~20W/m2 = 0.6 mm/day



Does precipitation from these shallow clouds matter?

Precipitation Histogram of JJA for 1991-1995
for the Rhine catchment area with a regional climate model
(RACMO) (25km resolution)



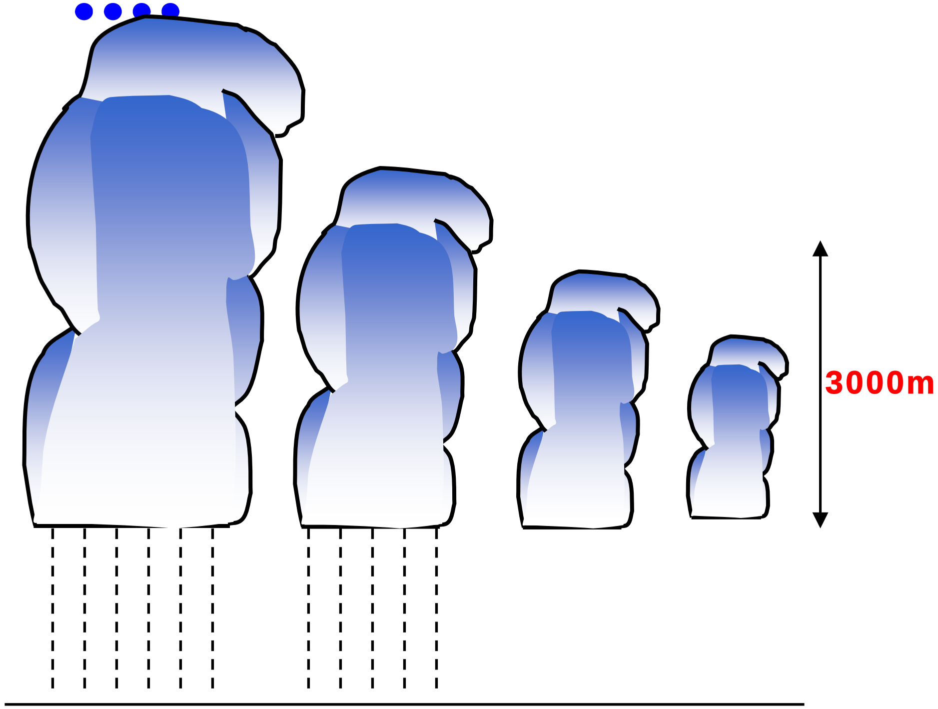
Ctl (23r4) :

- **Too few low precipitation rate events.**
- **Too many high precipitation rate events**

Ctl (31r1) :

- **Too many low precipitation rate events.**
- **Too few high precipitation rate events**
- **Lower extreme events!!**

Howcome?

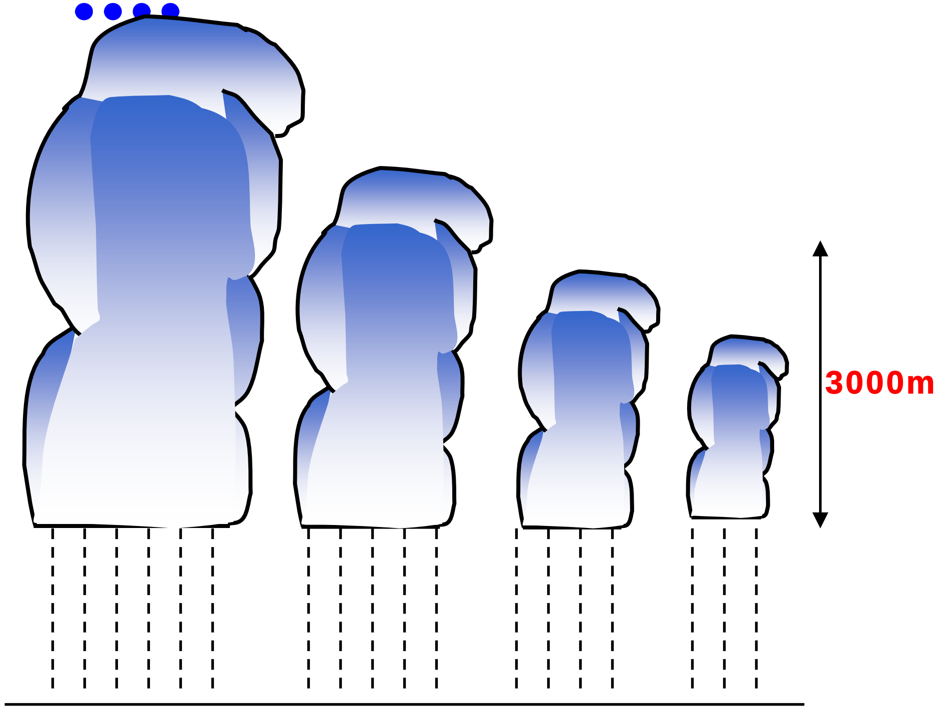


Control (23r4) :

clouds shallower than 3000m are not allowed to precipitate:

- **Obviously reduces the “moderate rain intensity events”**
- **Allows more extreme rain events to build up.**

As opposed to..........



New (25r4) :

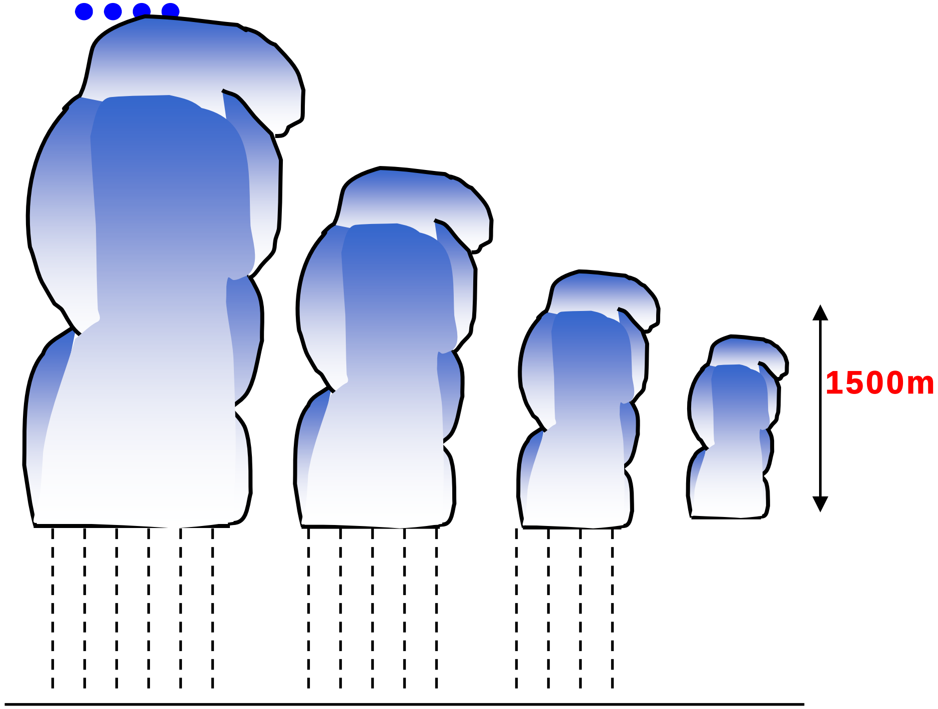
In which all clouds are allowed to precipitate (if enough ql):

- **Obviously encourages the “moderate rain intensity events”**
- **Prohibits more extreme rain events to build up.**

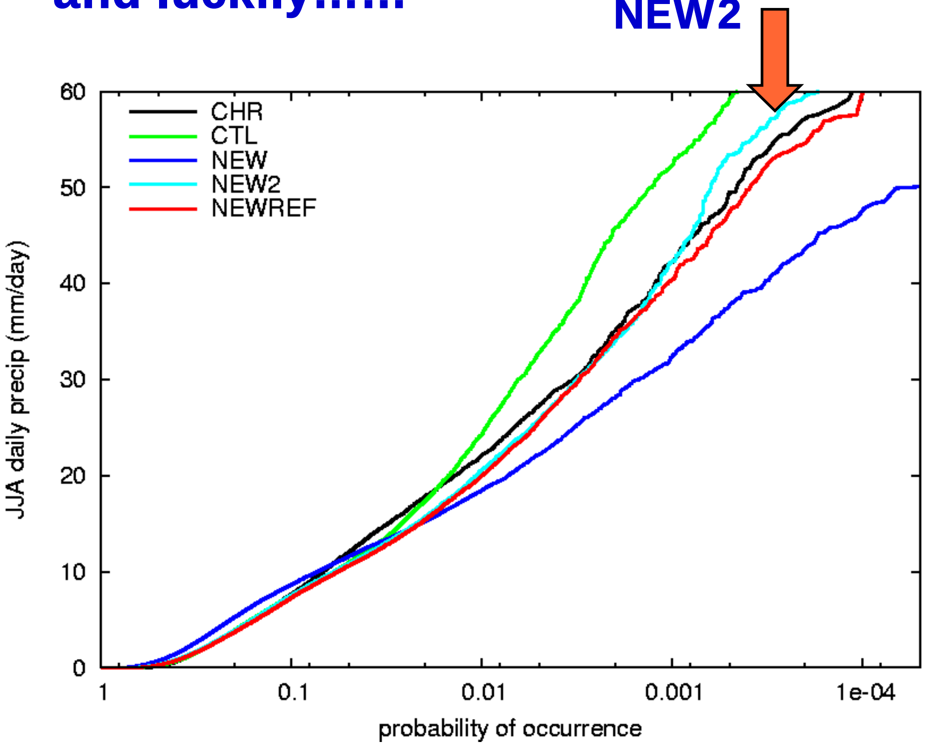
••••

So as a (temporary) fix:

.....One can prohibit clouds of 1500m to precipitate



and luckily......



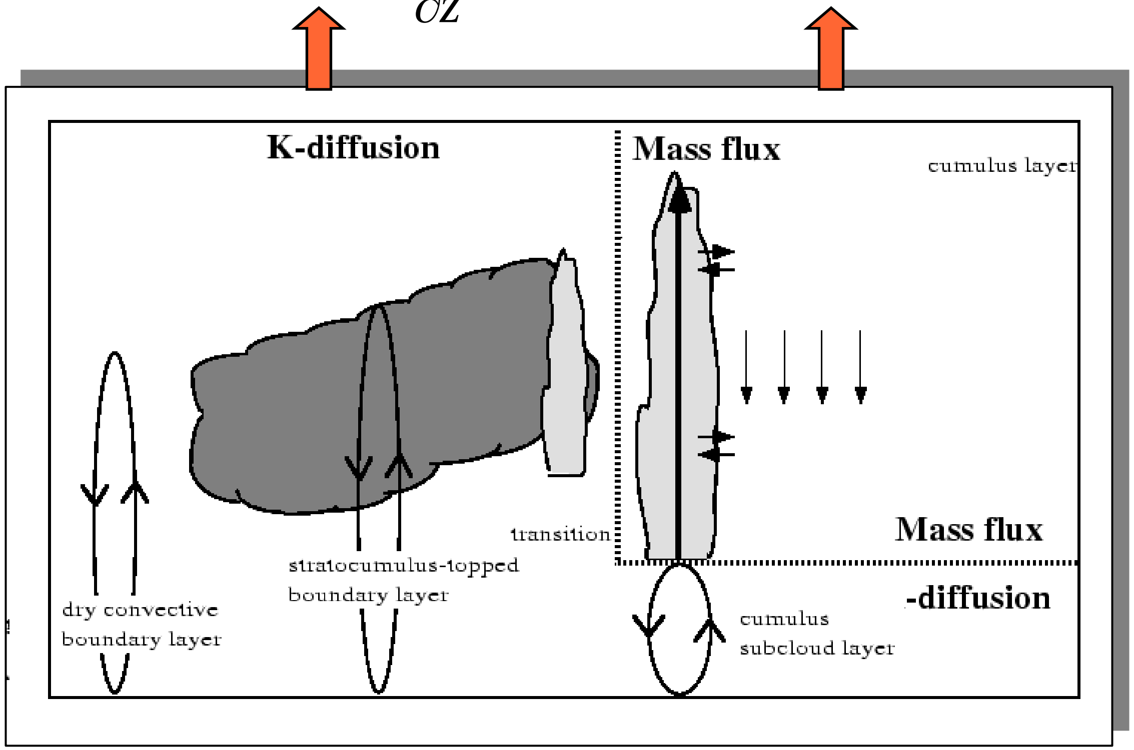
- This merely shows the sensivity of the overall precipation statistics to the precipitation efficiency of shallow clouds!!

Standard (schizophrenic) parameterization approach:

$$\overline{w'\phi'} \cong -K \frac{\partial \overline{\phi}}{\partial z}$$

$$\overline{w'\phi'} \cong M(\phi_u - \overline{\phi})$$

$$\frac{\partial \overline{\phi}}{\partial t} \cong -\frac{\partial}{\partial z}\left(\overline{w'\phi'}\right) + \overline{S}$$

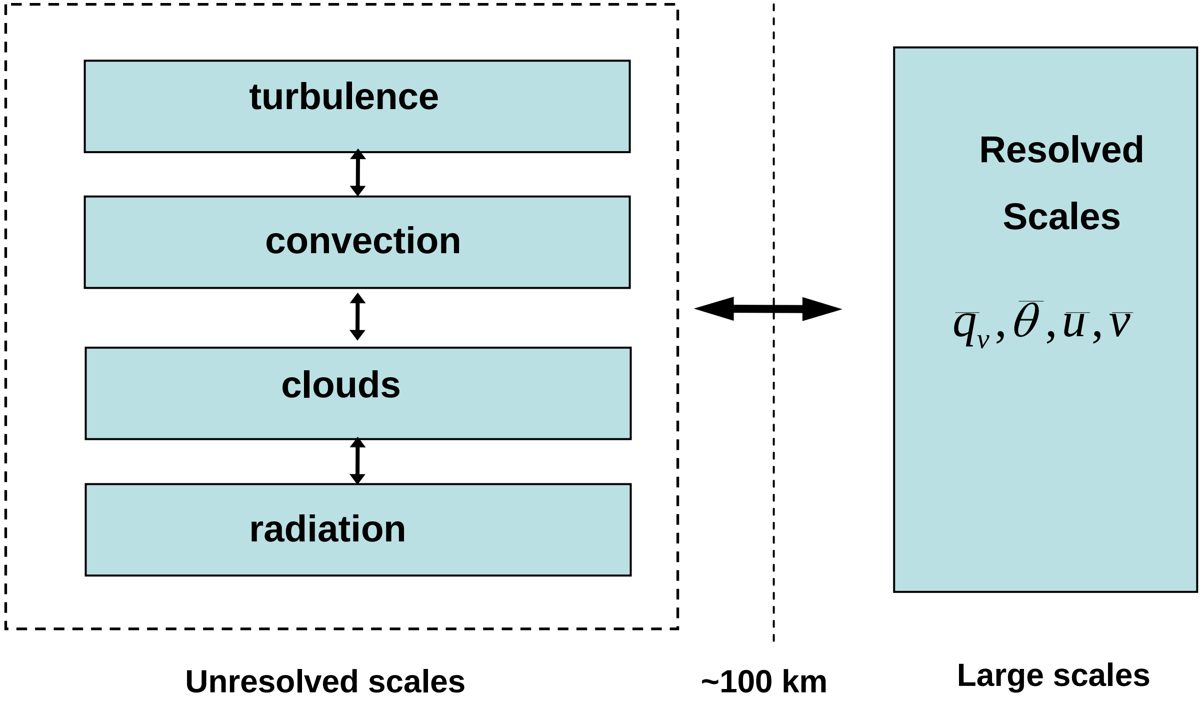


This unwanted situation can lead to:

- **Double counting of processes**
- **Inconsistencies**
- **Problems with transitions between different regimes:**
  - **dry pbl ➔ shallow cu**
  - **scu ➔ shallow cu**
  - **shallow cu ➔deep cu**

The "parameterization dish" looks perhaps a bit messy.

Intermezzo (2)



turbulence

convection

clouds

radiation

Resolved

Scales

$\overline{q}_v, \overline{\theta}, \overline{u}, \overline{v}$

Unresolved scales

~100 km

Large scales

Increase consistency between the parameterizations!

How?

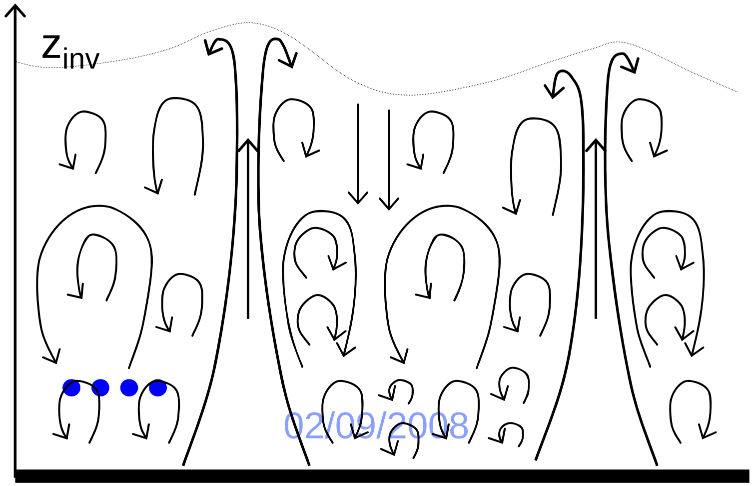
Eddy-Diffusivity/Mass Flux approach : a way out?

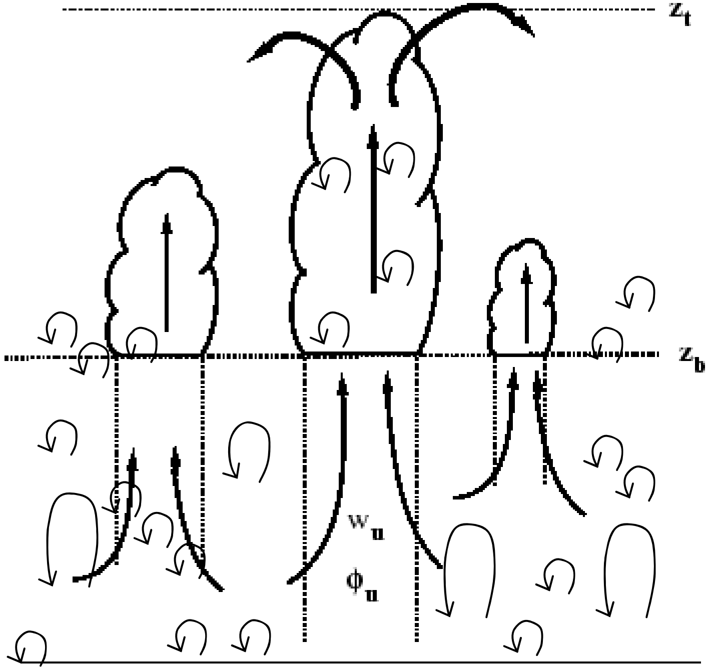
- Nonlocal (Skewed) transport through strong updrafts in clear and cloudy boundary layer by advective Mass Flux (MF) approach.

- Remaining (Gaussian) transport done by an Eddy Diffusivity (ED) approach.

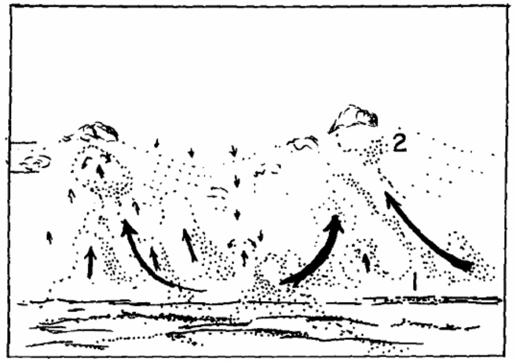
Advantages :

- One updraft model for : dry convective BL, subcloud layer, cloud layer.
- No trigger function for moist convection needed
- No switching required between moist and dry convection needed

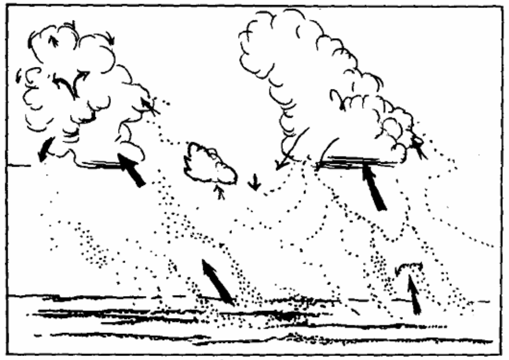




LeMone & Pennell (1976, MWR)



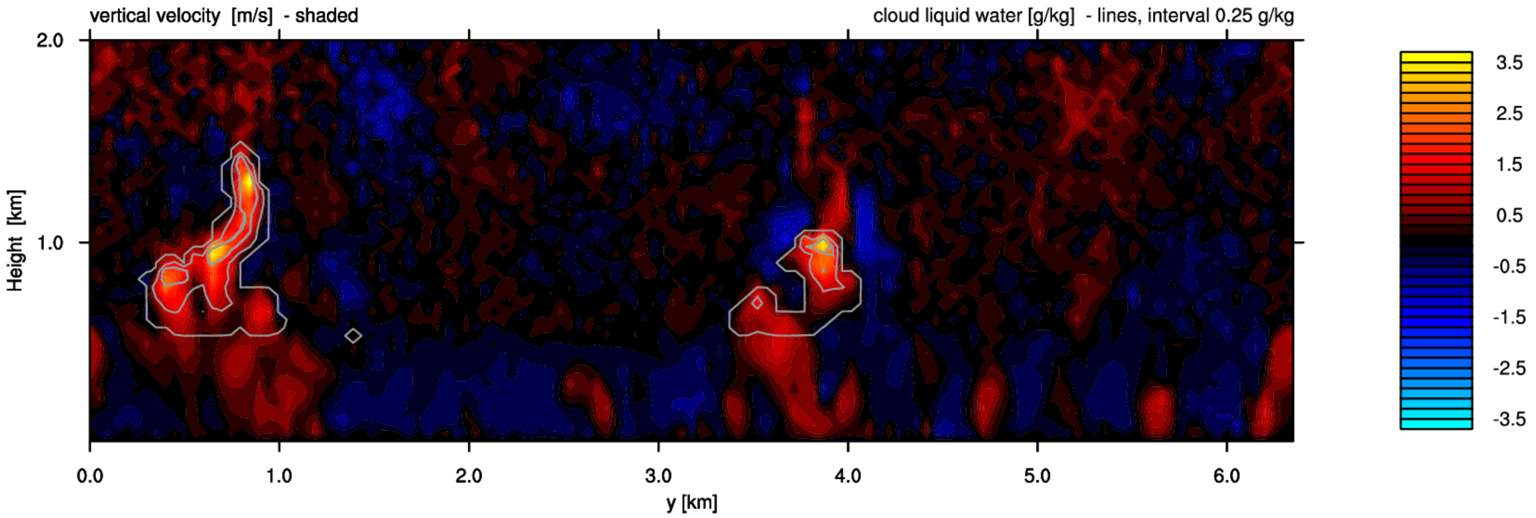
(a)



(b)

Cumulus clouds are the condensed, visible parts of updrafts that are deeply rooted in the subcloud mixed layer (ML)

LES BOMEX vertical cross-section

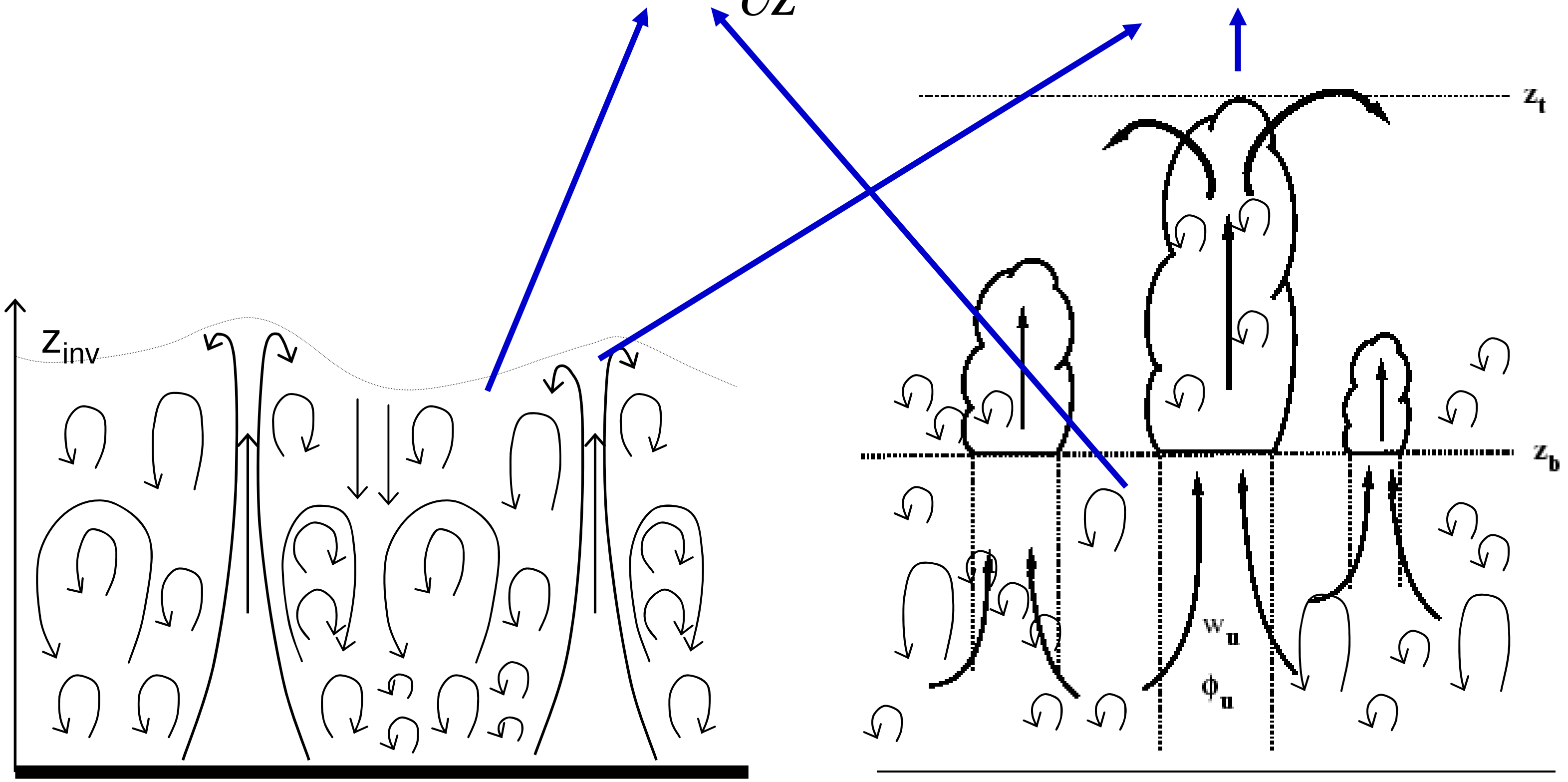


The (simplest) Mathematical Framework :

••••

$$\overline{w'\phi'} = a_u \overline{w'\phi'}^{u} + (1 - a_u)\overline{w'\phi'}^{e} + a_u w_u (\phi_u - \phi_e)$$

$$\cong \quad -K\frac{\partial \overline{\phi}}{\partial z} + \quad M(\phi_u - \overline{\phi})$$



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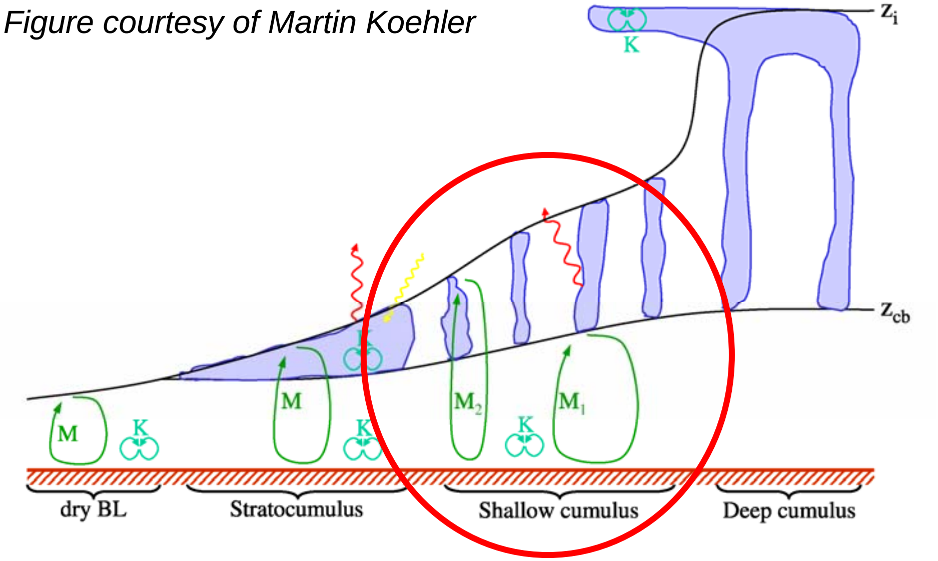
Cumulus Topped Boundary Layer

Neggers, Kohler & Beljaars accepted for JAS 2008

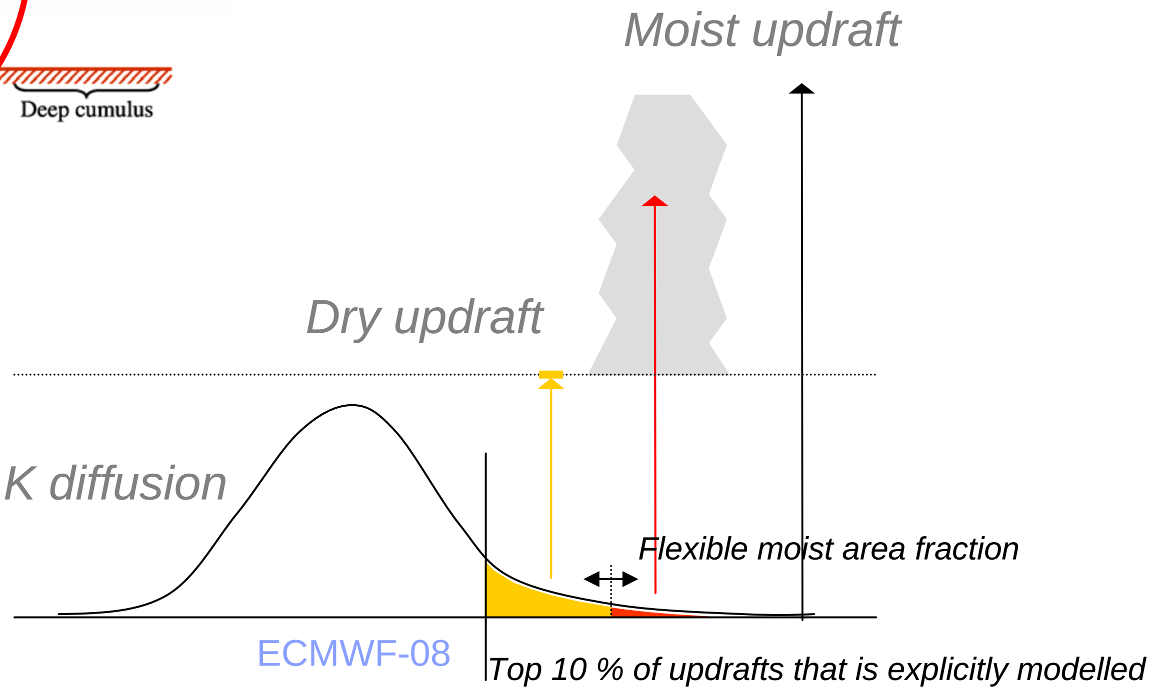
alternatives: Lappen and Randall JAS 2001

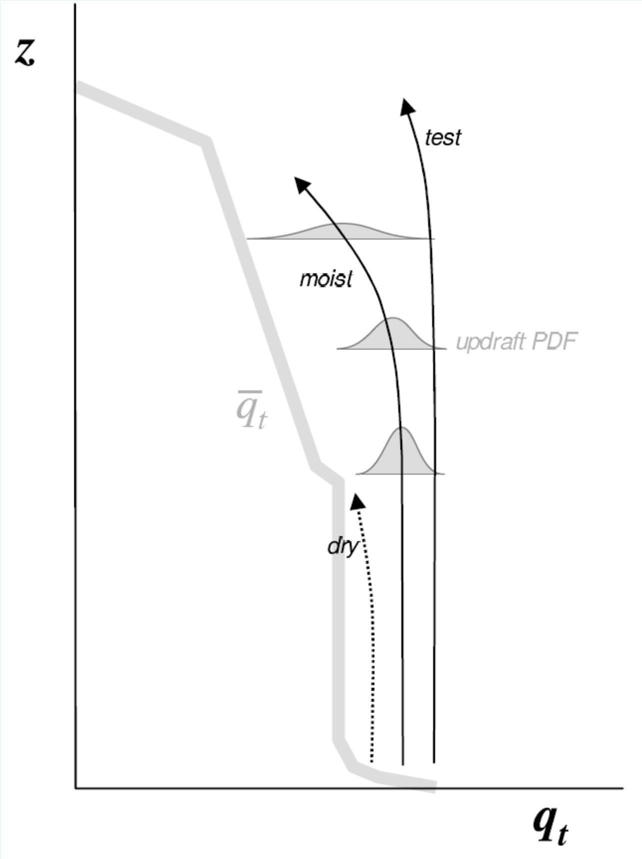
Rio and Hourdin JAS 2008

Figure courtesy of Martin Koehler



$$\overline{w'\phi'}_{PBL} = -K\frac{\partial\overline{\phi}}{\partial z} + \sum_{i=1}^{N} M_i(\phi_i - \overline{\phi})$$





- Assume a Gaussian joint PDF(θl,qt,w) shape for the cloudy updraft.

- Mean and width determined by the multiple updrafts

- Determine everything consistently from this joint PDF

$$a, w_u, \theta_{l,u}, q_{t,u}$$

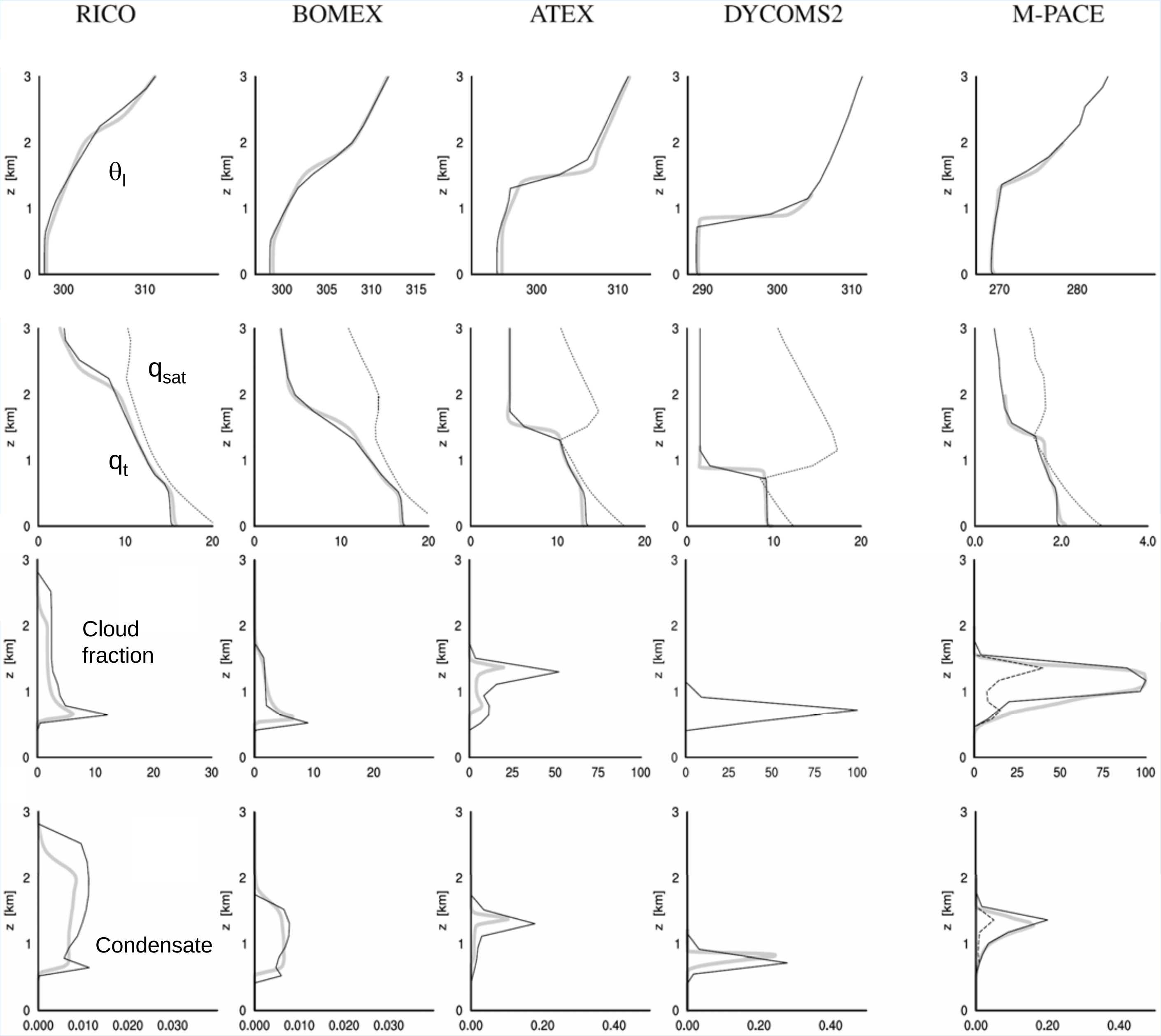
An reconstruct the flux:

$$\overline{w'\psi'} = a_u w_u \left(\psi_u - \overline{\psi}\right)$$

Remarks:

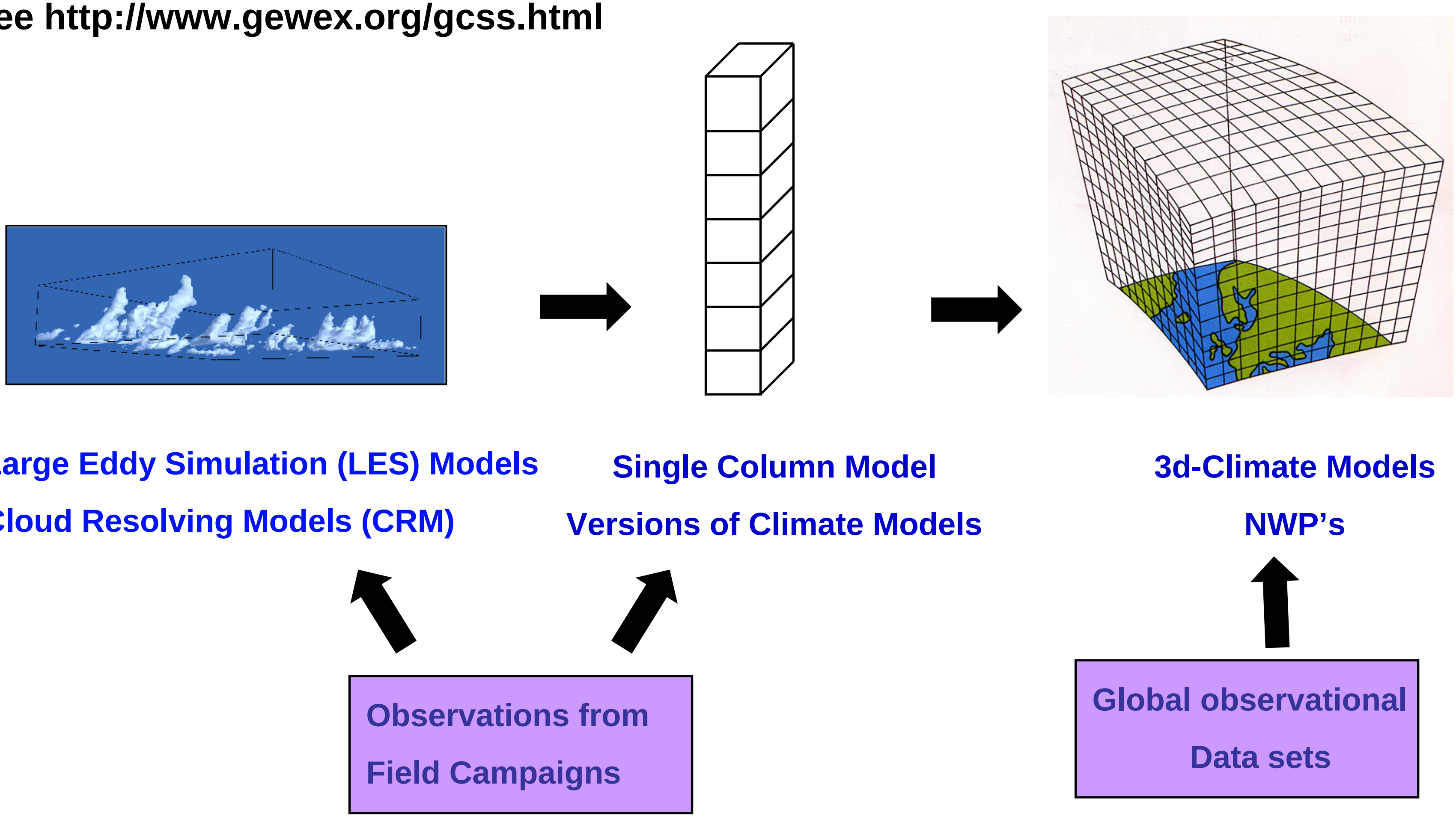
- No closure at cloud base
- No detrainment parameterization
- Pdf can be used for cloud scheme and radiation

Tested for a large number of GCSS Cases……………….



A slow, but rewarding Working Strategy

See http://www.gewex.org/gcss.html



Large Eddy Simulation (LES) Models

Cloud Resolving Models (CRM)

Single Column Model

Versions of Climate Models

3d-Climate Models

NWP's

Observations from

Field Campaigns

Global observational

Data sets

Development

Testing

Evaluation

But… Many open problems remain

Conceptually on process basis

- Convective Momentum Transport
- Influence of Aerosols/Precipitation on the (thermo)dynamics of Scu and Cu
- Mesoscale structures in Scu and Shallow Cu
- Transition from shallow to deep convection (deep convective diurnal cycle in tropics)
- What controls the low cloud fraction

Parameterization

- Vertical velocity in convective clouds
- Convection on the 1km~10km scale. (stochastic convection)
- Microphysicis (precip)
- Transition regimes.

Climate

Determine and understand the processes that are responsable for the uncertainty in cloud-climate feedback.