

# Assimilation of cloud/precipitation data at regional scales

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## Introduction

### • The Model

- Weather Research and Forecasting (WRF) community model.
- Advanced Research WRF (ARW) non-hydrostatic dynamical core

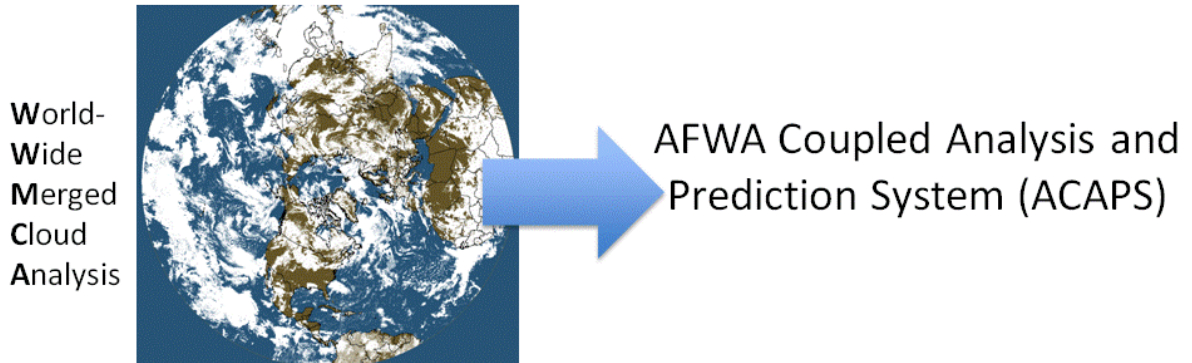
### • The Data Assimilation (DA)

- Data Assimilation Research Testbed (DART): EnKF
- WRF Data Assimilation (WRFDA): 3DVar, FGAT, 4DVar
- Gridpoint Statistical Interpolation (GSI): 3DVar, 4DVar under development

### • The Operational Applications

- Air Force Weather Agency (AFWA)
- NOAA (“Rapid Refresh”)

# Introduction

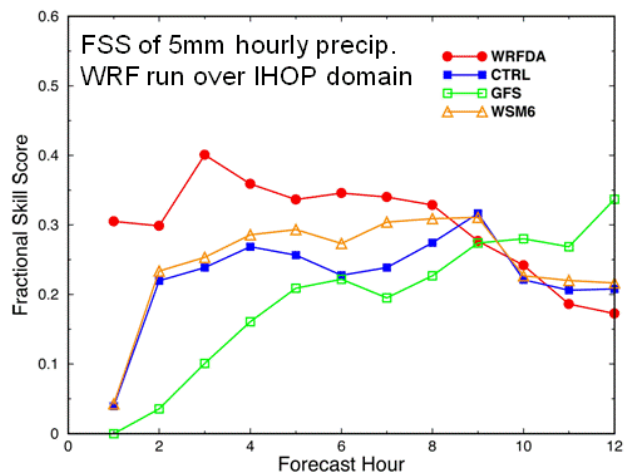
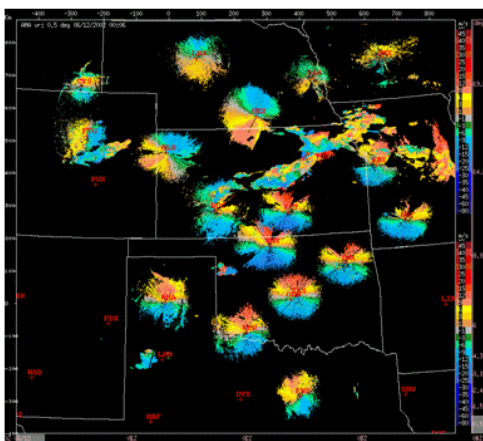


**SCOPE:** *Develop an analysis and prediction system of 3D cloud properties combined with the dynamical variables.*

## WRFDA 3DVAR and radar radial velocity IHOP one-week retrospective study

- CTRL Initialization by NCEP **ETA** analysis
- GFS Initialization by NCEP **GFS** analysis
- WRFDA WRF 3DVAR **radar radial velocity**
- WSM CTRL with **different microphysics**

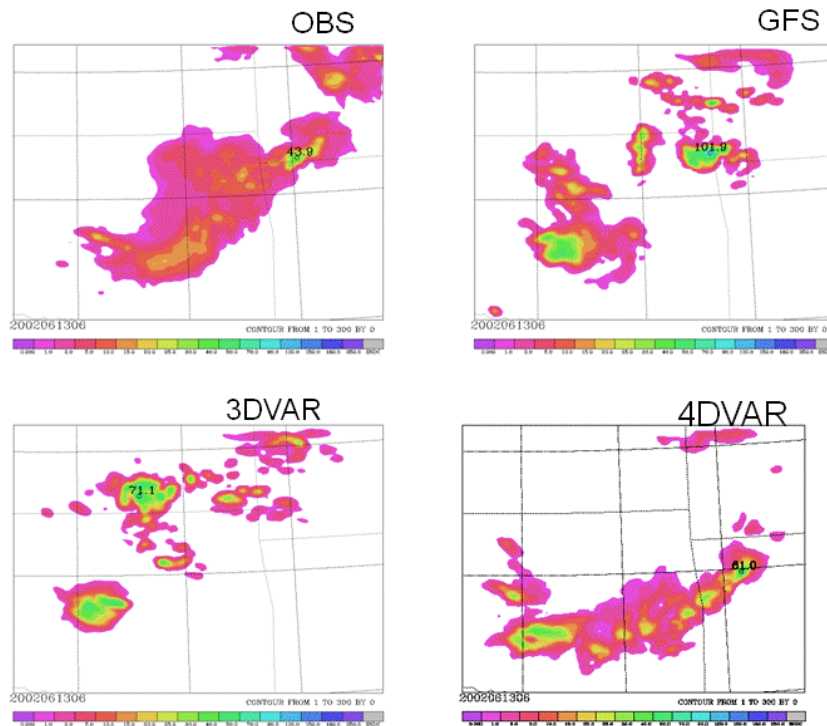
Radial velocity 2002061200 UTC



- Radar data assimilation improves the precipitation forecast up to 8 hours
- Forecast is more sensitive WRT initial conditions than physics (microphysics is most sensitive among all physics tested)

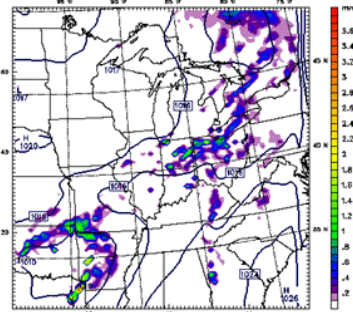
*(Jenny Sun)*

# Hourly precipitation at 0600 UTC 13 June

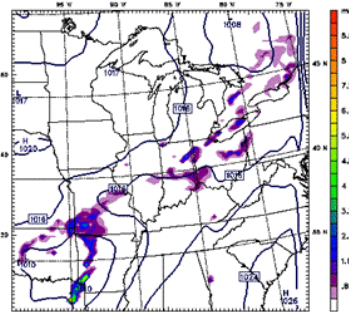


# Simulated SSMIS radiances

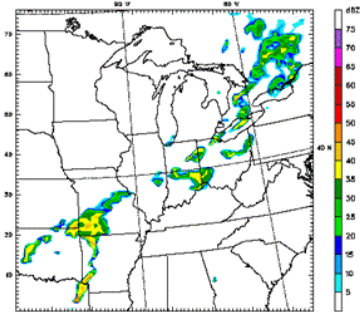
Column-Integrated cloud water



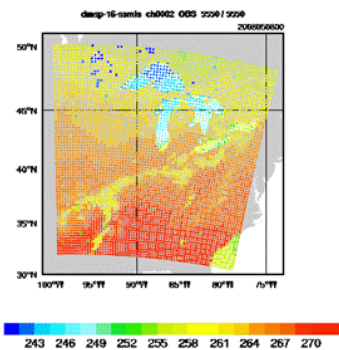
Column-Integrated rain water



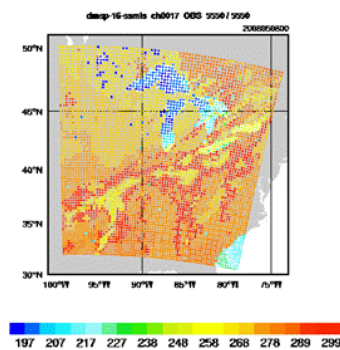
Radar Reflectivity



Simulated Ch2 Tbs



Simulated Ch17 Tbs



Model = "truth" for SSMIS radiance simulation

Only liquid hydrometeors considered

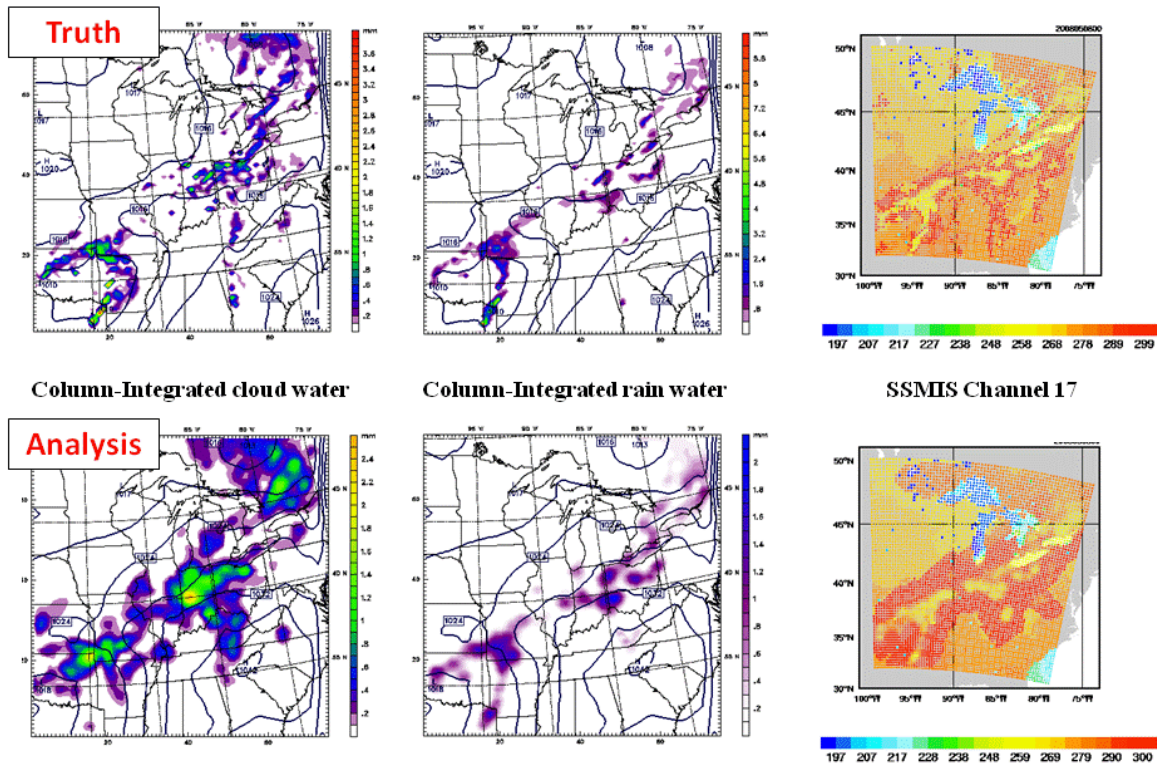
Simulated SSMIS radiances (ch 1~6, 8~18) at each grid-point using CRTM

(Zhiqun Liu)

# Assimilation of simulated SSMIS radiances in WRF 3DVAR

- Use total water  $Q_t = Q_{wv} + Q_{clw} + Q_{rain}$  as a control variable (instead of individual hydrometeors)
- Use a warm-rain microphysics scheme's TL&AD for partitioning  $Q_t$  increment into  $Q_{wv}$ ,  $Q_{clw}$  &  $Q_{rain}$ . (Xiao et al., 2007)
- CRTM as cloudy radiance observation operator
- Minimization starts from a cloud-free background, this scenario can be realistic for less accurate cloud/precip. forecast in the real world
- Perfect background for other variables (T,Q etc.)
- Perfect observations (no noise added to the simulated Tbs)
- 2 outer-loops

## Simulated SSMIS radiances

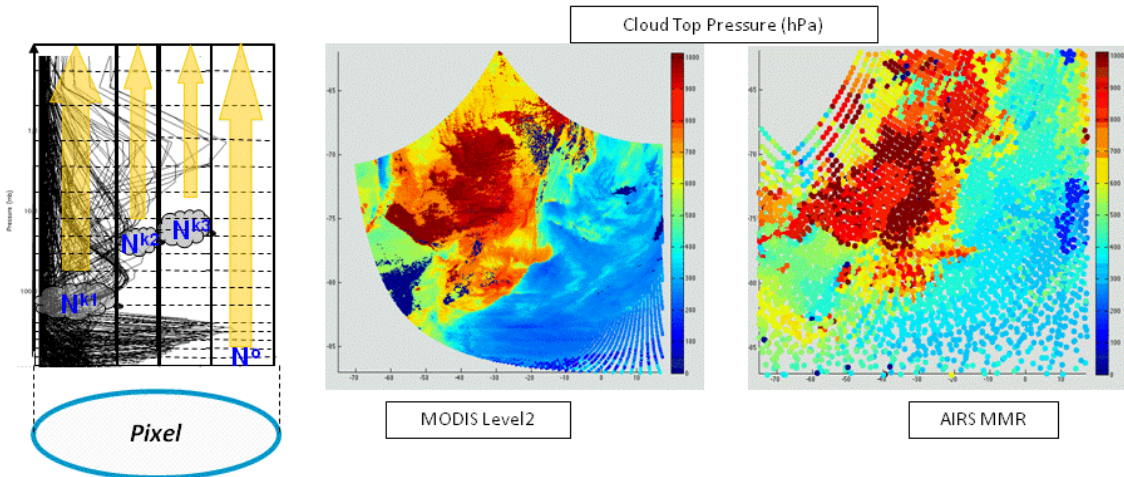


# Infrared Cloudy Radiances

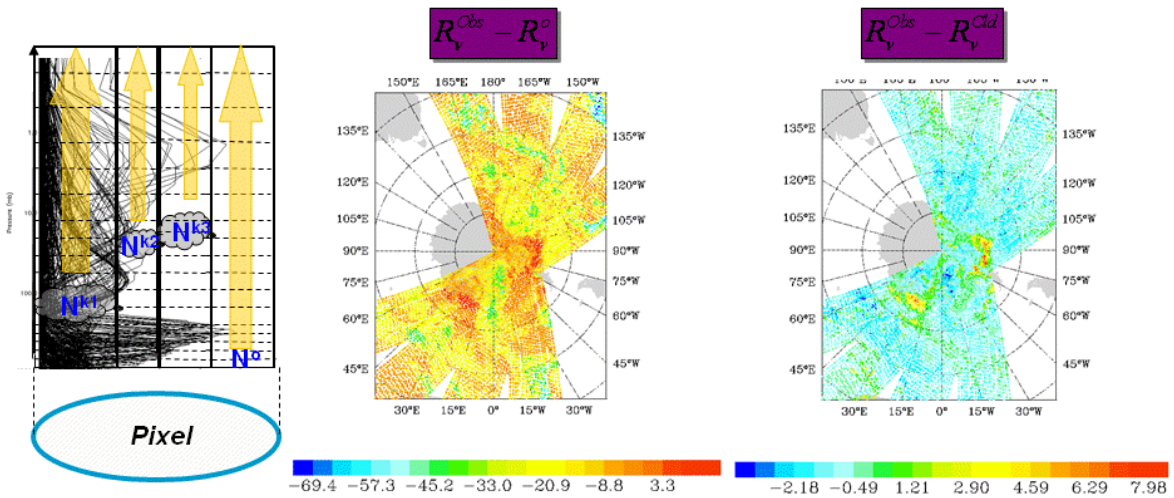
Cloud fractions  $N^k$  are adjusted *variationally* to fit observations:

$$R_v^{Cld} = N^o R_v^o + \sum_{k=1}^n N^k R_v^{*k}$$

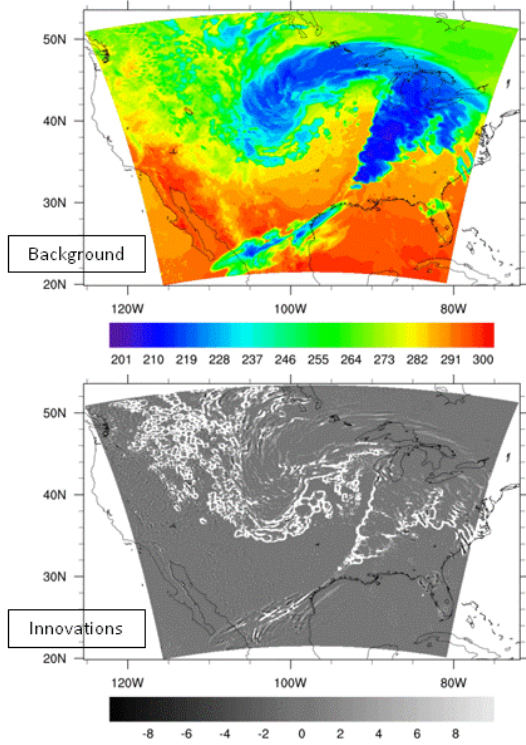
$$J(N) = \frac{1}{2} \sum_v \left( \frac{R_v^{Cld} - R_v^{Obs}}{R_v^o} \right)^2 \text{ with } \begin{cases} 0 \leq N^k \leq 1, \forall k \in [0, n] \\ N^o + \sum_{k=1}^n N^k = 1 \end{cases}$$



## IR Cloudy Radiances (linear obs operator)

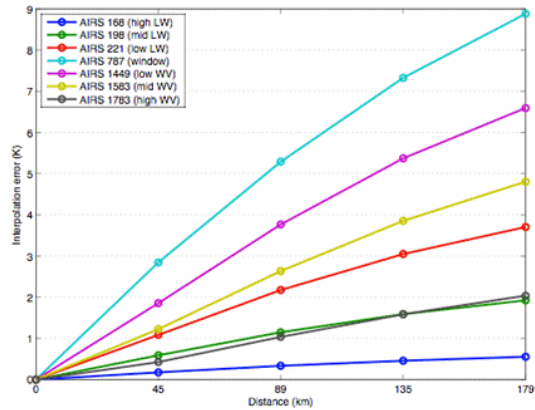


# Representativeness Error



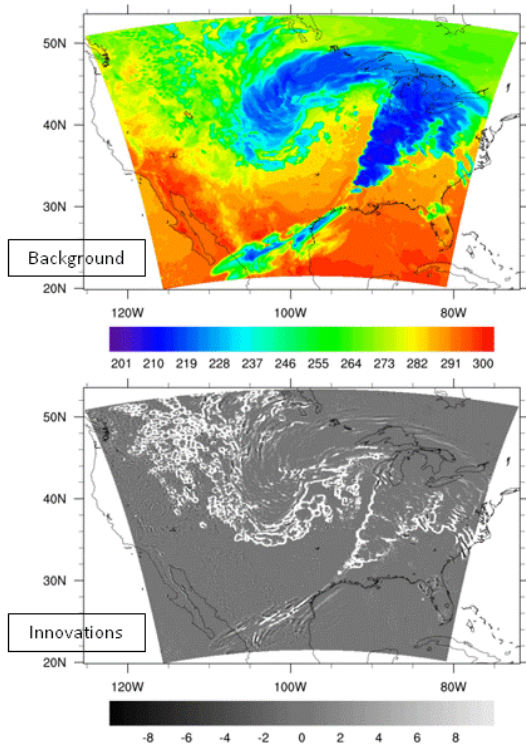
Simulated mismatch in resolution:

- Perfect observations (high resolution)
- Perfect Background (lower resolution)



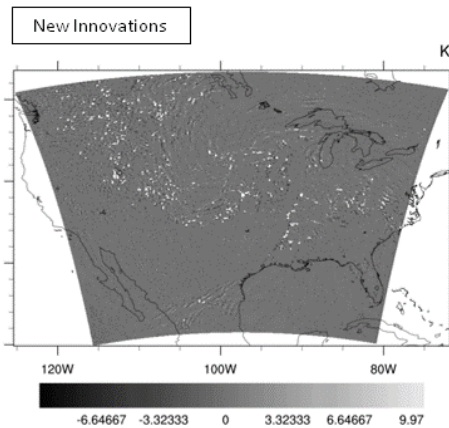
11

# Representativeness Error

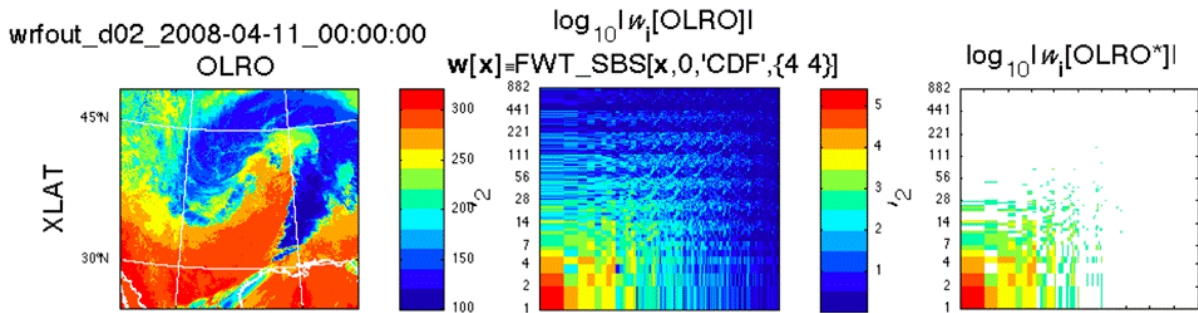


New interpolation scheme:

1. Automatic detection of sharp gradients
2. New "proximity" for interpolation



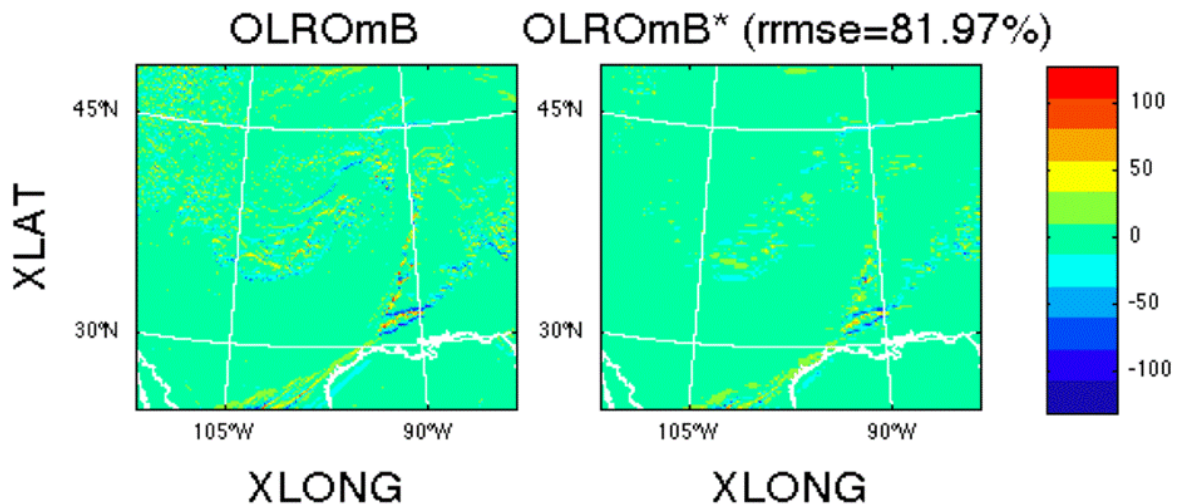
## Biorthogonal wavelet transform can *isolate* observation-background differences scale-by-scale while preserving physical-space localization



By sorting and comparing  $|w_i|$  for obs.  $\mathbf{y}_o$  & background  $\mathbf{y}_b$  we can isolate a multi-scale subset  $i \in I$  (right) from which *equivalent* representations  $\mathbf{y}_o^*$  and  $\mathbf{y}_b^*$  of  $\mathbf{y}_o$  and  $\mathbf{y}_b$  can be reconstructed...

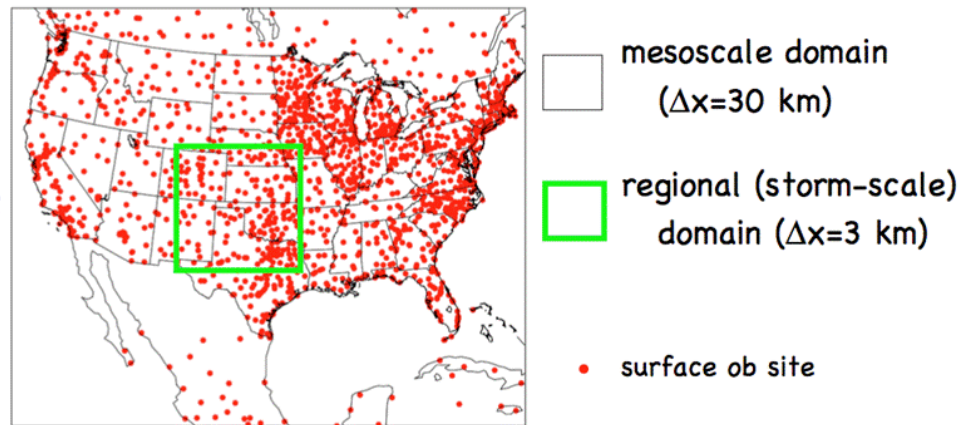
(Aimé Fournier)

## Reduction in representativeness error within observation-background differences



The raw  $\mathbf{y}_o - \mathbf{y}_b$  (left) includes errors due to  $\mathbf{y}_o$  and  $\mathbf{y}_b$  coming from completely different representations, that (hypothetically) have been *reconciled* by the foregoing wavelet-coefficient selection procedure.

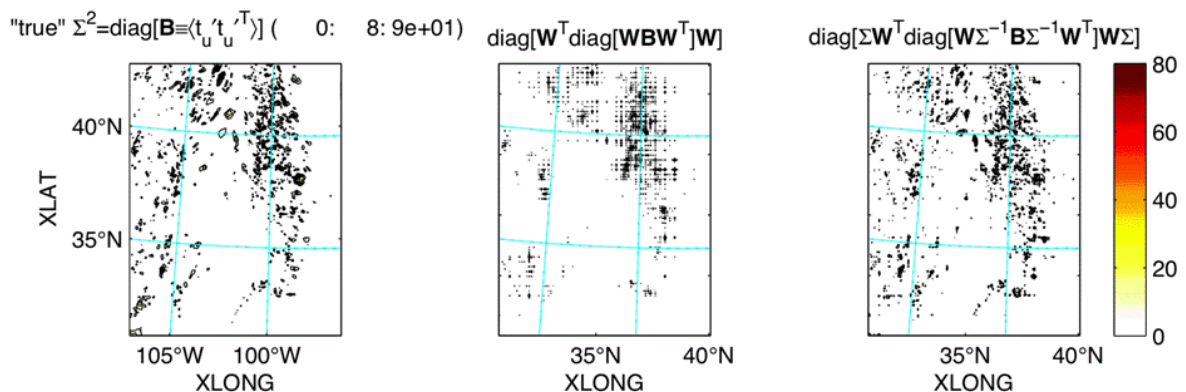
WRF  
ensemble



- 30-km ensembles are initialized at 1200 UTC on the day of interest.
  - Multi-physics ensemble
    - PBL (3 schemes), cumulus (3 schemes), shortwave radiation (2 schemes)
  - Ensemble mean and boundary conditions from 1200 UTC NAM
  - Spatial perturbations from an ensemble Kalman filter applied to observations (sounding, surface, aircraft) during the previous 2.5 days
    - Perturbations are mesoscale and flow dependent.
    - Grid-scale dynamics and parameterization diversity increase ensemble spread.
    - Observations decrease ensemble spread.
- Each 30-km ensemble member provides initial and boundary conditions for a 3-km ensemble member. (David Dowell)

## Wavelet representation of Background Error Covariance Matrix

Background covariance can be *efficiently* modeled by assuming diagonality of the wavelet-coefficient covariance matrix (Fisher & Andersson, Deckmyn & Berre).

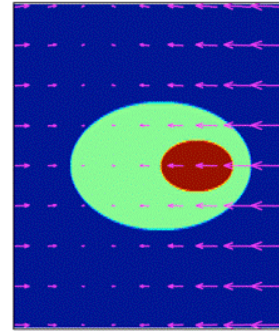


- The normalization with  $\Sigma^2 = \text{diag} \mathbf{B}$  (left) yields a model with *fewer* artifacts (right) than does  $\Sigma = \mathbf{I}$  (center) (as found by D&B earlier).
- In these plots  $\mathbf{x}$  is unbalanced temperature anomaly in a 30-member ensemble computed by Dowell with horizontal resolution  $N=450 \times 350$ . (Aimé Fournier)



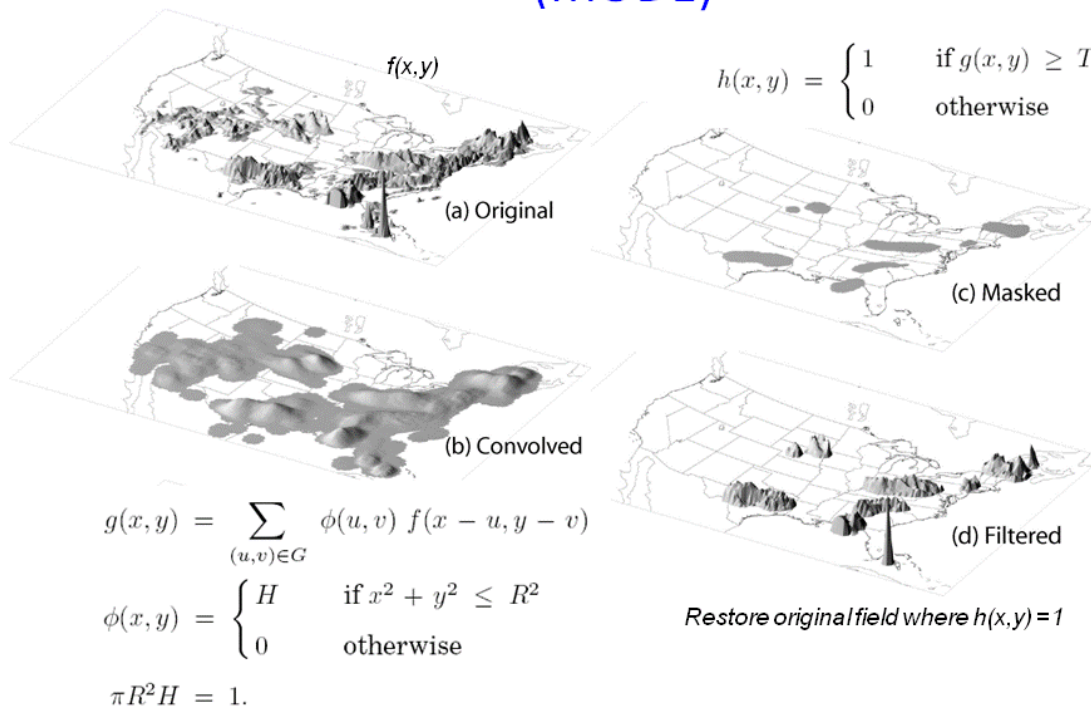
## Diagnostic Verification Methods

1. Object-based
2. Field Deformation
3. Neighborhood (fuzzy)
4. Scale Decomposition
5. Variograms

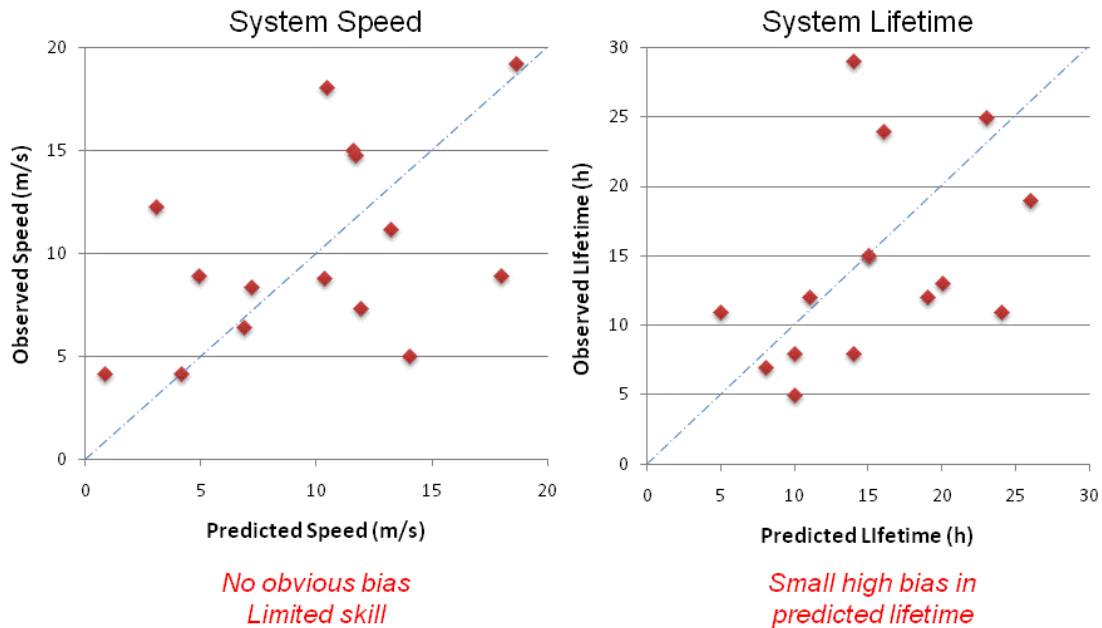


(Chris Davis - <http://www.ral.ucar.edu/projects/icp>)

## Method of Object-based Diagnostic Evaluation (MODE)



## Attributes of Rain Systems



ICP Workshop, August 24-25, 2009

19

## Recommendations

- Will traditional DA methods work for clouds? (e.g. non-linear, non-Gaussian)
- Focus on model error (e.g. microphysics, RTM)
  - info on model deficiencies
  - new DA techniques
- Leverage Ensemble / Variational experience
- Modular codes
  - increase flexibility
  - facilitate collaborations